



Tropical cyclogenesis

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RA I WMO - Training course on tropical cyclones 2023

Tropical cyclogenesis

Mechanisms, basin pattern and equatorial waves



**METEO
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I) Basic concepts and pathways to genesis

- *Definition*
- *Environmental conditions*
- *Historical theories*
- *The need for a vorticity precursor*
- *Interactions with wind shear*

II) Basin configurations

- 1) *Monsoon trough*
- 2) *Near equatorial trough*
- 3) *Climatology*
- 4) *Dynamics of the MT and NET*

III) Equatorial waves

- 1) *Definition and presentation*
- 2) *Link with TCG*

IV) A few advanced notions

- 1) *The Marsupial Paradigm*
- 2) *Vortical Hot Towers route to TCG*
- 3) *TCG at the intersection of large- and meso-scales*

Tropical cyclogenesis

Basic concepts and pathways to genesis



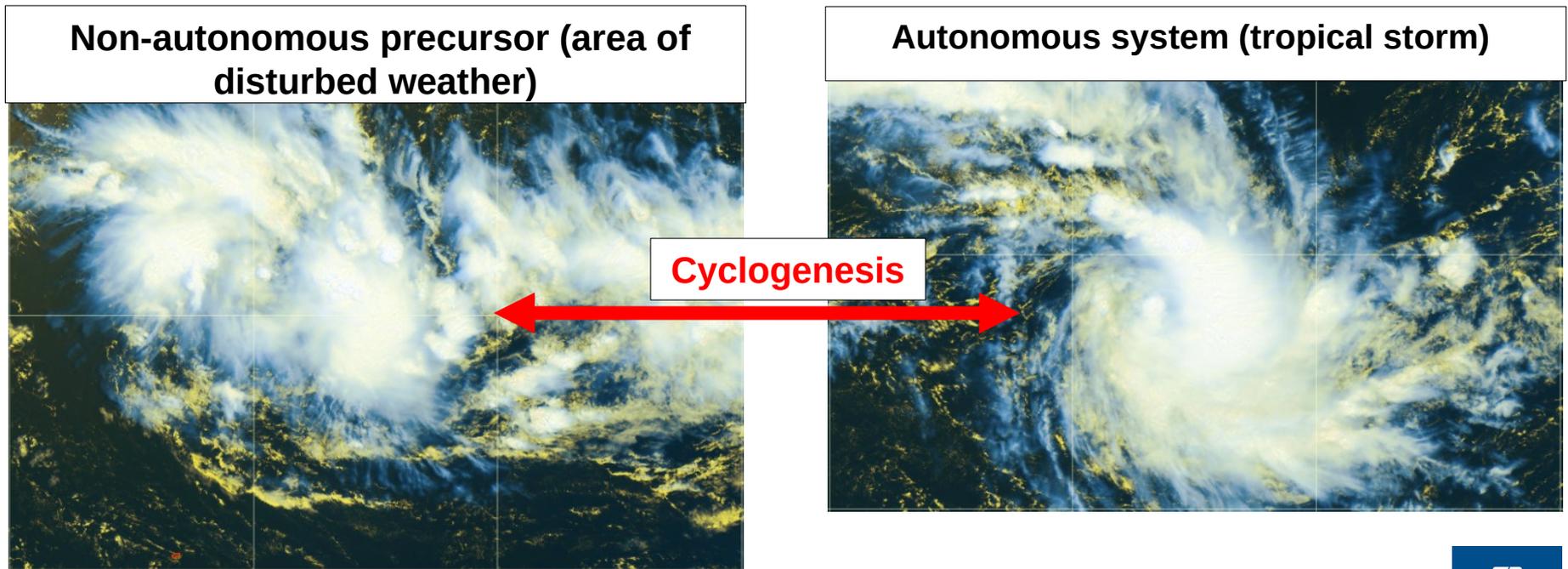
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FRANCE**



Cyclogenesis : definition

Definition : A tropical cyclogenesis (TCG) is carried out when a low pressure area has become an **autonomous** system. It no longer needs help from its environment to develop, through "**environmental forcing**".

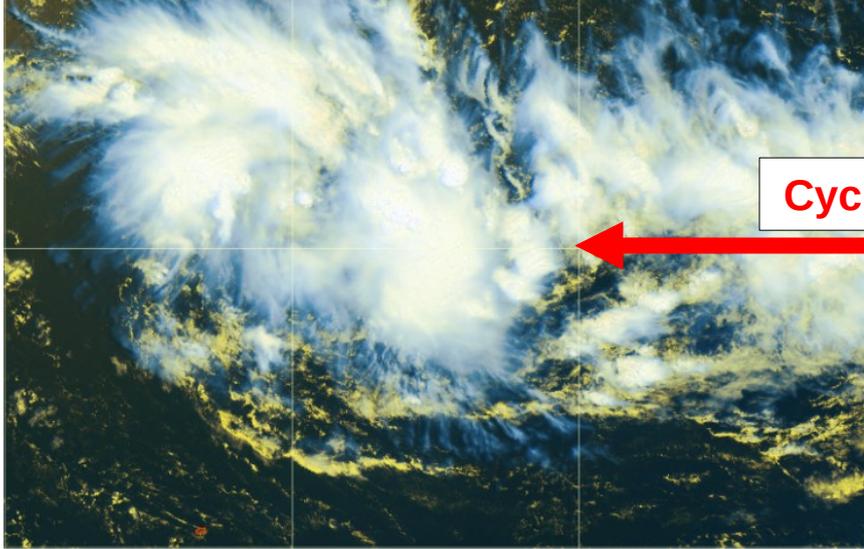
In operational terms, cyclogenesis is carried out when the stage of a tropical storm is reached. This process usually takes several days.



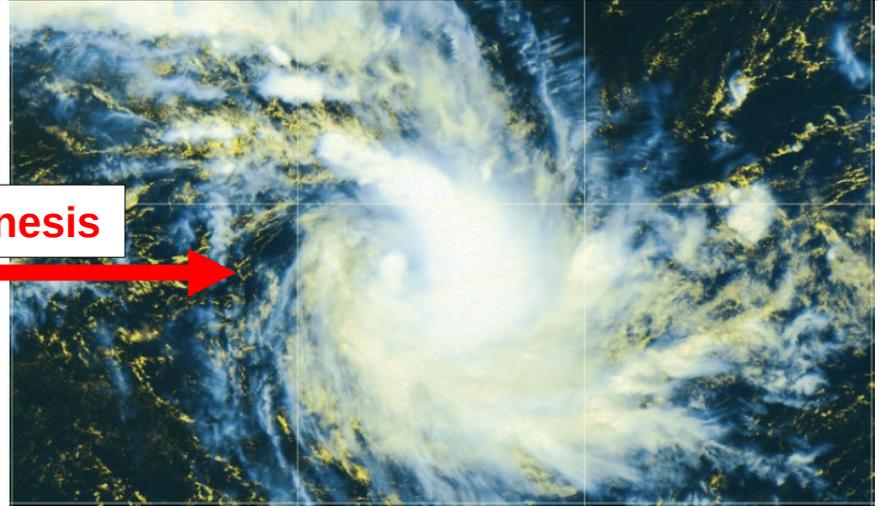
Src : MF Archives

Cyclogenesis : definition

Non-autonomous precursor (area of disturbed weather)



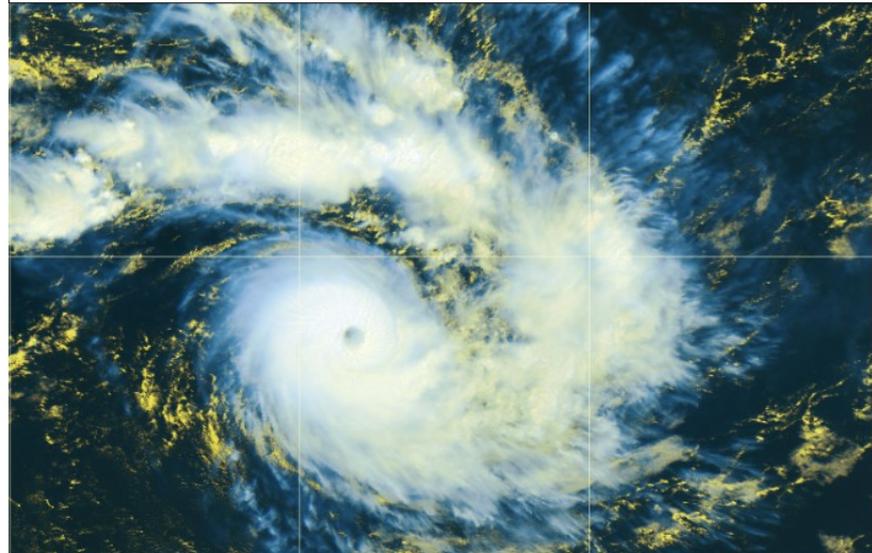
Autonomous system (tropical storm)



Cyclogenesis



Mature autonomous system



Intensification



Src : MF Archives

Cyclogenesis : basic concepts

Necessary conditions for cyclogenesis (Gray, 1968) :

- ✓ Sufficient **ocean energy** [Sea Temp. > **26°C** over at least 60 m depth]
- ✓ Generalized instability allowing deep convection
- ✓ Mid-tropospheric humidity (700/400 hPa layer)
- ✓ Latitude > **5°**
- ✓ Low vertical wind shear (**less than 15kt**)
- ✓ Vorticity of low layers (precursor)

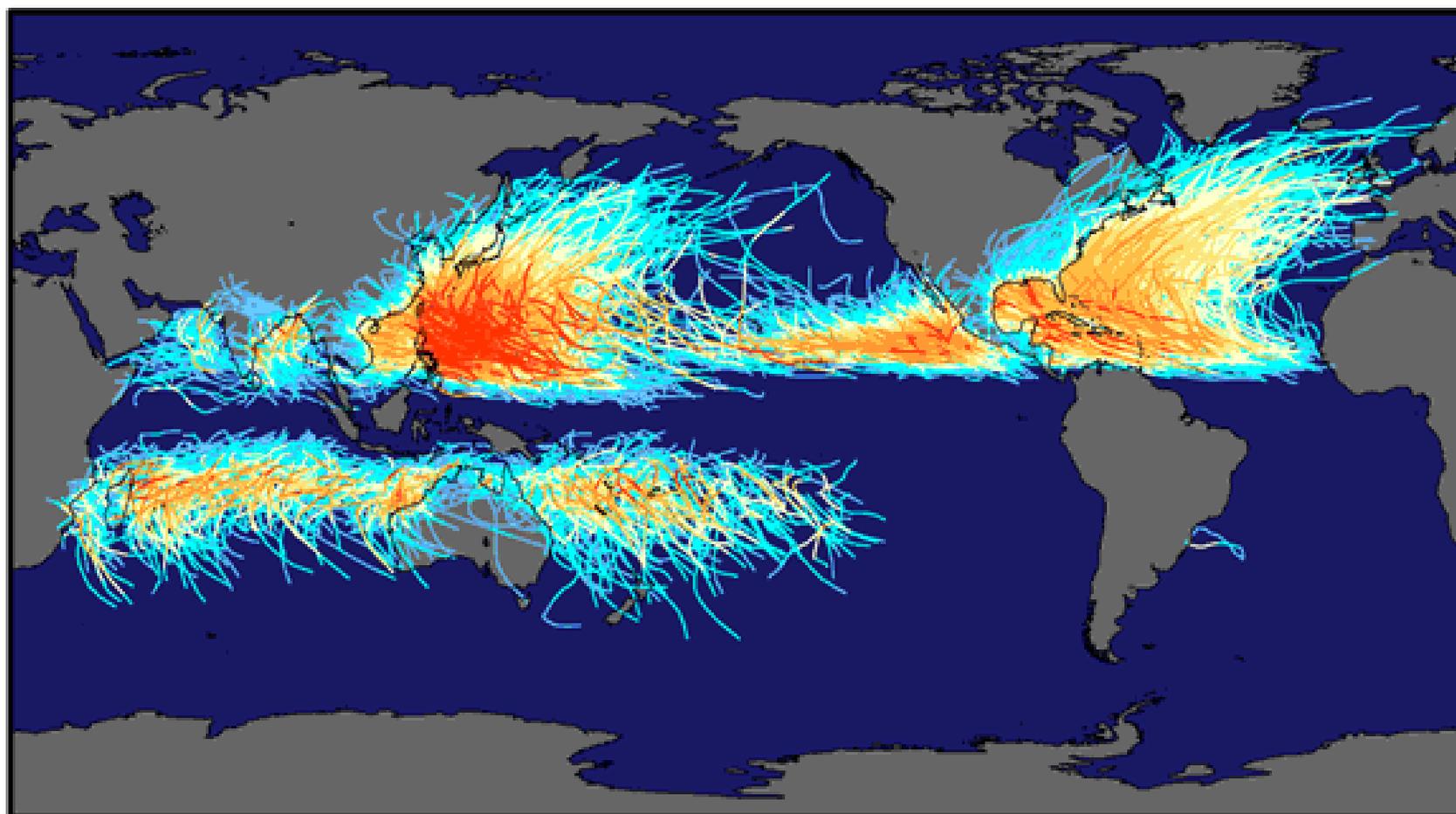
Cyclogenesis : not that simple...

Necessary conditions for cyclogenesis (Gray, 1968) :

- ✓ Sufficient **ocean energy** [Sea Temp. > **26°C** over at least 60 m depth] } **Fuel**
- ✓ Generalized instability allowing deep convection } **Conducive atmospheric environment**
- ✓ Mid-tropospheric humidity (700/400 hPa layer) }
- ✓ Latitude > **5°** } **Conducive dynamical environment**
- ✓ Low vertical wind shear (**less than 15kt**) }
- ✓ Vorticity of low layers (precursor) } **Spark !**

Cyclogenesis : basic concepts

Tracks and Intensity of Tropical Cyclones, 1851-2006

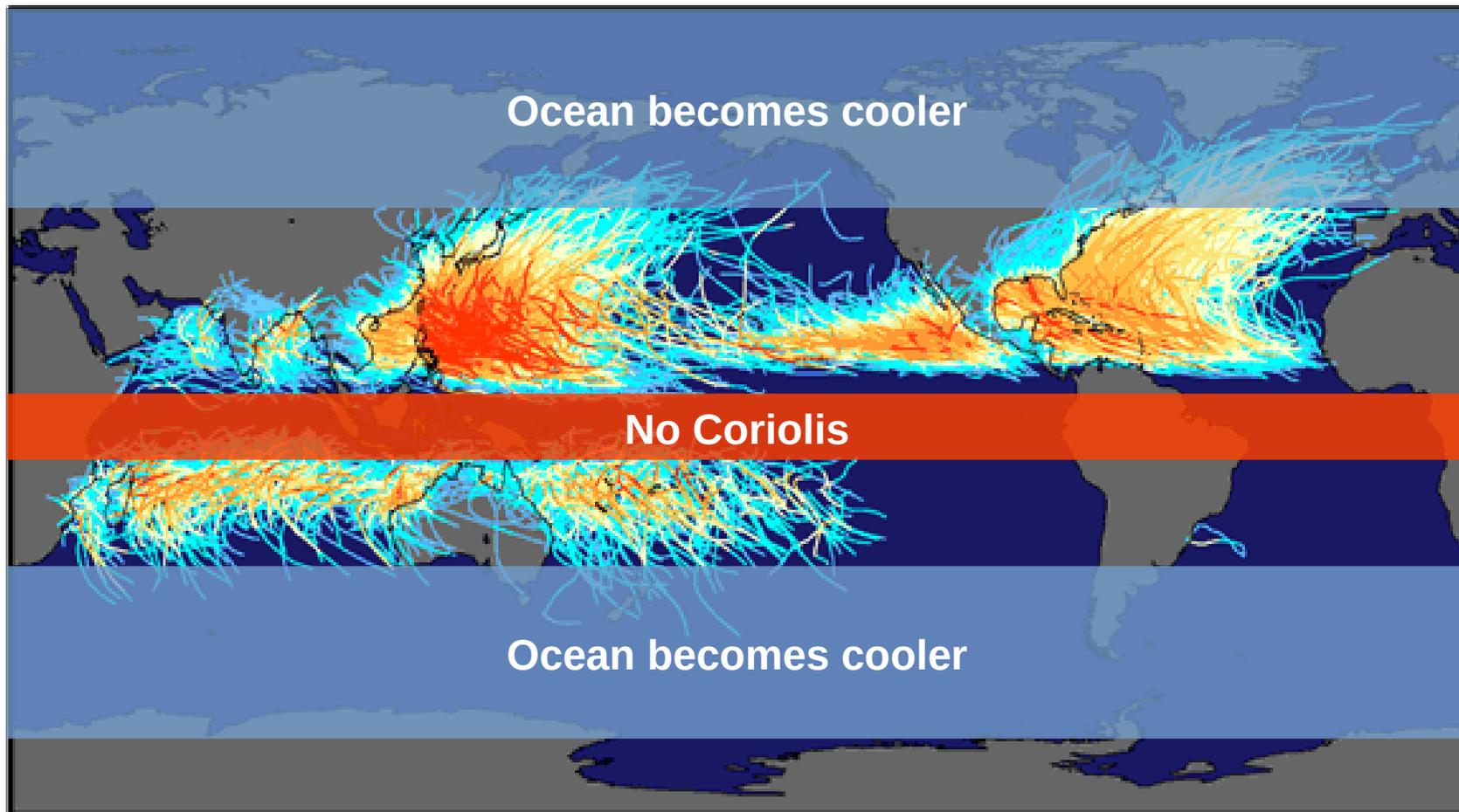


Saffir-Simpson Hurricane Intensity Scale

Robert A. Rohde, UC Berkeley / NASA's Earth Observatory

Cyclogenesis : basic concepts

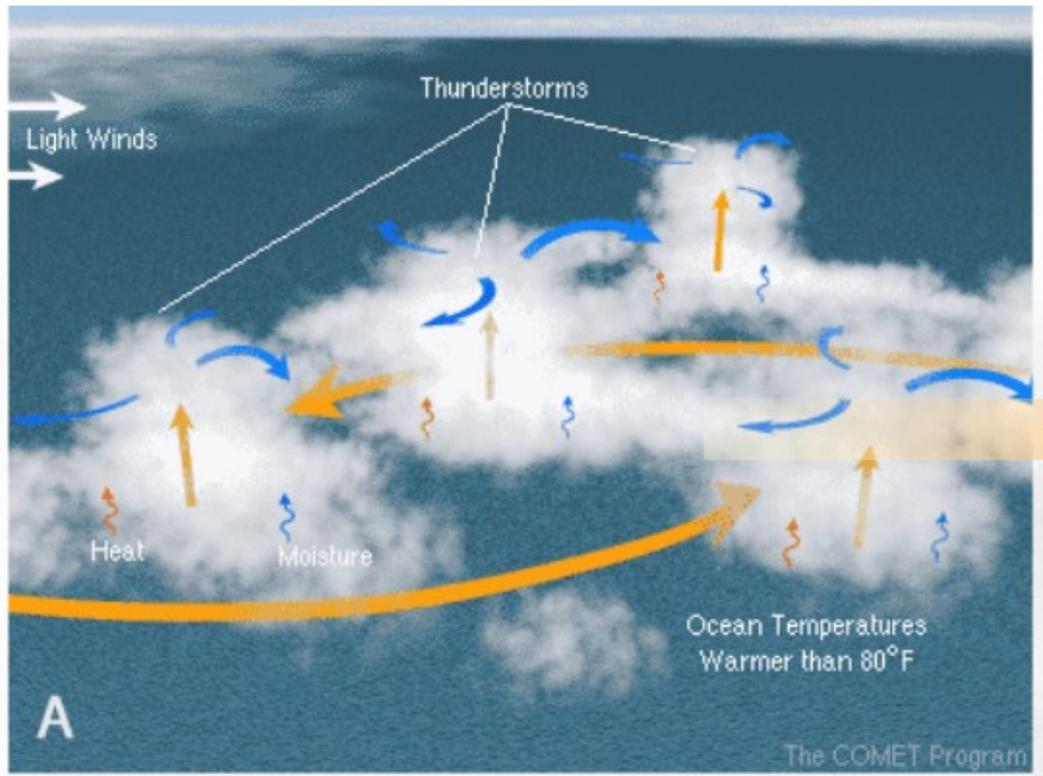
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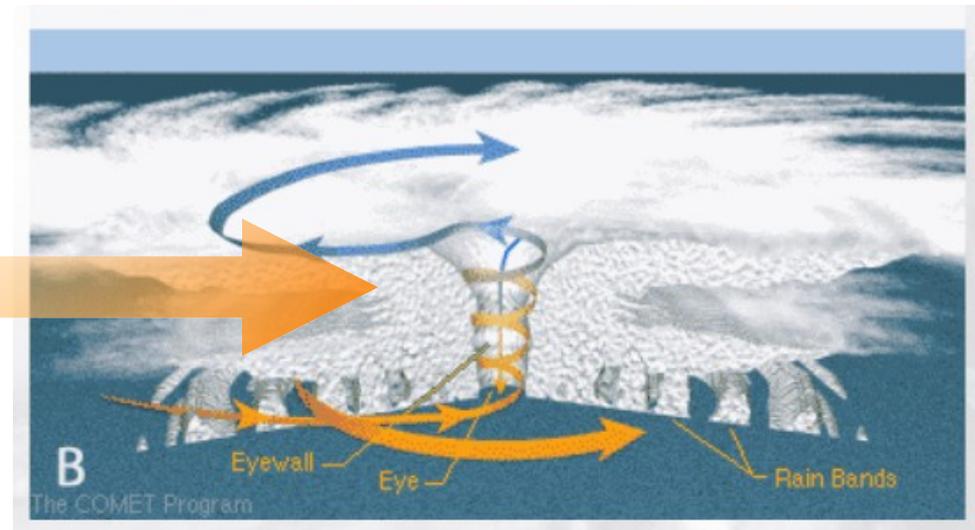
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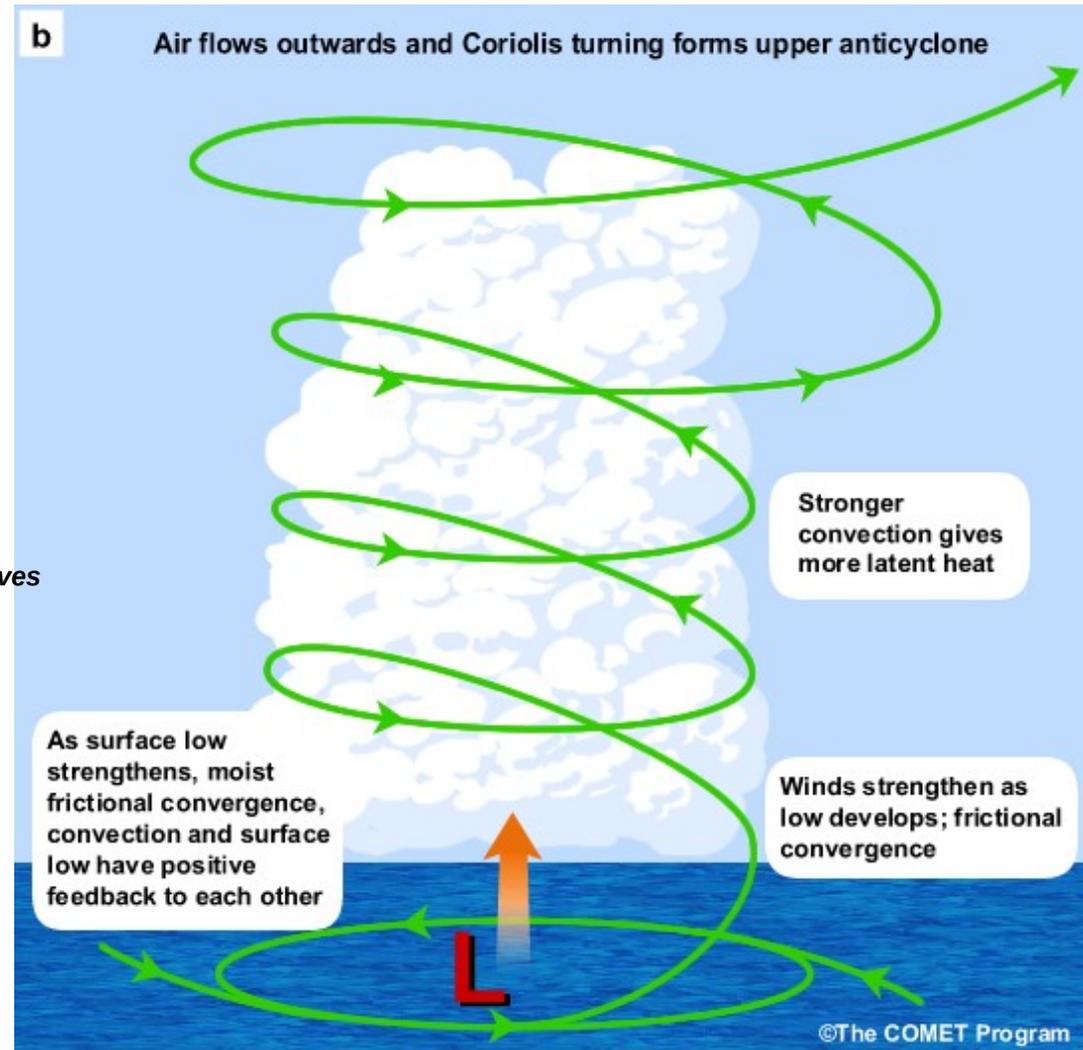
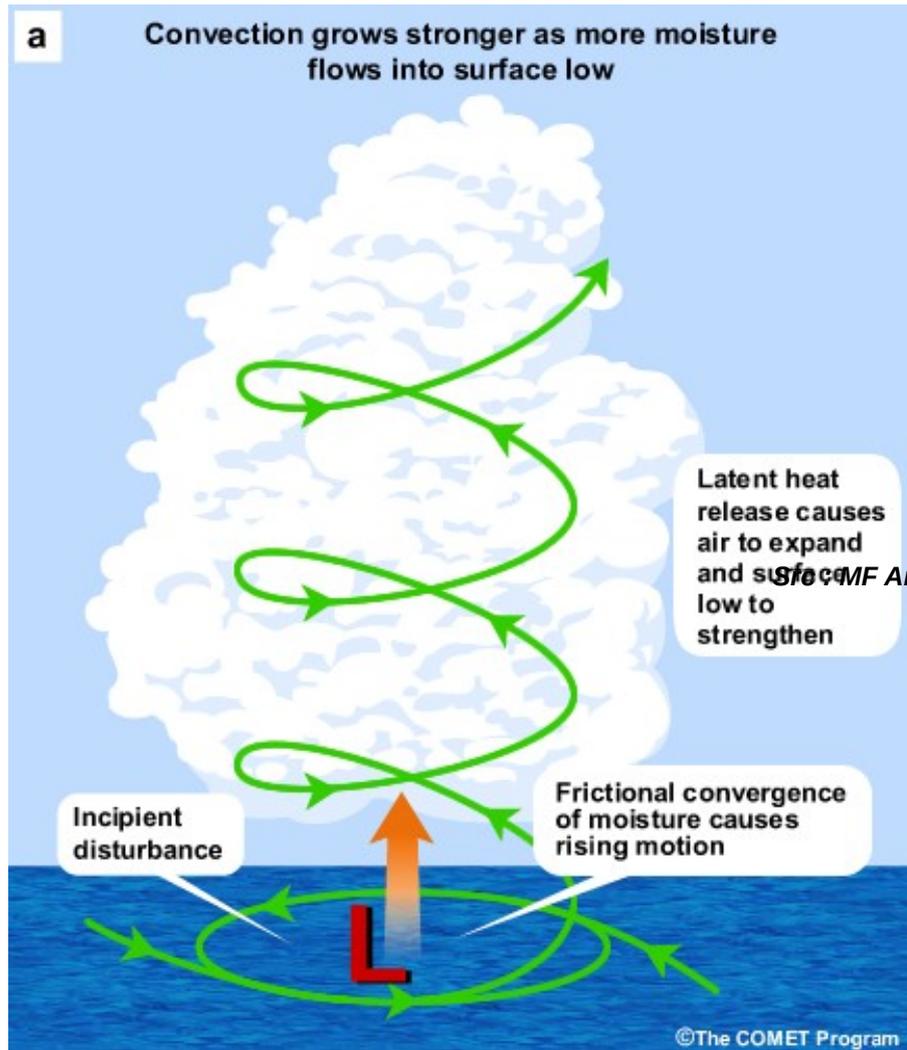
How do we go from a few MCS to a self-sustained TS ?



Src : COMET Program



How do we go from a few MCS to a self-sustained TS ?

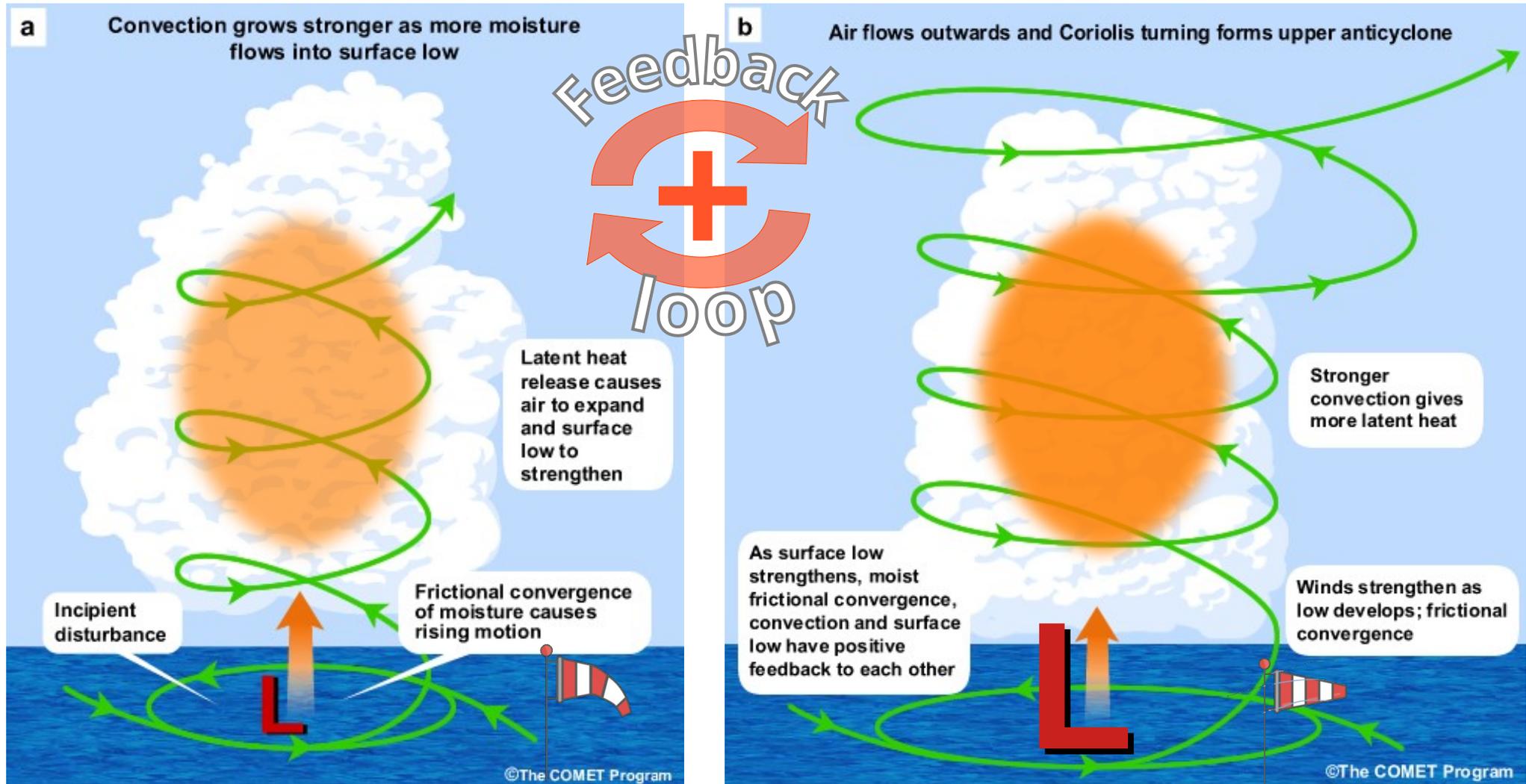


Src : COMET Program

CISK Theory (Ooyama, Charney & Eliassen, 1964) :

- One of the first description of a pathway to genesis
- Simple scheme with an axisymmetric vortex

How do we go from a few MCS to a self-sustained TS ?

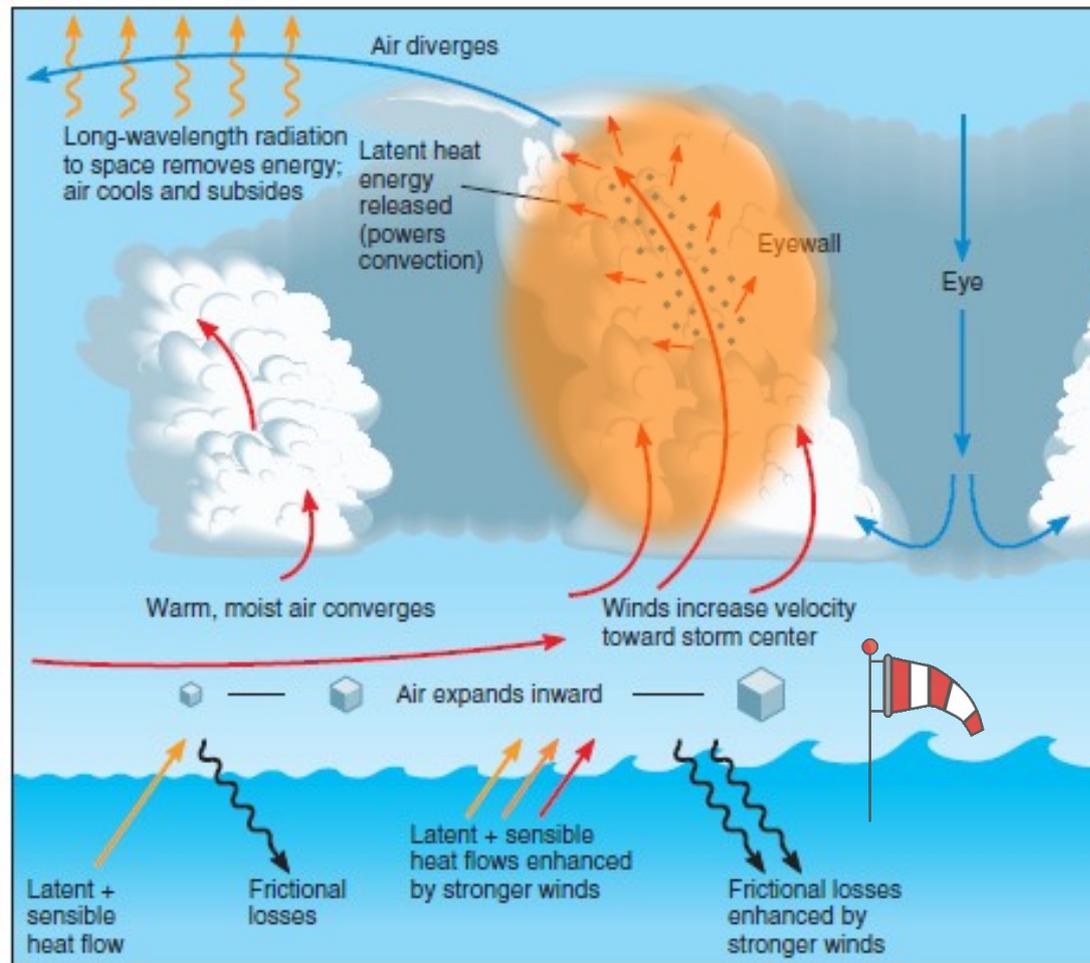


Src : COMET Program

CISK Theory (Ooyama, Charney & Eliassen, 1964) :

- One of the first description of a pathway to genesis
- Simple scheme with a simple 2-D axisymmetric vortex

How do we go from a few MCS to a self-sustained TS ?



Src : Hurricane FlashCards

Figure 11.10 Energy flow within a hurricane.

WISHE Theory (Emanuel, 1986) :

- Based on the role of the heat fluxes between the ocean and the boundary layer.
- May not be as crucial in more realistic 3D simulations (and thus, in real world situations)

Cyclogenesis : not that simple...

Necessary conditions for cyclogenesis (Gray, 1968) :

- ✓ Sufficient **ocean energy** [SST > **26°C** over at least 60 m depth] } **Fuel**
- ✓ Generalized instability allowing deep convection } **Conductive atmospheric environment**
- ✓ Mid-tropospheric humidity (700/400 hPa layer) } **Conductive dynamical environment**
- ✓ Latitude > **5°**
- ✓ Low vertical wind shear (**less than 15kt**)
- ✓ Vorticity of low layers (precursor) } **Spark !**

Cyclogenesis : basic concepts

Necessary conditions for cyclogenesis (Gray, 1968) :

- ✓ Sufficient **ocean energy** (at least 60 m depth) } **Fuel**
 - ✓ Generative circulation (50 hPa) } **Conducive atmospheric environment**
 - ✓ Low-level wind shear (**less than 15kt**) } **Conducive dynamical environment**
 - ✓ Vertical shear of low layers (precursor) } **Spark !**
- ... but not sufficient !**

Cyclogenesis : basic concepts

TABLE 4. Average TCC density and genesis productivity by ocean basin (1982–2009). Values for the Indian Ocean basin (i.e., North Indian and South Indian regions) are only for the 1998–2009 period. The boldface indicates a higher than global value in Table 2 because the 1982–97 period is included.

Region	Density (TCC $\times 10^{-6}$ km $^{-2}$ yr $^{-1}$)	Genesis productivity (%)
North Atlantic	9.02	6.0
South Atlantic	2.57	0.1
Western Pacific	10.95	12.4
Eastern Pacific	11.47	6.0
South Pacific	6.59	3.8
North Indian	10.51	7.0
South Indian	5.93	8.5
Global	7.69	7.1

Hennon et al, 2012

Fraction of persistent convective clusters (>24hrs) that resulted in cyclogenesis

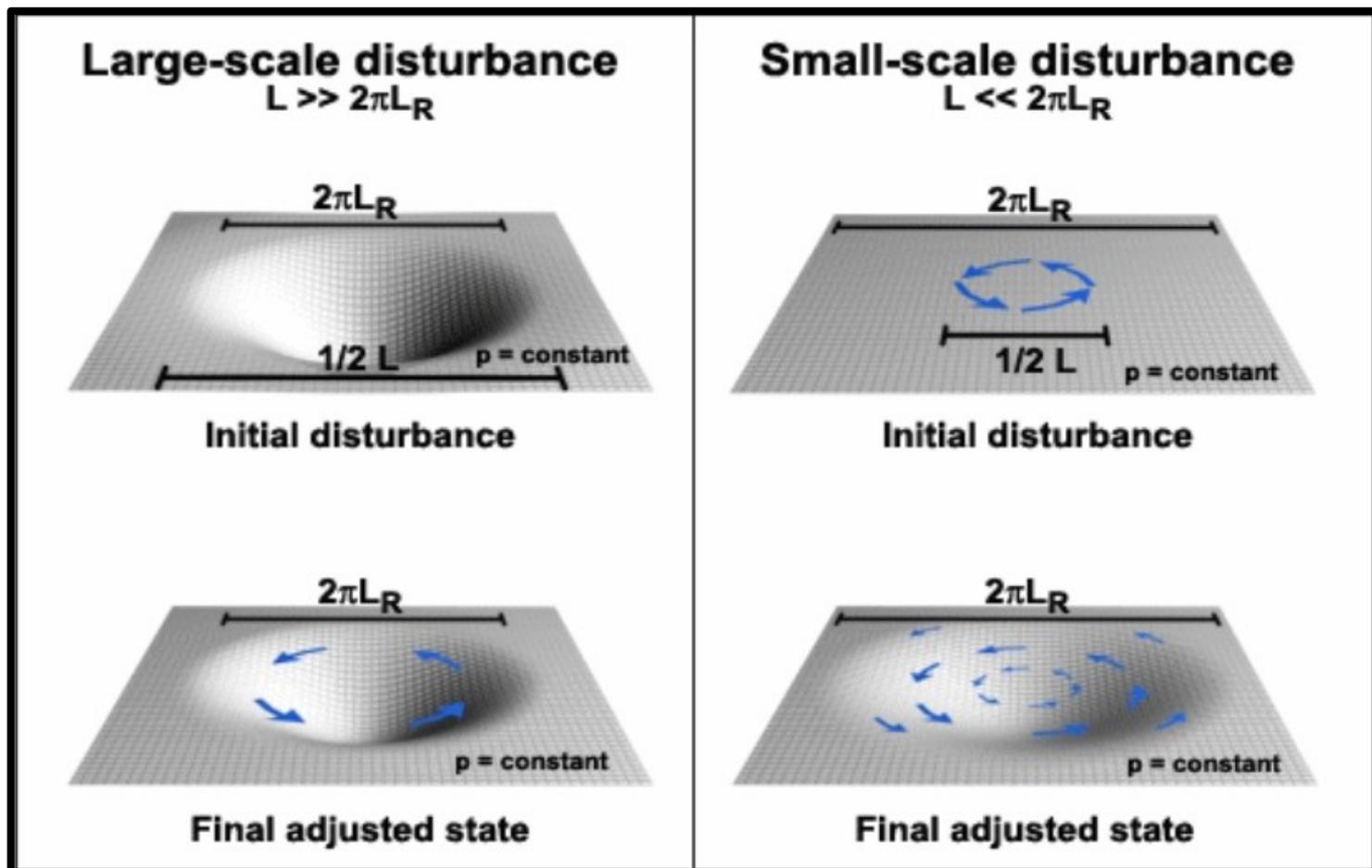
Why do we need a precursor :

The Rossby radius of deformation L_R

$$L_R = \frac{N \cdot H}{f + \xi}$$

| Buoyancy / Gravity waves
| Rotational effects

N, Brunt Väisälä freq.
H, vertical depth
f, Coriolis parameter
ξ, relative vorticity



- **Why do we need a precursor :**
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→ Defines the critical horizontal dimension beyond which a perturbation will result in a significant change in the wind and pressure fields (through geostrophic adjustment).

→ In the tropics, L_R is, by nature, very large (> 10 000 km) since the Coriolis parameter (*f*) is low ...

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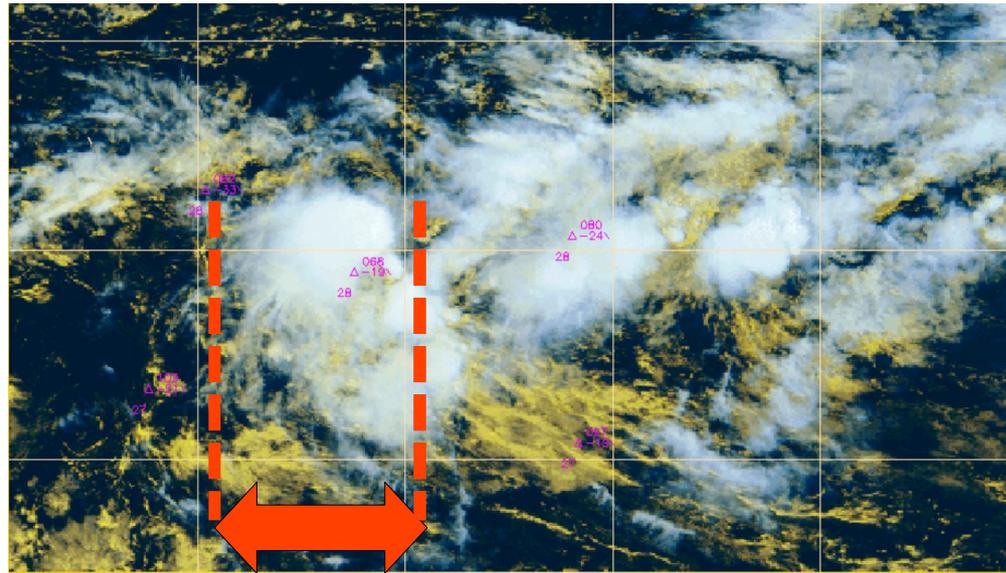
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→ In the tropics, L_R is, by nature, very large (> 10 000 km) since the Coriolis parameter (f) is low ...

So we need a significant ξ to compensate the lack of f !

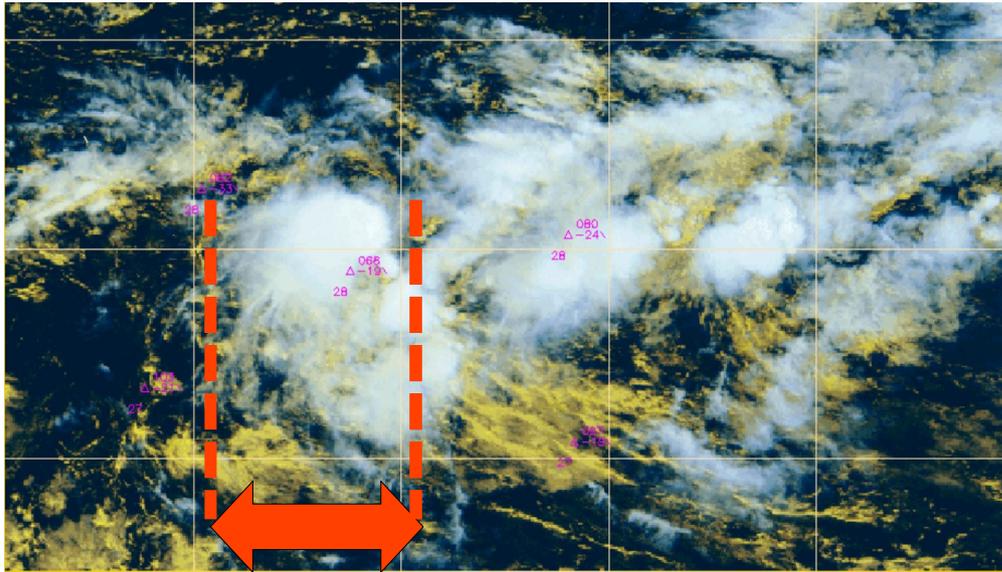
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$L < L_R$

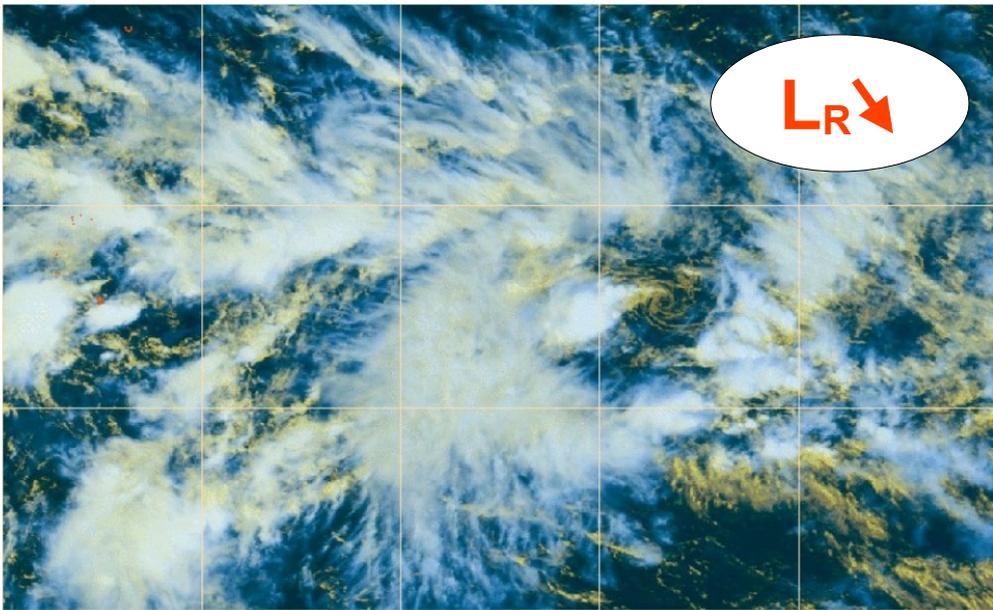
Src : MF Archives

- Why do we need a precursor :
The Rossby radius of deformation L_R

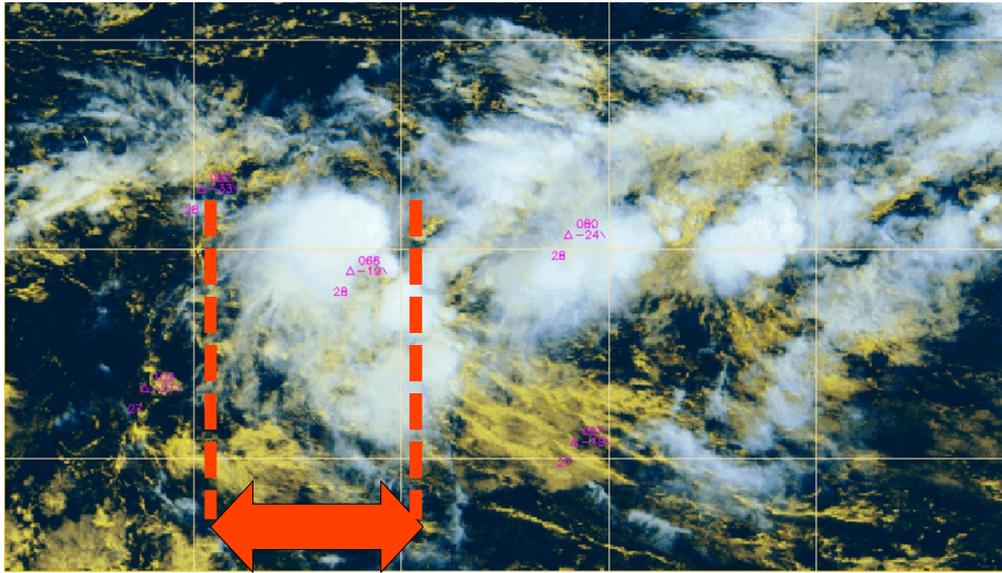


$L < L_R$

Src : MF Archives

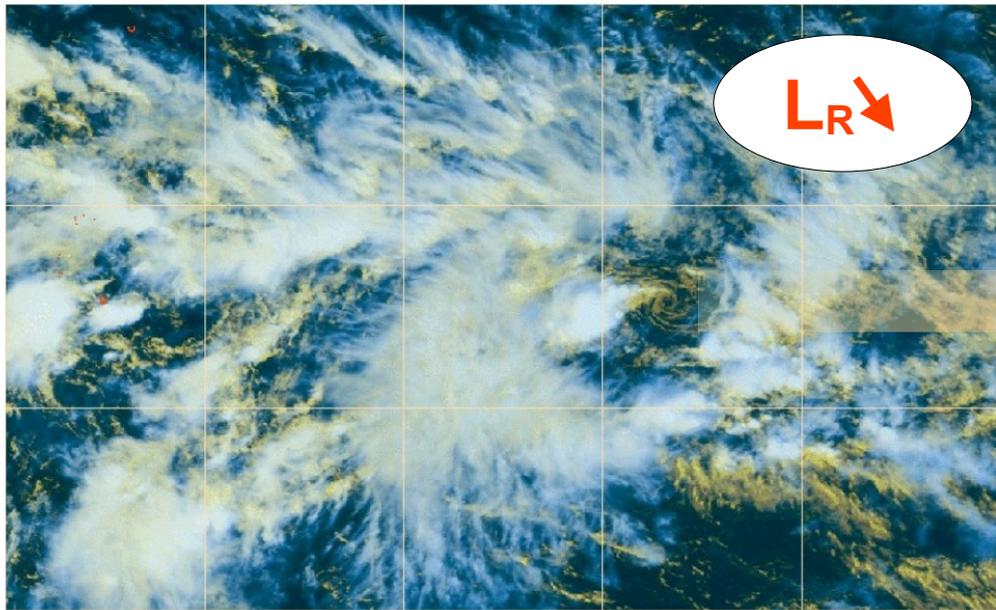


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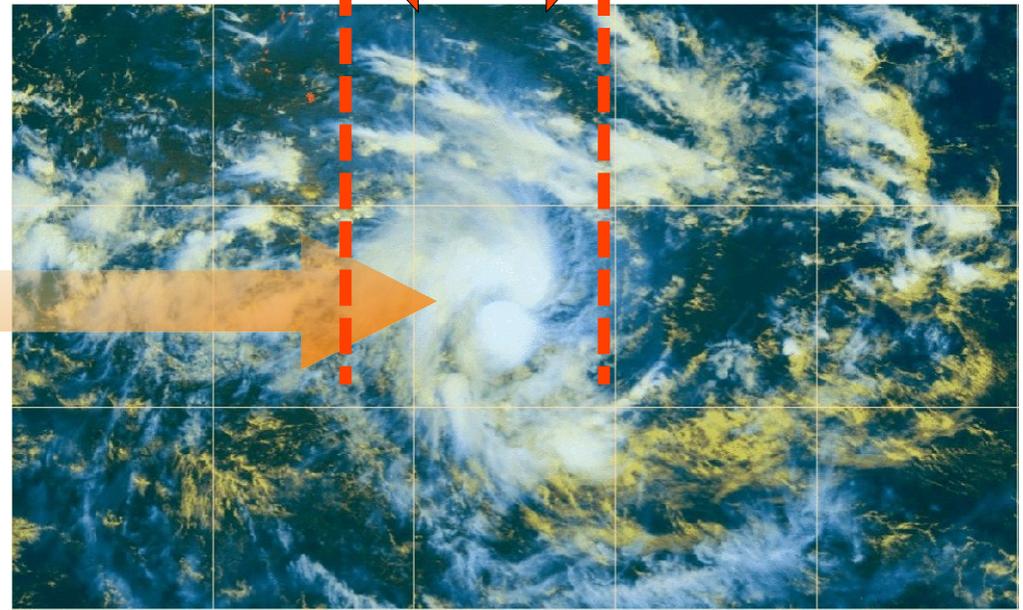


$L < L_R$

Src : MF Archives

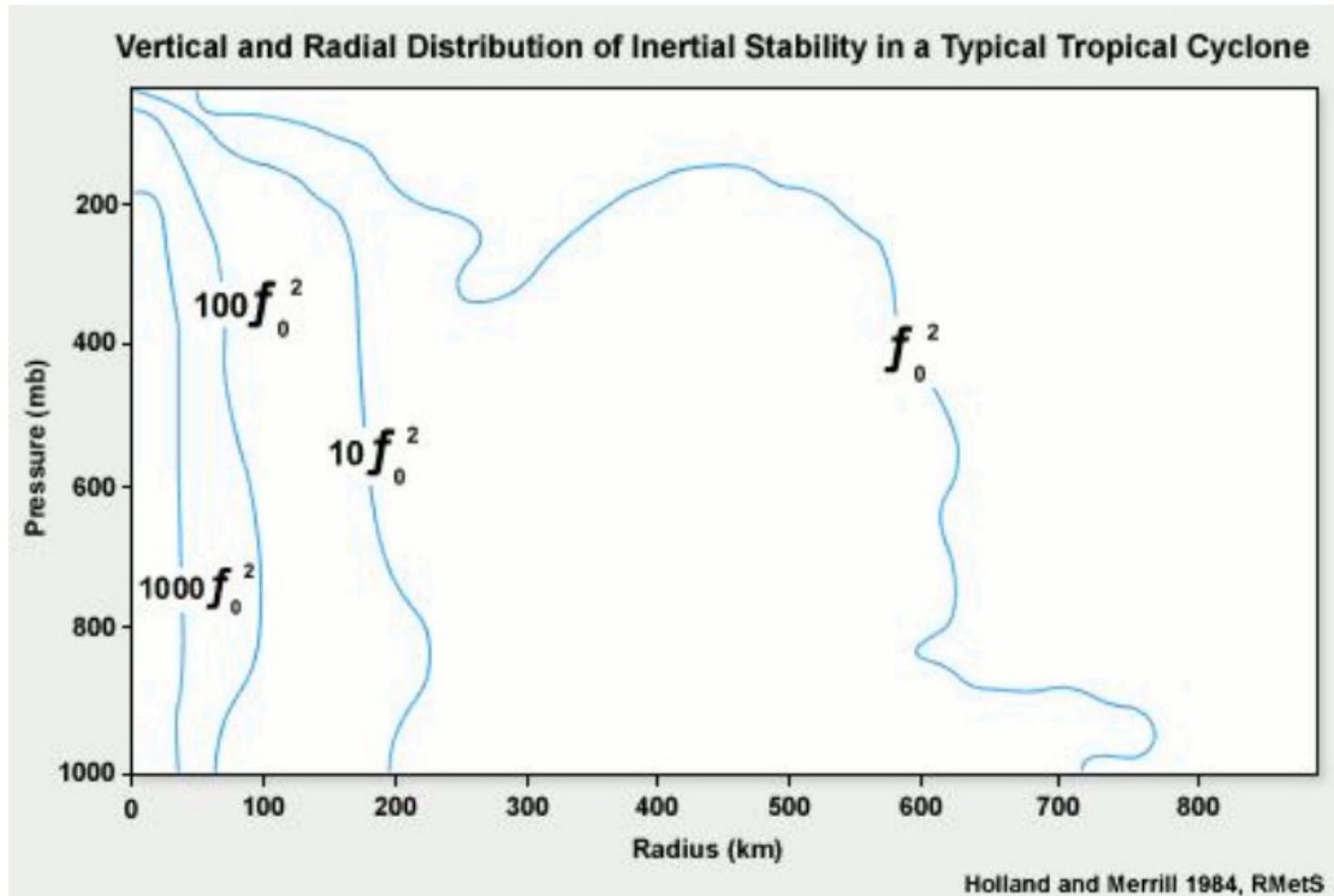


L_R



$L > L_R$

Cyclogenesis : the importance of vortex stability



Src : COMET Program

Figure 4.2. Vertical and radial distribution of the inertial stability in a typical tropical cyclone. To illustrate the contribution to inertial stability of the storm winds compared to its environment, the inertial stability values have been scaled by f_0 , the value of the Coriolis parameter at the storm center. Figure adapted by COMET from Holland and Merrill (1984) and obtained from http://www.meted.ucar.edu/tropical/textbook/ch10/tropcyclone_10_2_2_2.html.

Cyclogenesis : the importance of vortex stability

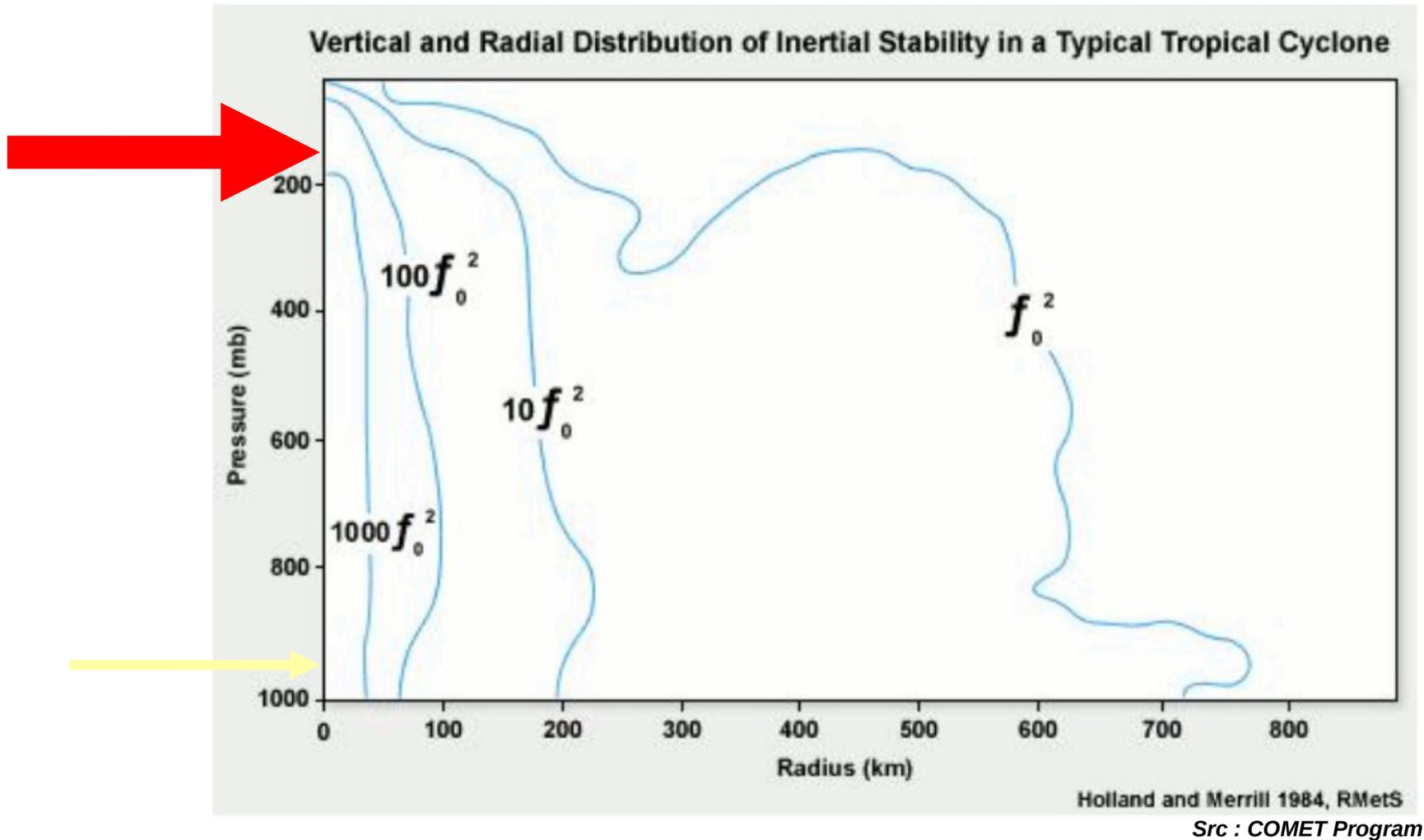
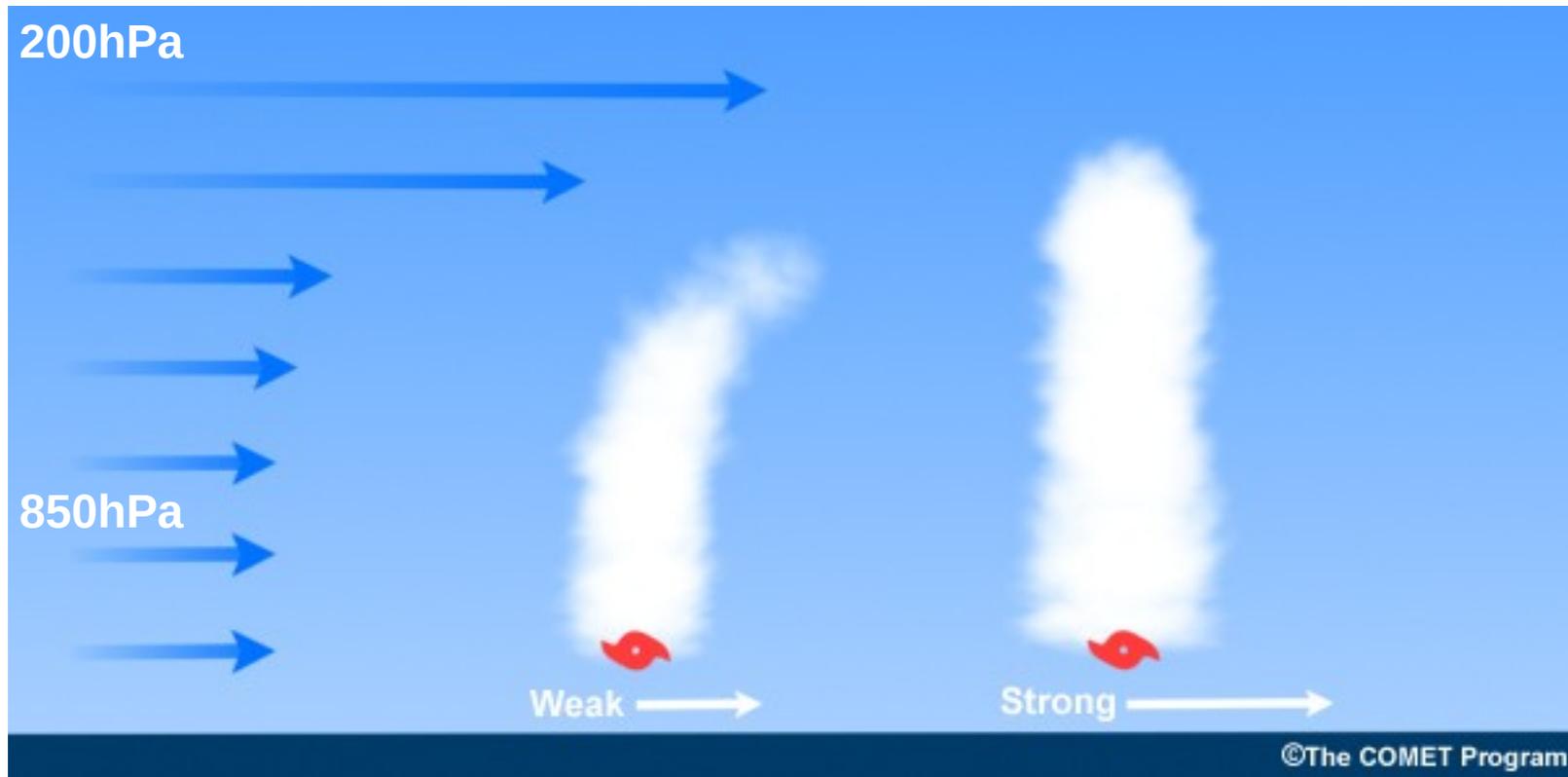


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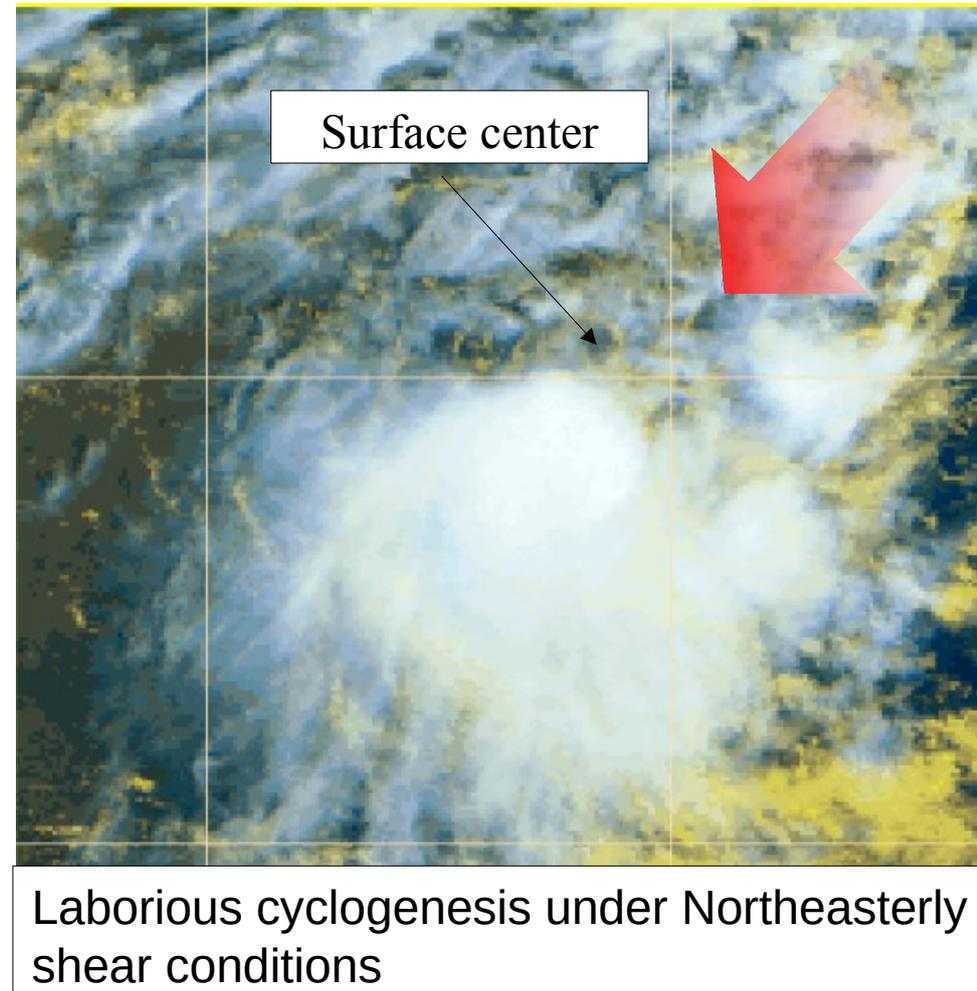
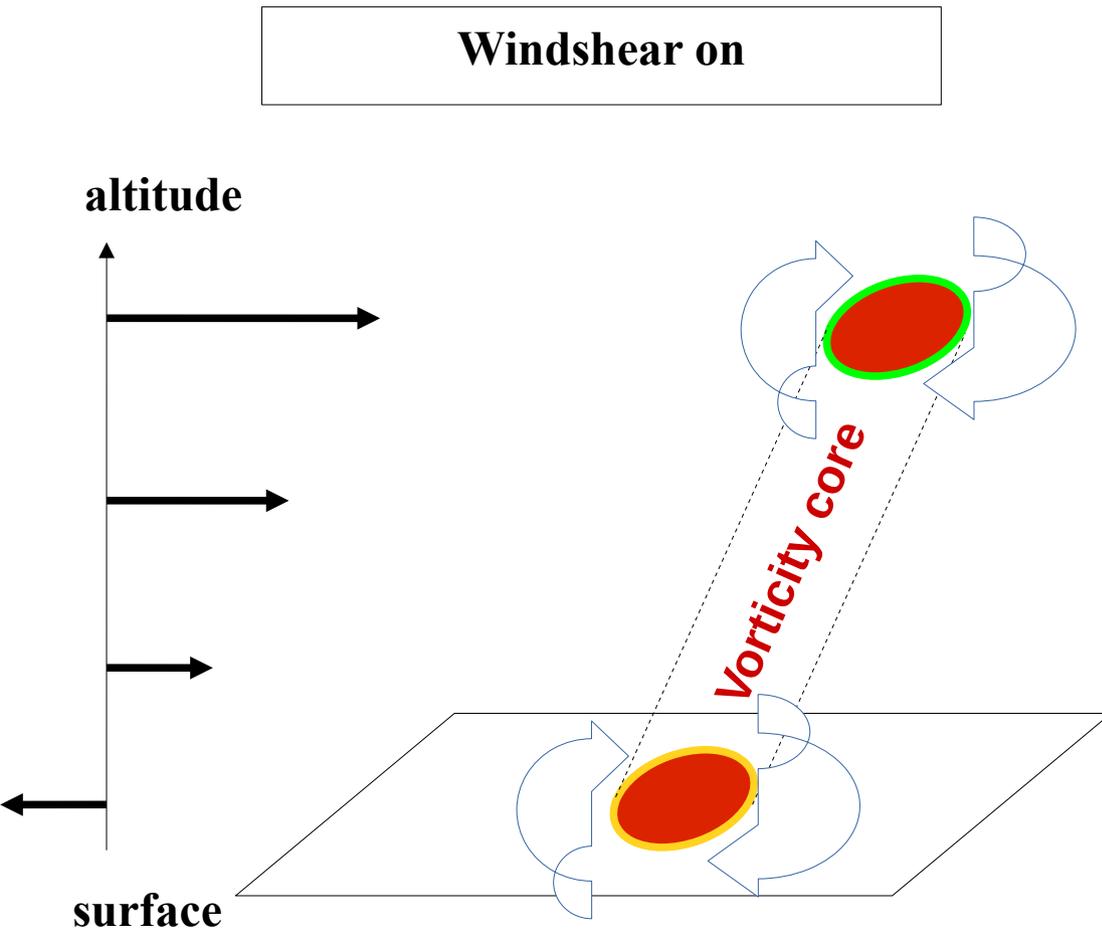
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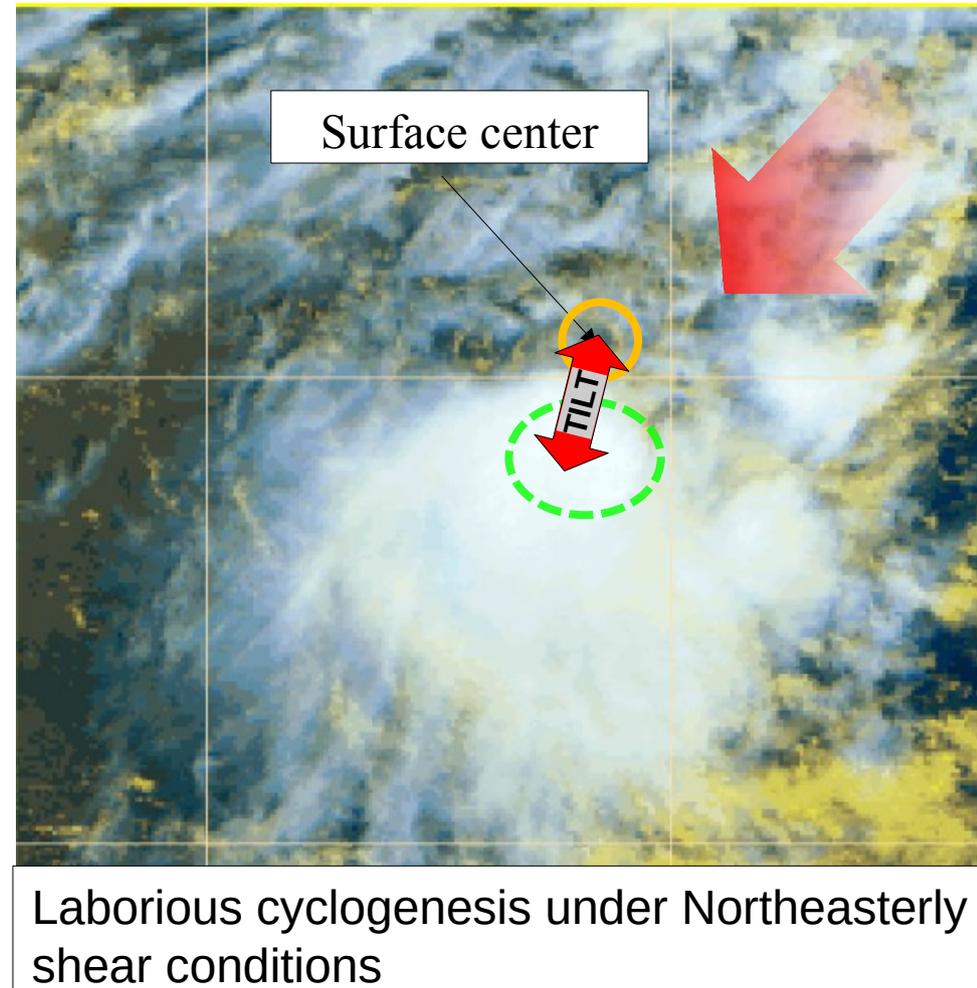
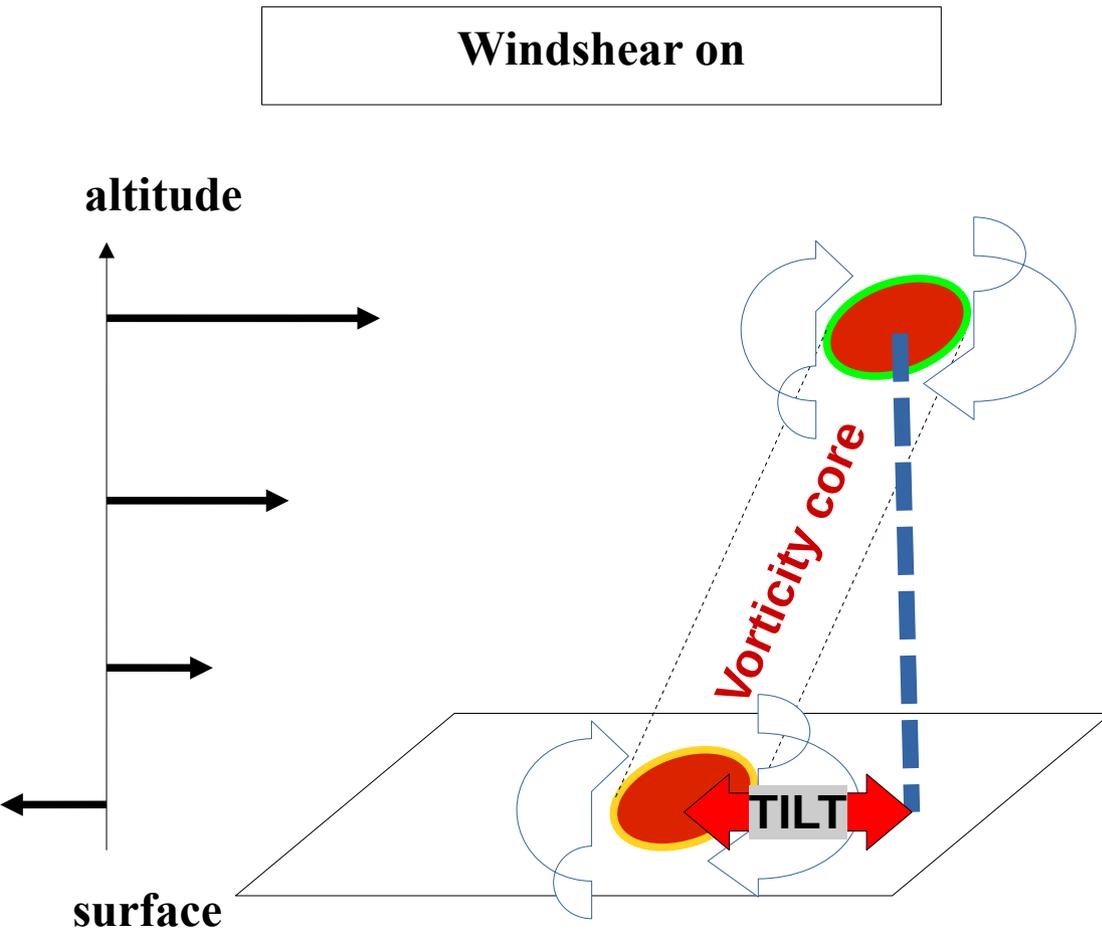
- **The windshear in the cyclogenesis process**

Src : MF Archives



- **The windshear in the cyclogenesis process**

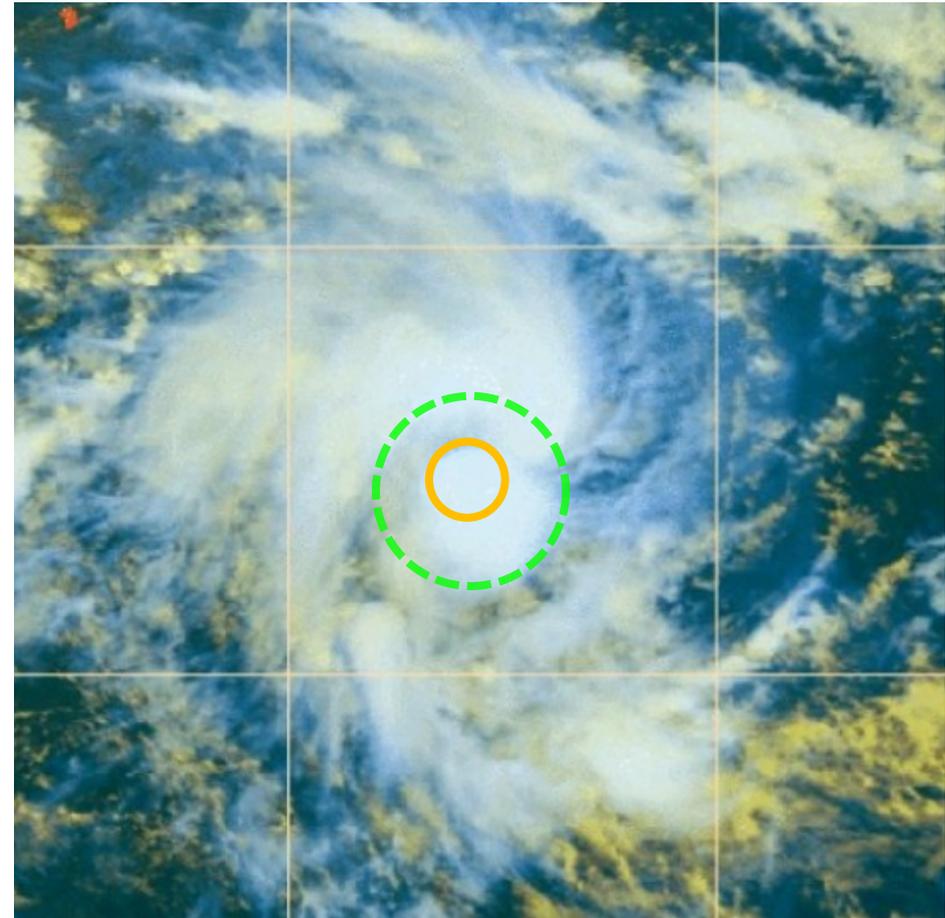
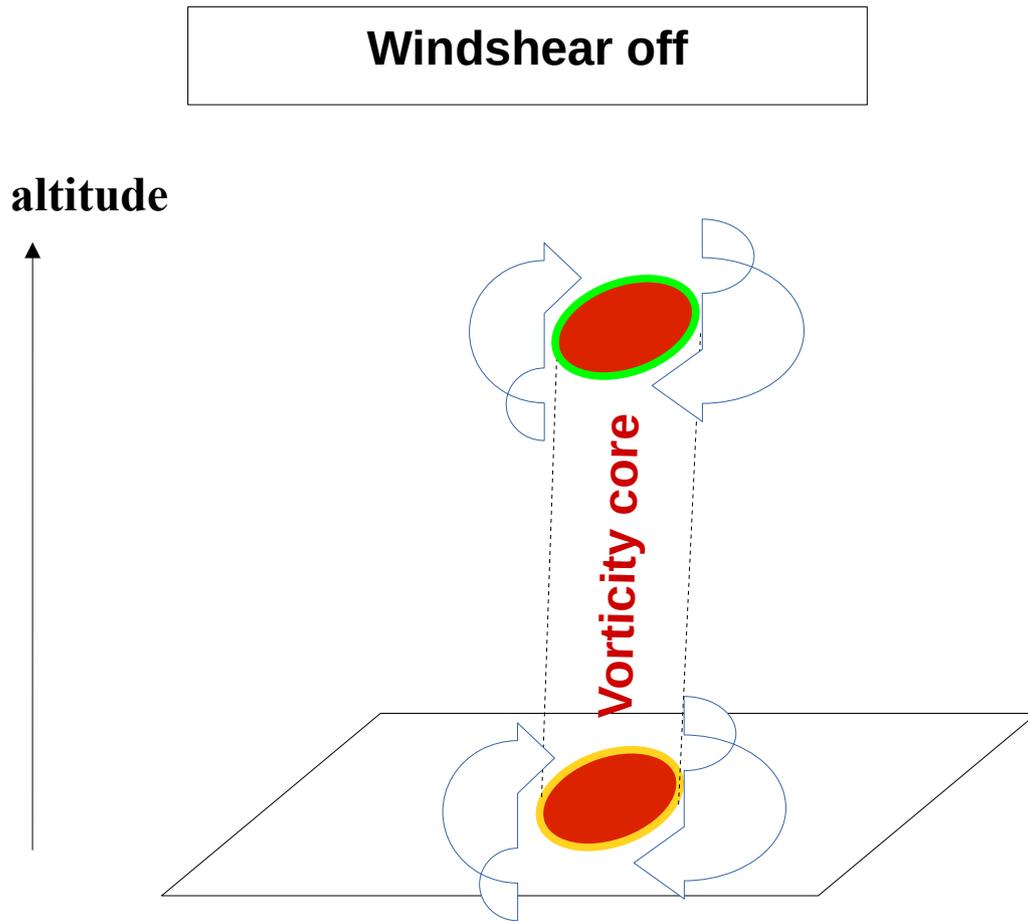
Src : MF Archives



→ Misaligned low and mid-levels vortices can slow down or even impede the TCG process

- **The windshear in the cyclogenesis process**

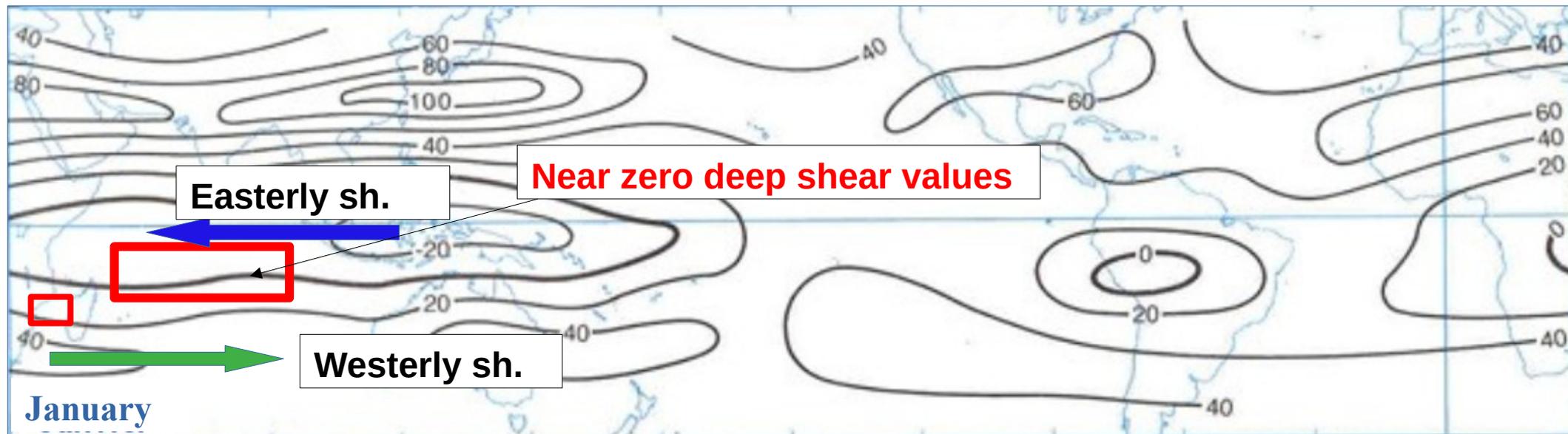
Src : MF Archives



→ When low and mid-levels vortices finally align (through strong convective bursts and/or weakening of the shear constraint), TCG occurs.

The windshear in the cyclogenesis process

Src : MF Archives



January climatology of the zonal component of the deep shear: U200 - U850 (in kt)

→ Weak vertical wind shear (< 20 kt) is observed within the SWIO development zones (especially for the cyclogenesis zone in the central Indian Ocean - this is less true for the secondary zone of the Mozambique Channel)

Tropical cyclogenesis

Basins configurations



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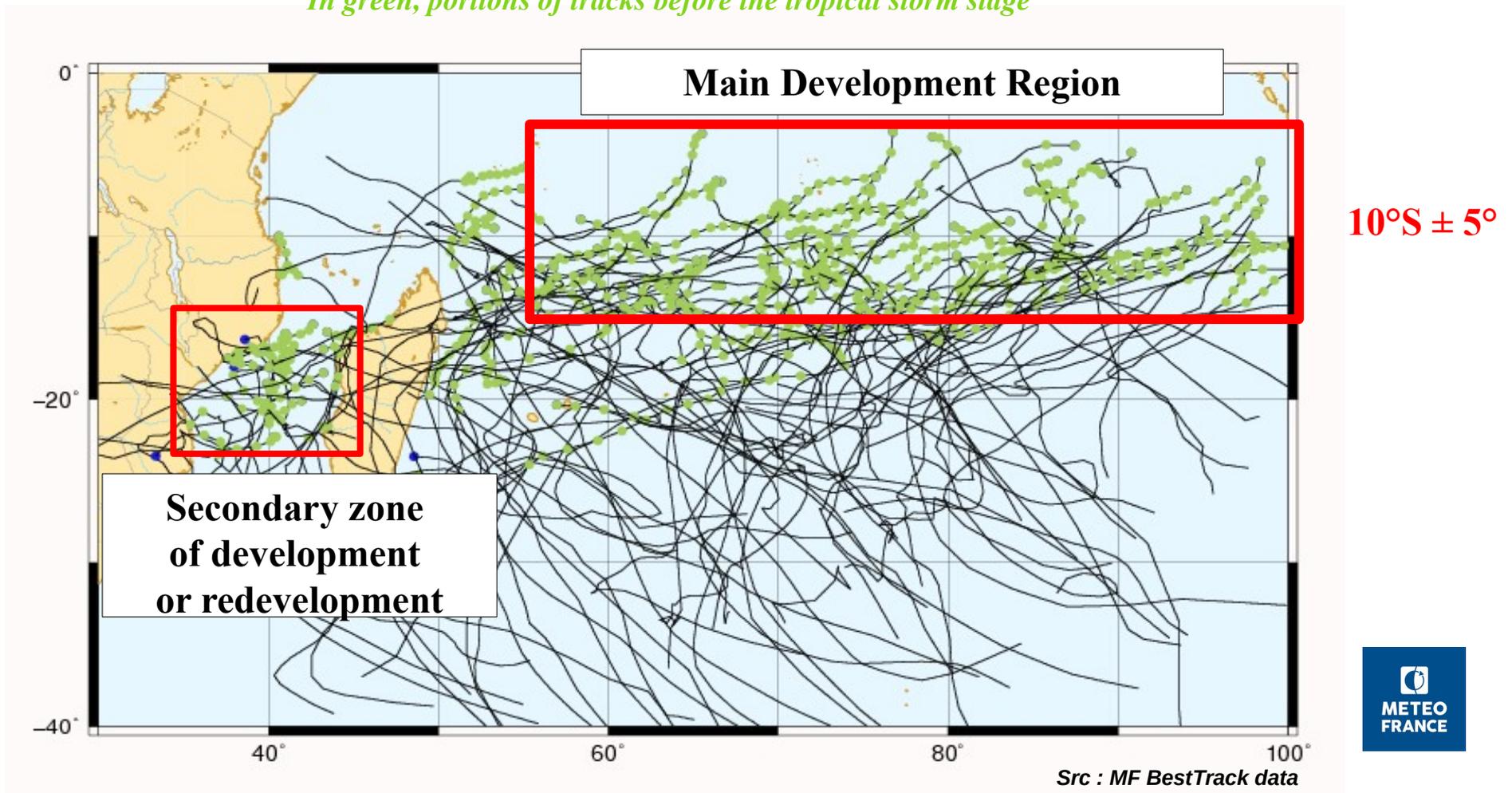
AI generated (Dall-E via Bing)

- **Cyclogenesis**

Definition : A tropical cyclogenesis is carried out when a low pressure area has become autonomous and no longer needs the help of its environment to develop, through "environmental forcing". In operational terms, cyclogenesis is carried out when the stage of a tropical storm is reached. It is a process that usually takes several days.

Location of cyclogenesis processes on the SWIO (2010/2011-2020/2021)

In green, portions of tracks before the tropical storm stage



- **Basin pattern / Bassin configuration**

Definition : Seasonal and specific low levels wind patterns over the tropical ocean. They are observed on a monthly timescale.

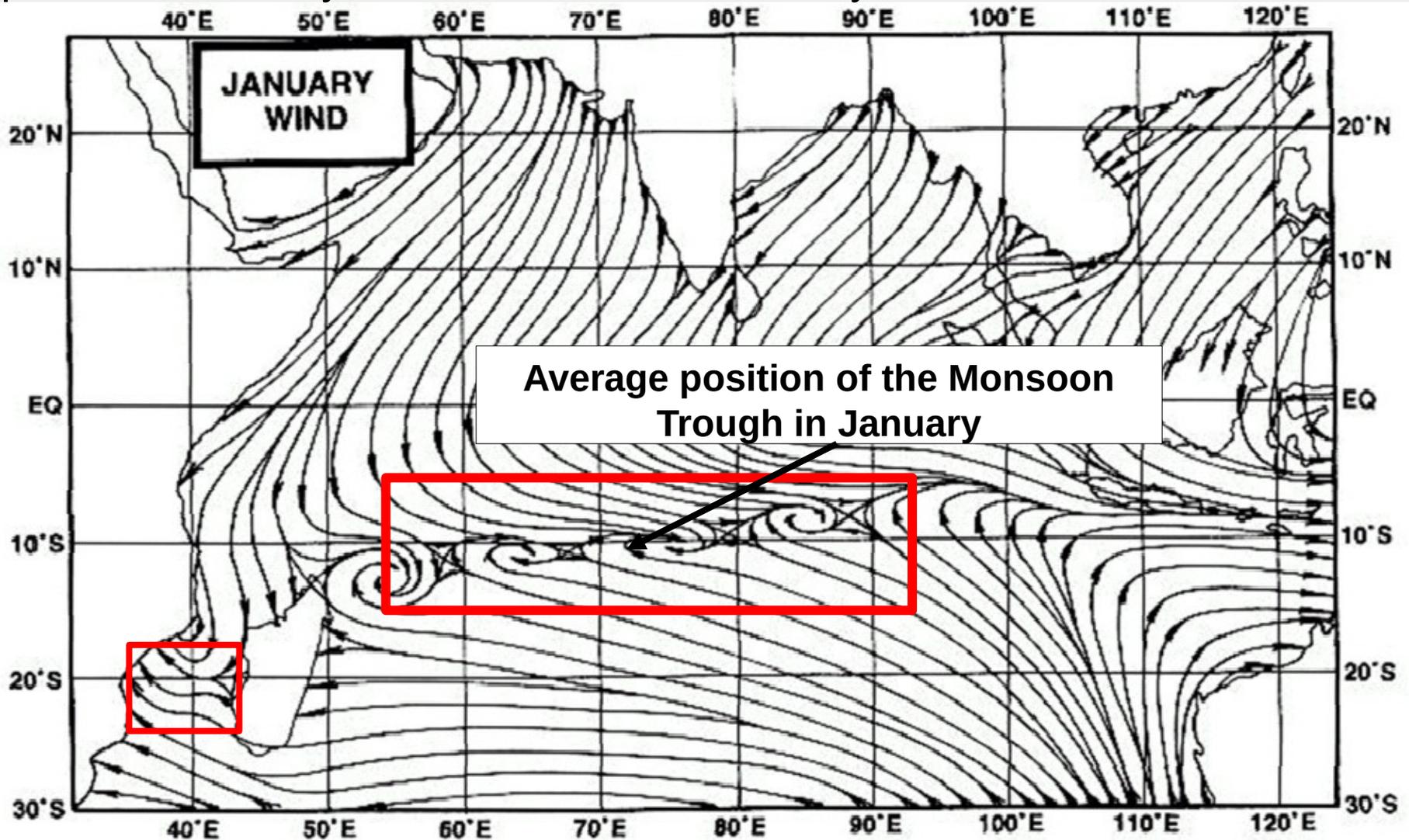


Figure 2.6. Mean surface level streamline analyses over the Indian Ocean for January (Sadler, 1975).

- **Basin Configuration**
Conceptual model of the Monsoon Trough (MT)

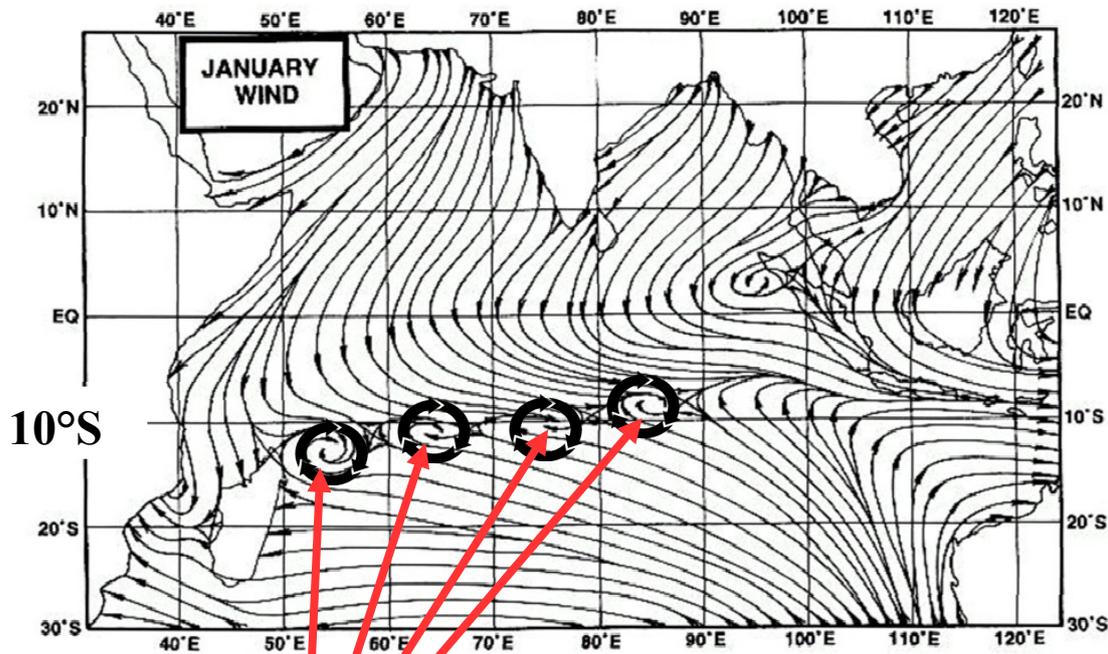
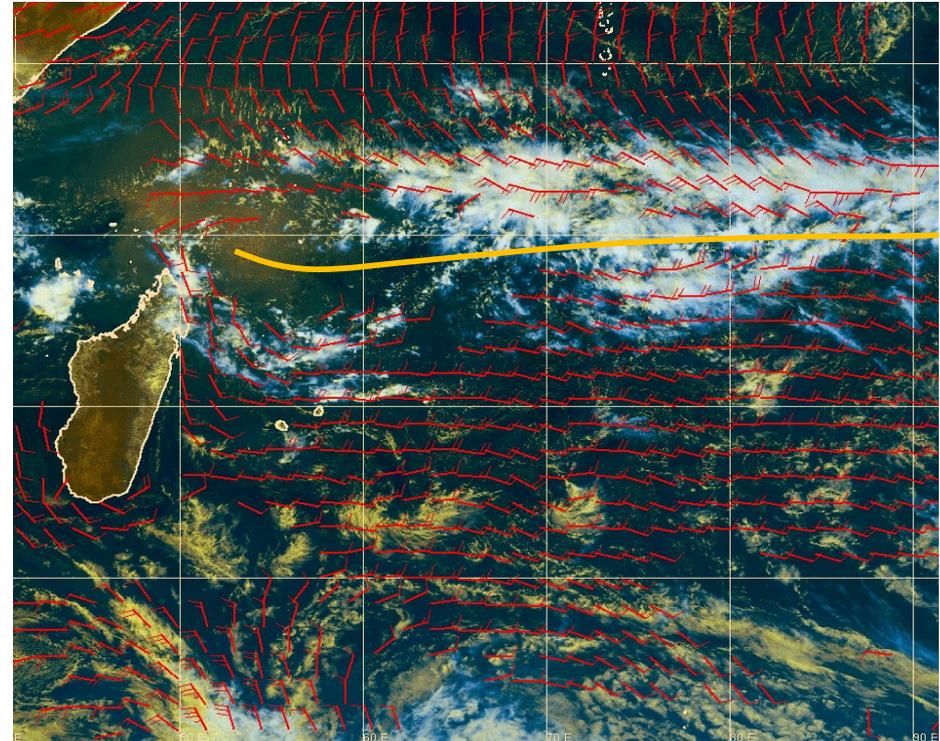


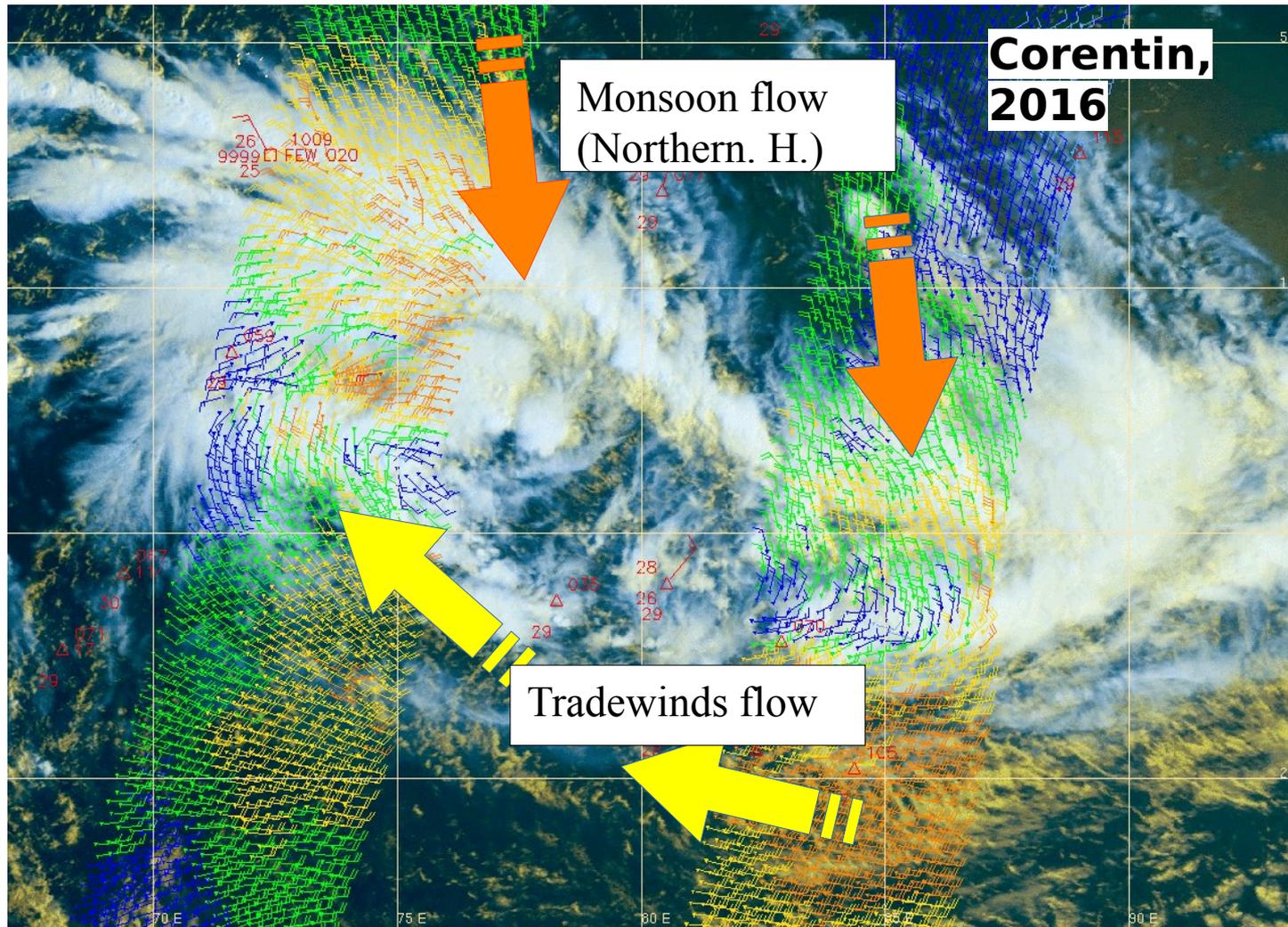
Figure 2.6 Mean surface level streamline analyses over the Indian Ocean for January (Saji et al., 1993).



Strongly enhance low levels vorticity in the Monsoon Trough. This is the most common configuration for **the heart of the season (~ January to March).**

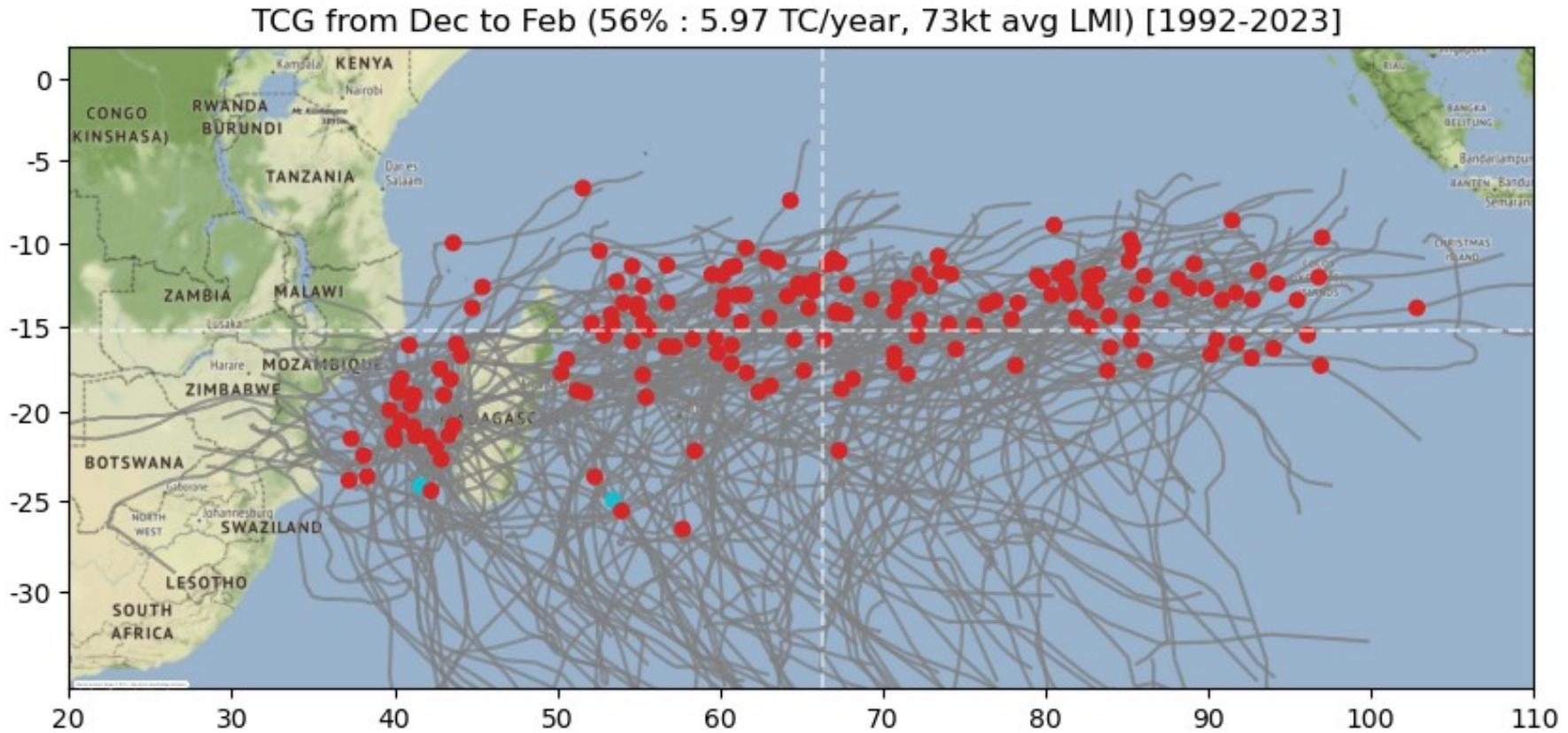
- The dynamics of the Monsoon trough, a synoptic or supra-synoptic low-level forcing

Src : MF Archives



Renforcement de la convergence entre les flux de moussons et d'alizés

- **Basin pattern / Bassin configuration**



TCG locations in **Red** (first point with intensity of a Moderate TS), in **Cyan** if Subtropical System.

Associated tracks in **Grey**

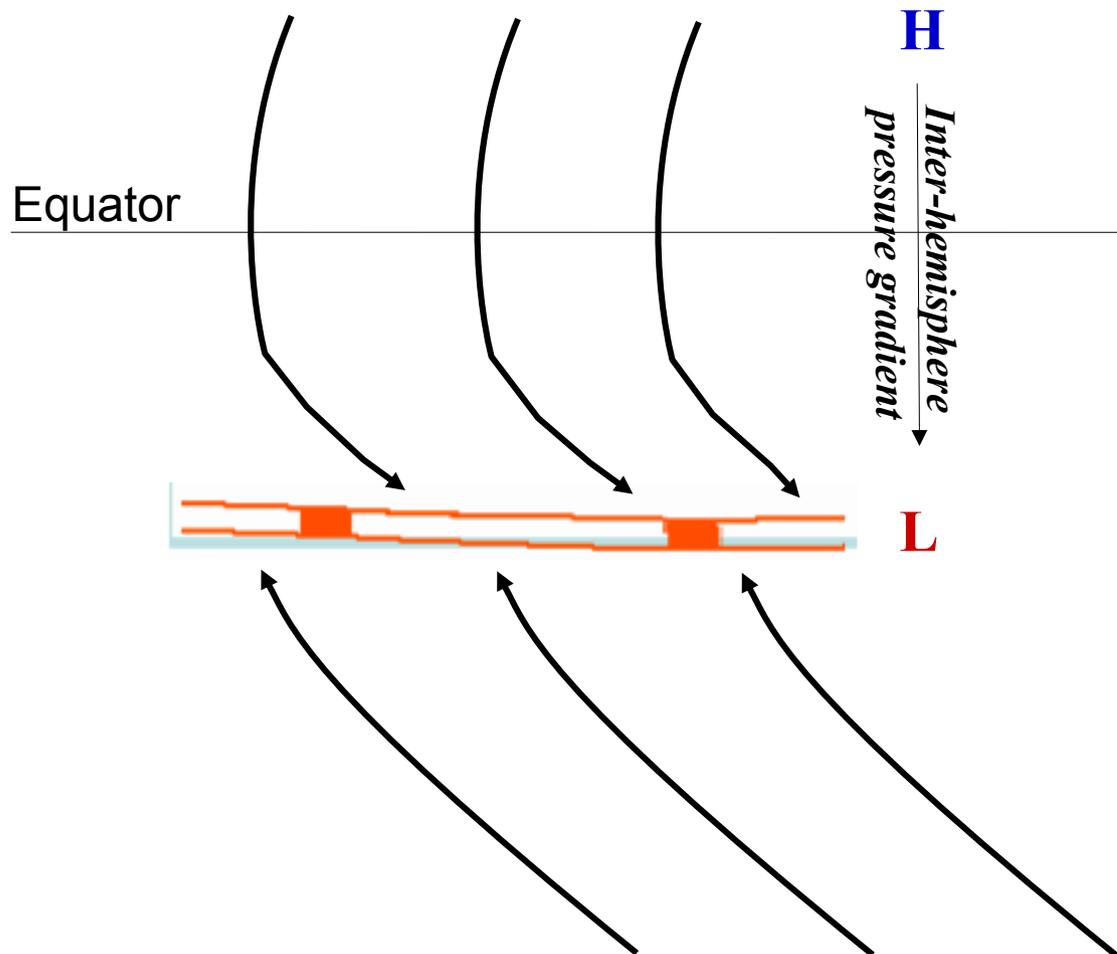
Src : MF BestTrack Data

Averaged lat and lons for the TCG of the sample shown by dotted white lines.

- **Basin Configuration**
Conceptual model of the Monsoon Trough (MT)

Overseas MT

Definition : Low tropospheric trough located in the mixing zone between the monsoon flow and the trade wind flow. The winds on the equatorial side of this trough have a strong meridional component as they cross the equator.

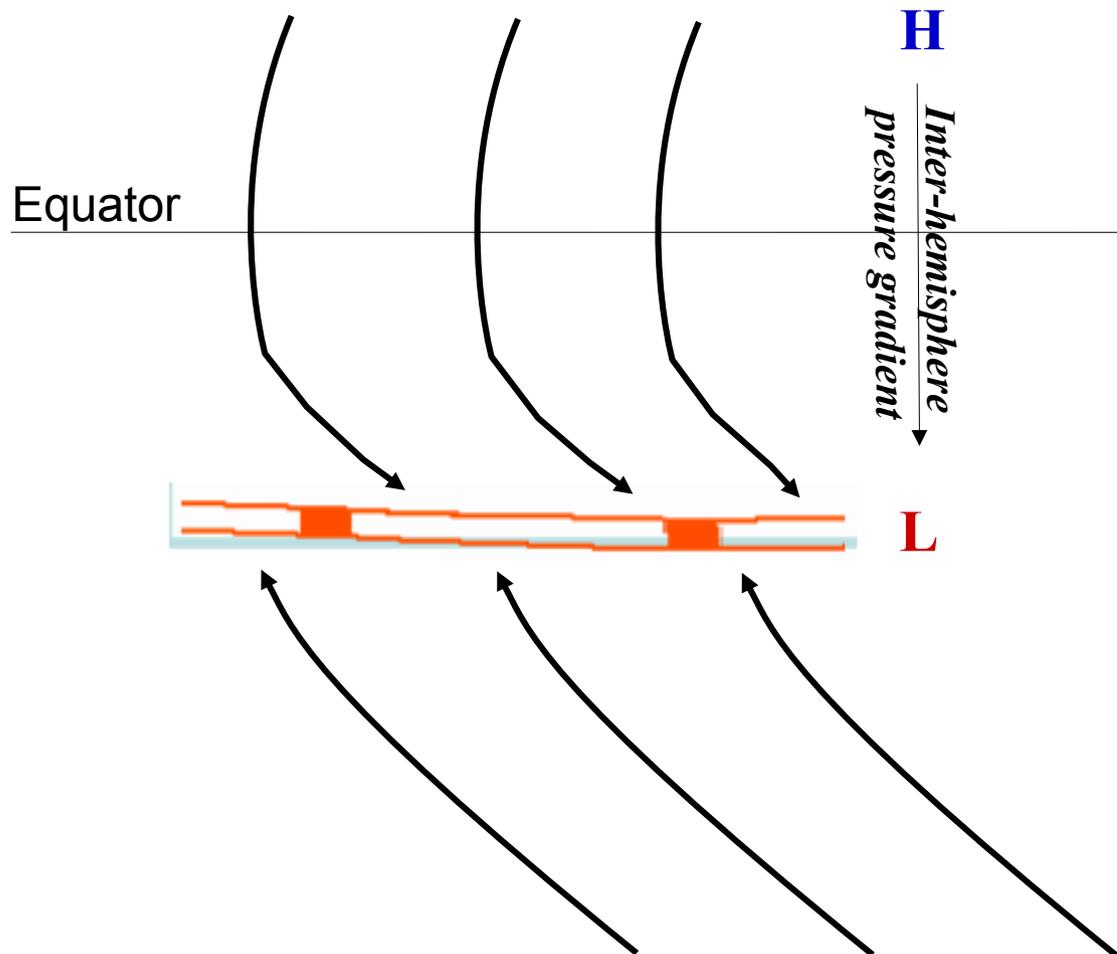


- **From the surface up to 850 hPa**
- **Strongest vorticity** along the axis
- **Large scale moisture convergence**

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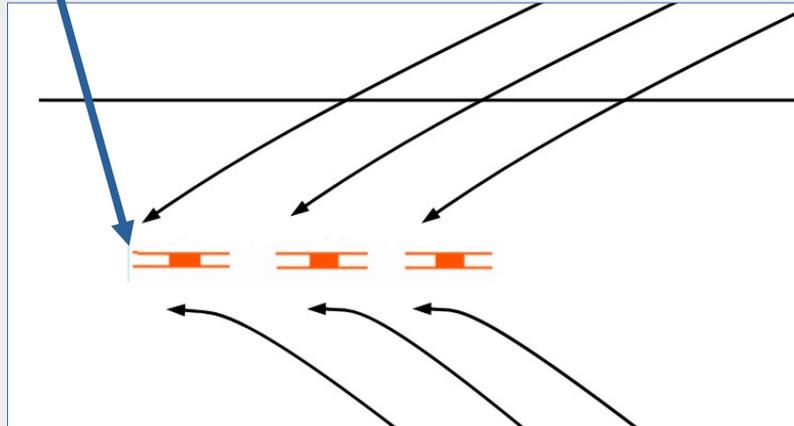
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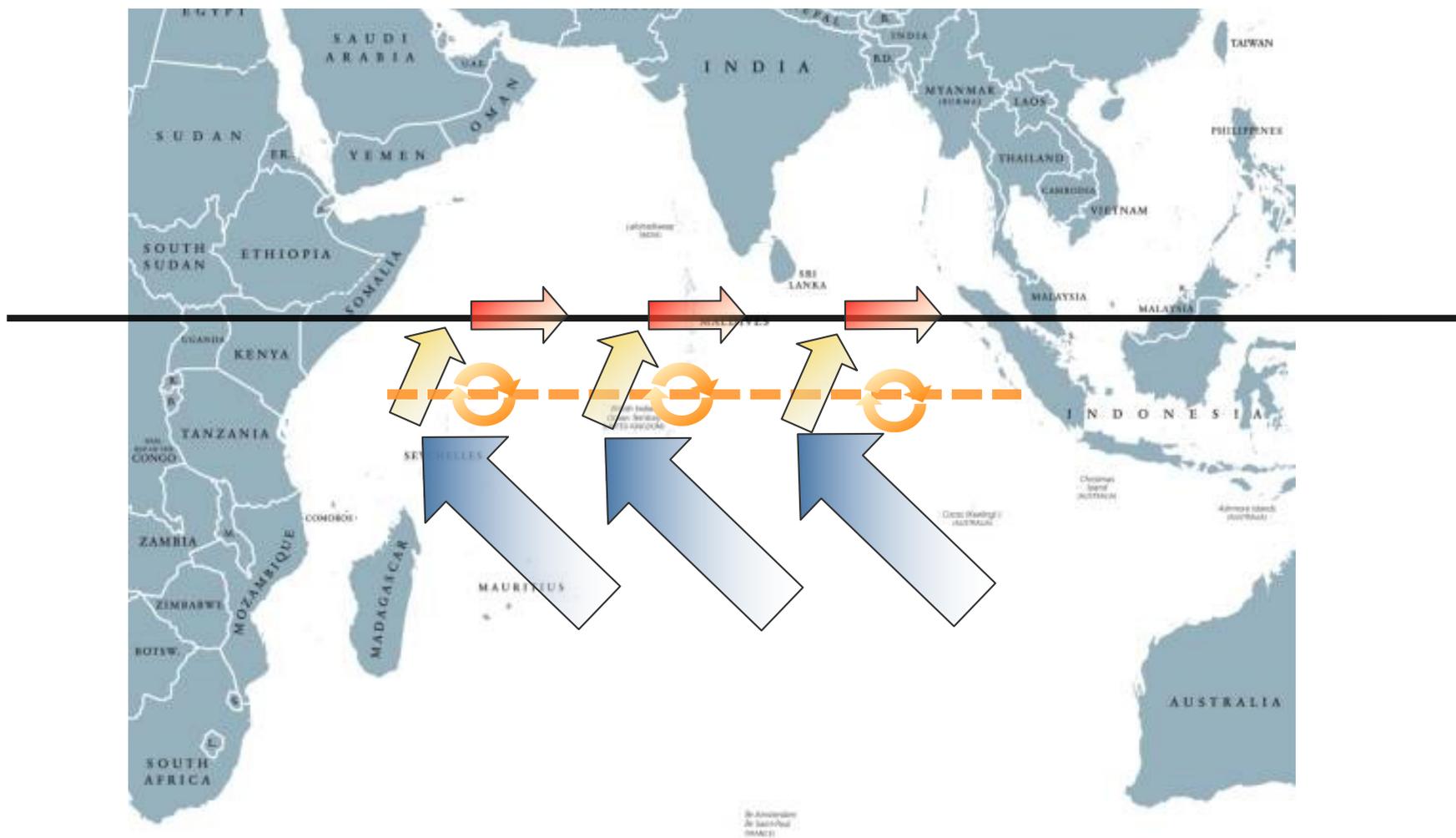
Definition : Low tropospheric trough located in the mixing zone between the monsoon flow and the trade wind flow. The winds on the equatorial side of this trough have a strong meridional component as they cross the equator.

This is not a MT. This is the Tradewinds Meteorological Equator.



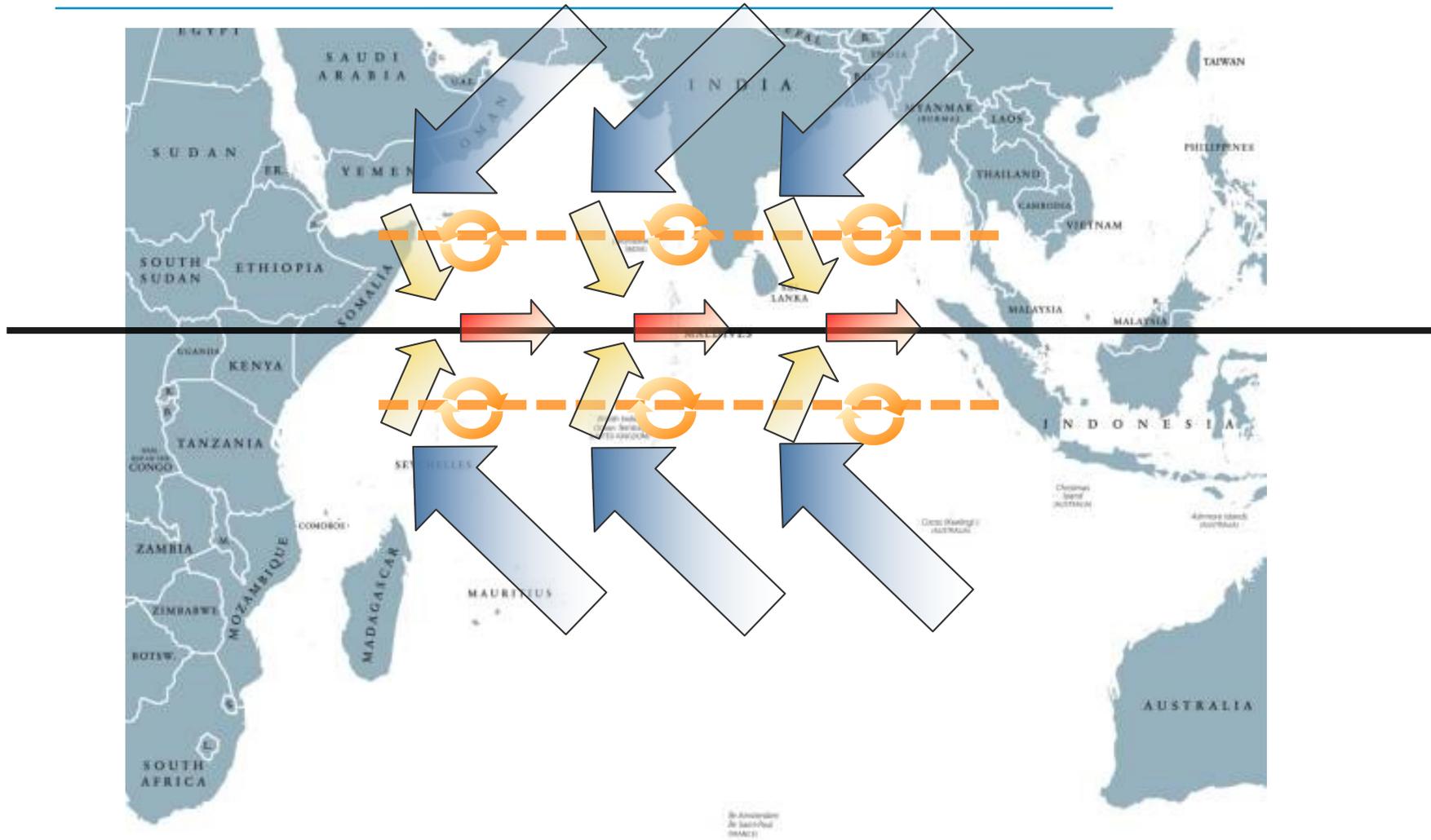
Surface up to 850 hPa
Turbidity along the axis
Moisture convergence

- **Basin Configuration**
Conceptual model of the Near Equatorial Trough (NET)



Definition : Low tropospheric trough with strong zonal winds on its equatorial side

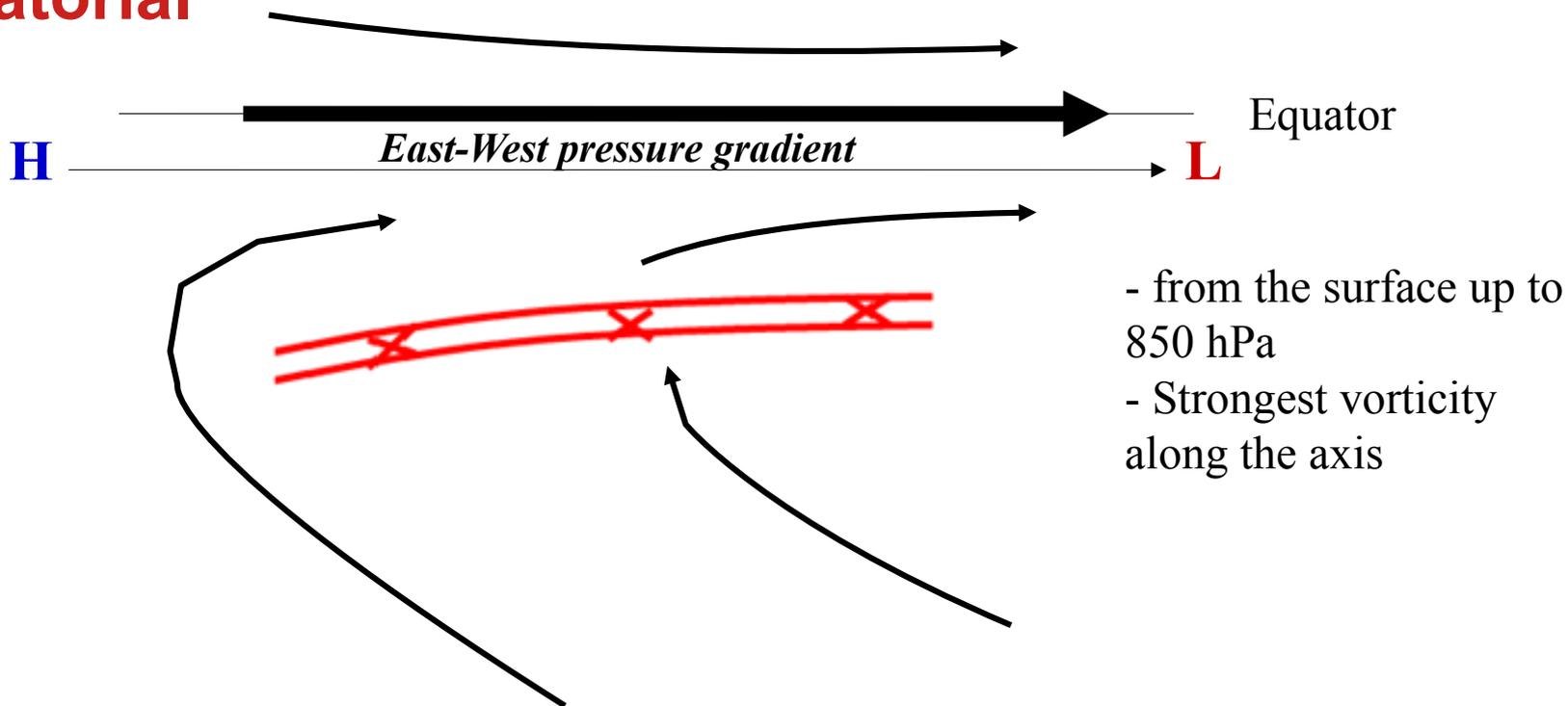
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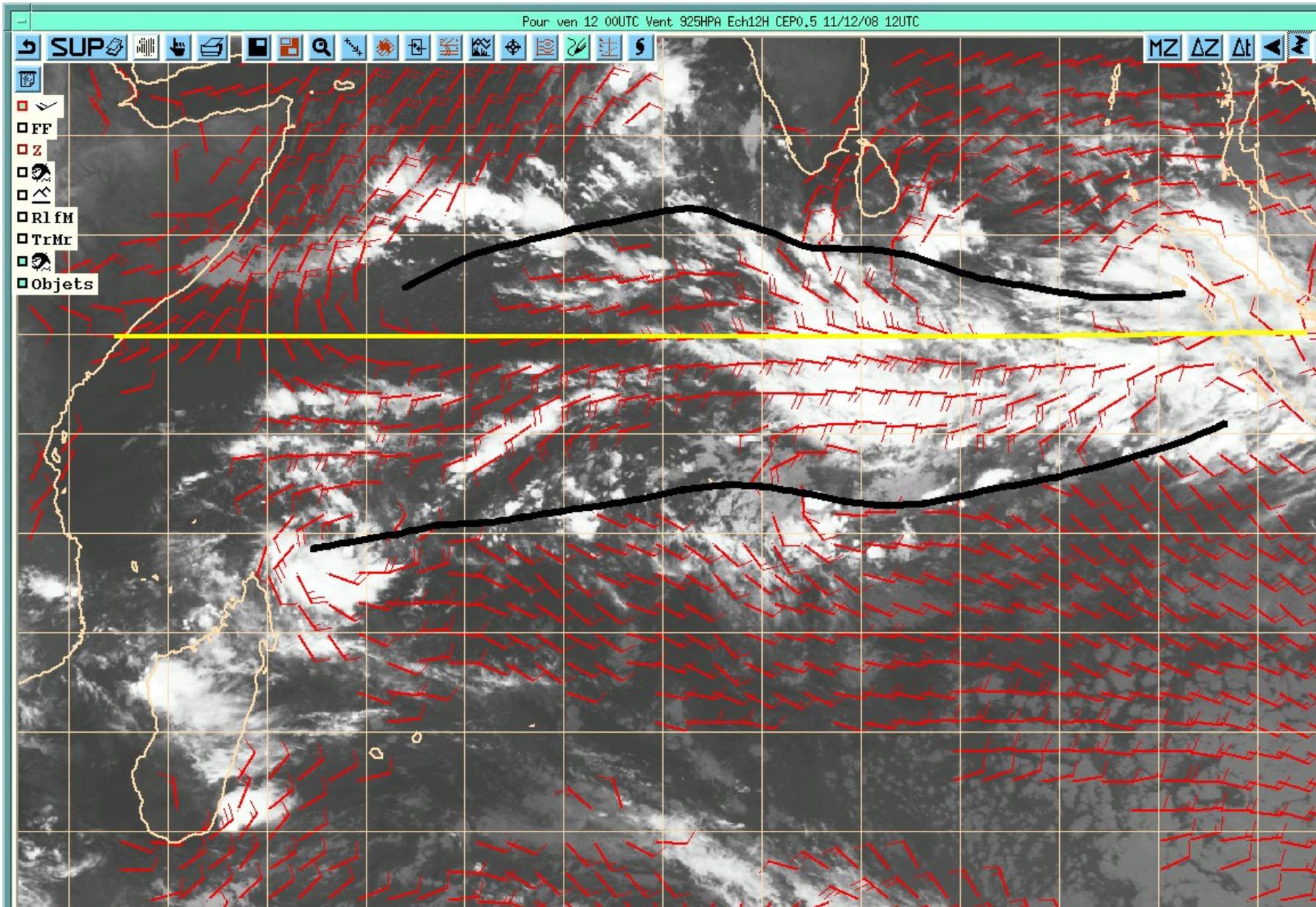
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Near Equatorial Trough

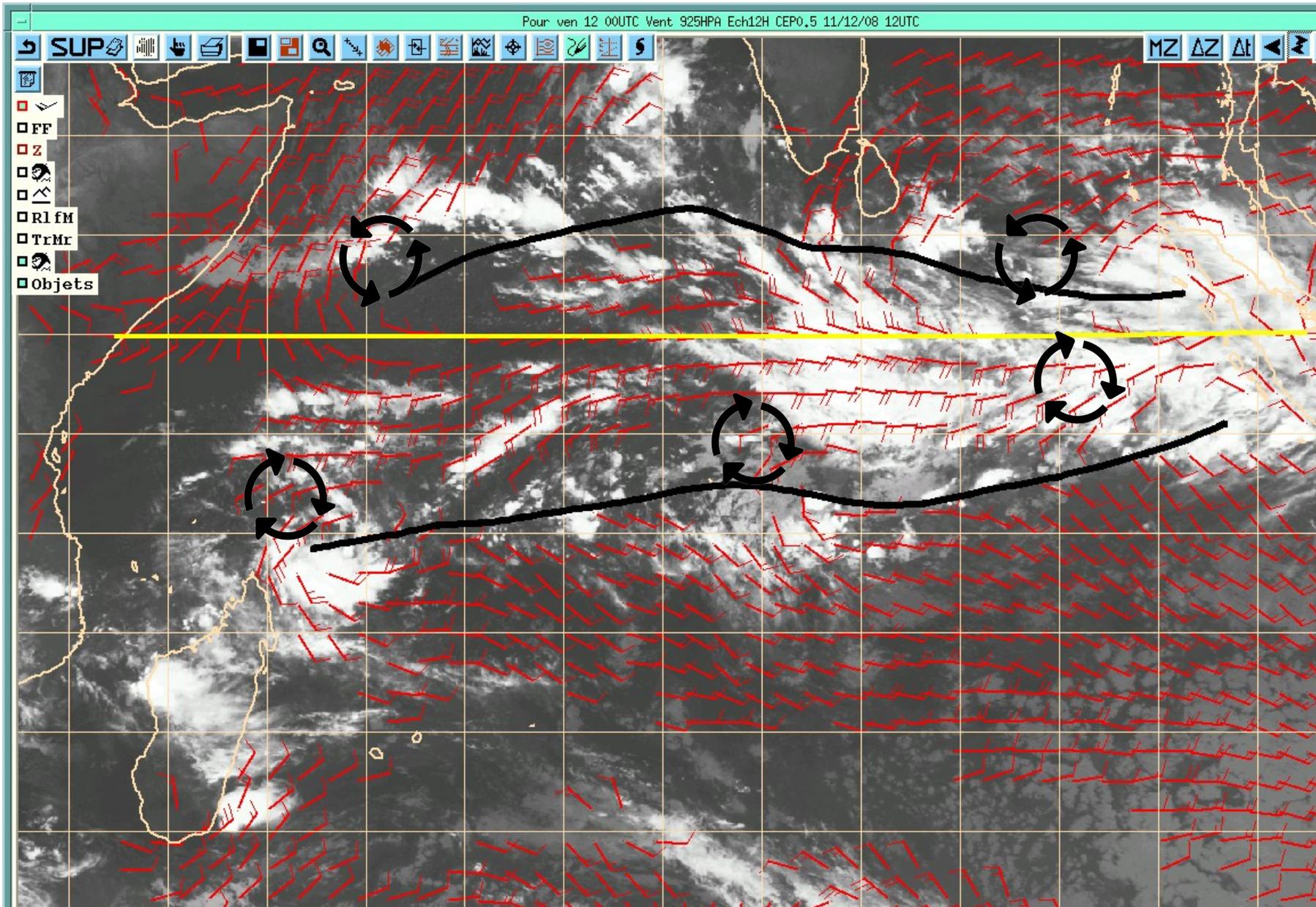


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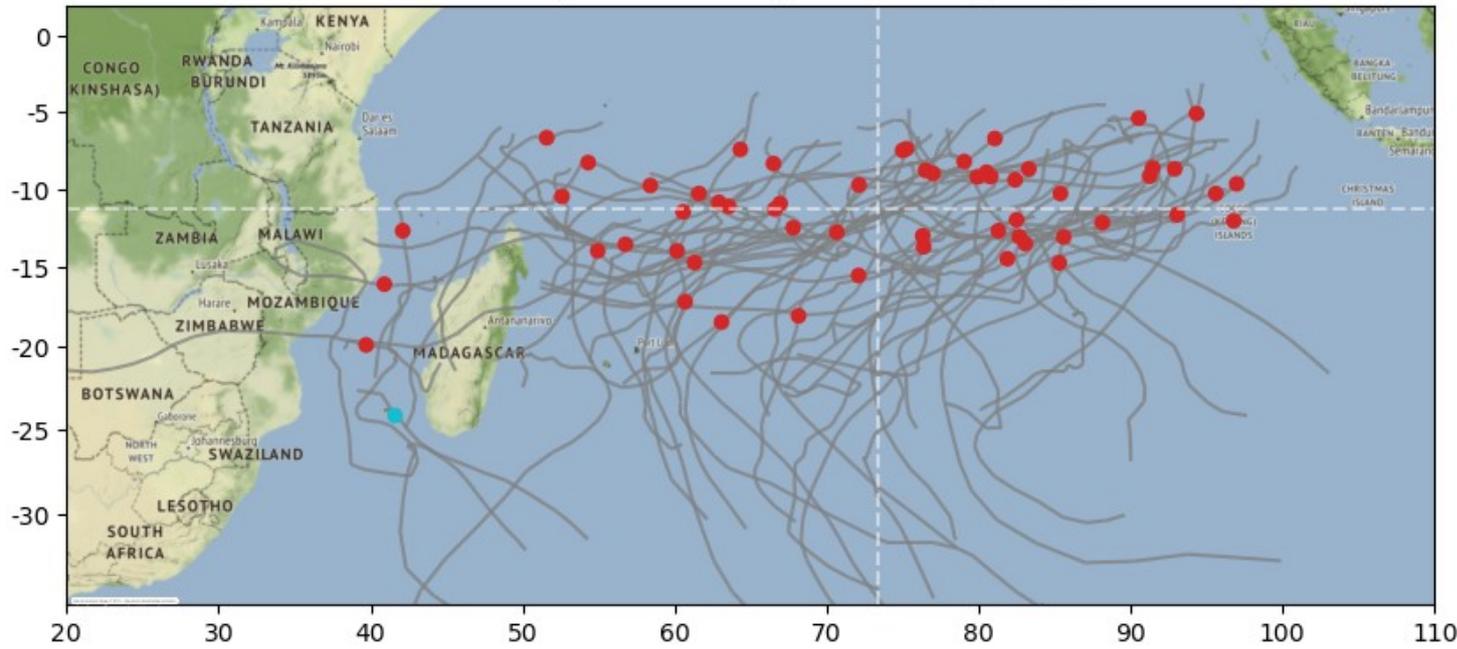
- **Basin Configuration**
Conceptual model of the Near Equatorial Trough (NET)



Basin Configuration

Conceptual model of the Near Equatorial Trough (NET)

TCG from Nov to Dec (18% : 1.93 TC/year, 73kt avg LMI) [1992-2023]

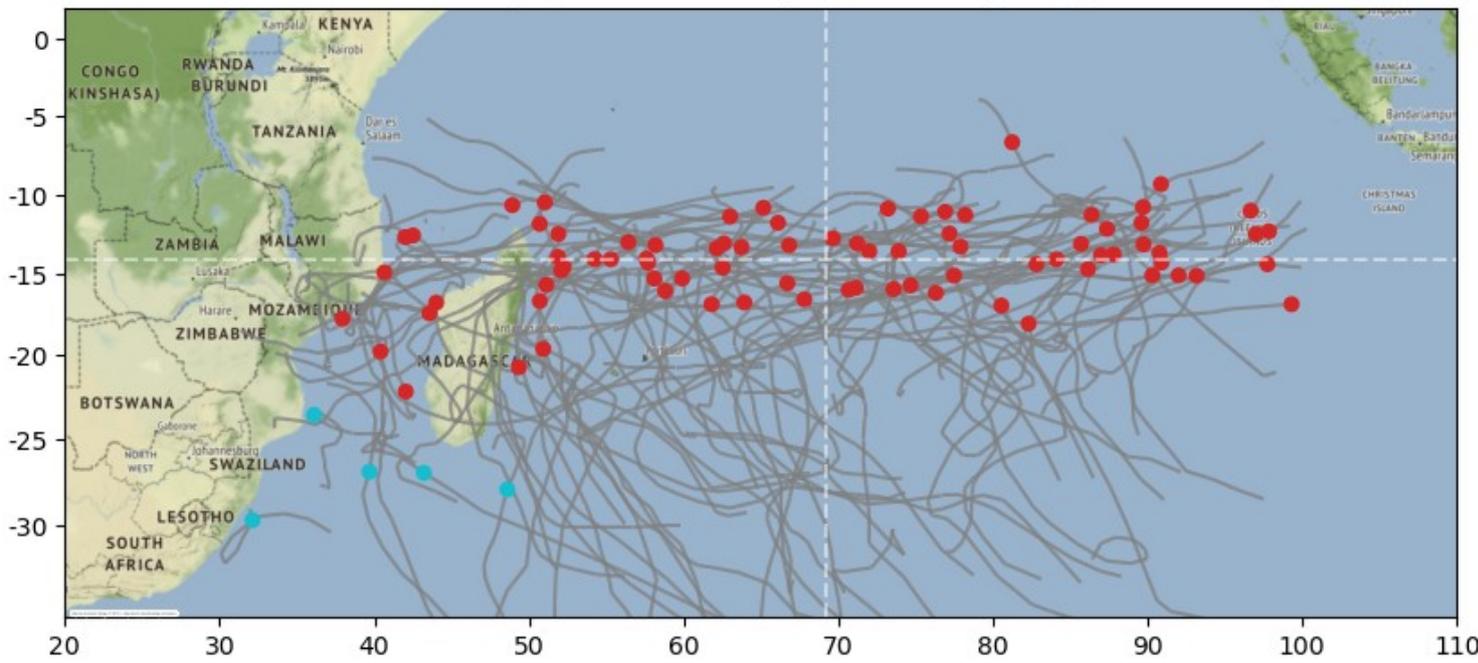


TCG locations in **Red** (first point with intensity of a Moderate TS), in **Cyan** if Subtropical System.

Associated tracks in **Grey**

Averaged lat and lons for the TCG of the sample shown by dotted white lines.

TCG from Mar to Apr (27% : 2.87 TC/year, 76kt avg LMI) [1992-2023]



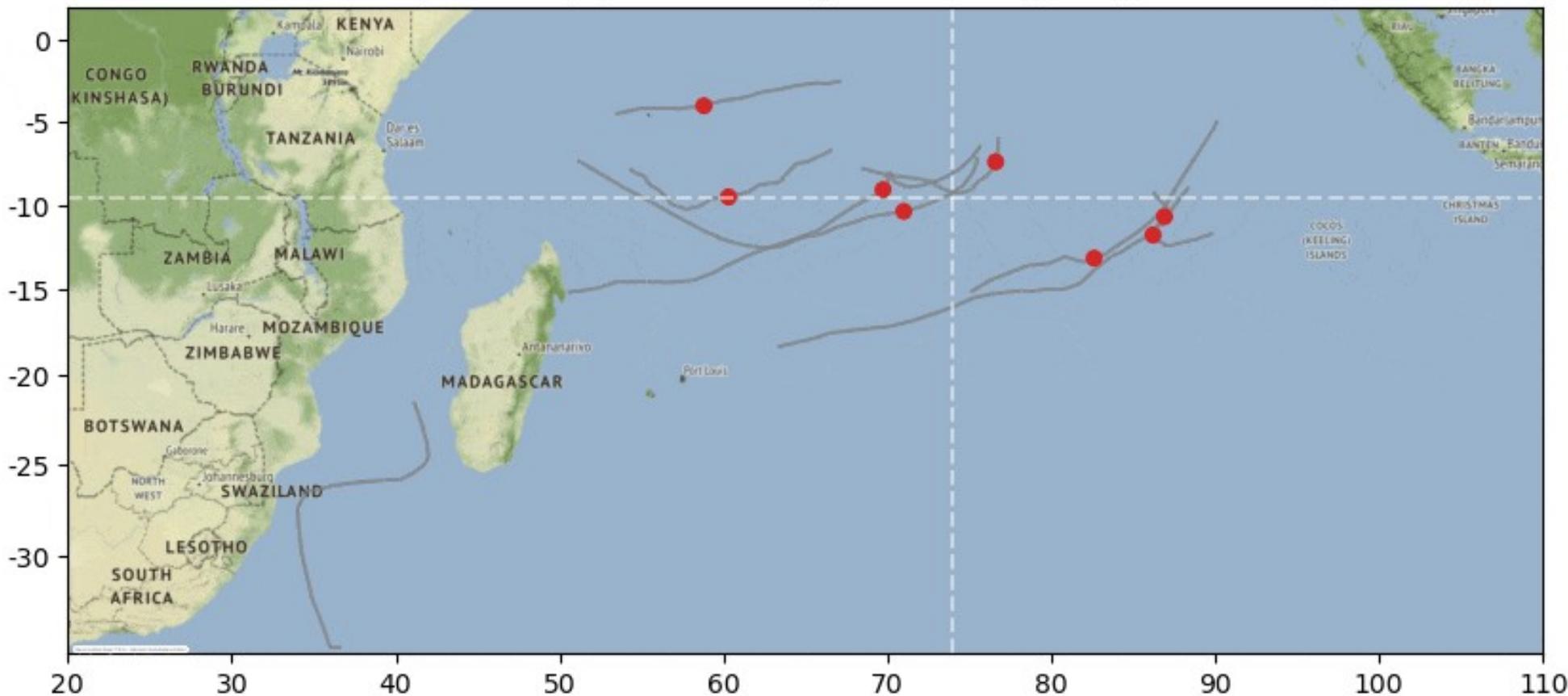
Src : MF BestTrack Data



Basin Configuration

Conceptual model of the Near Equatorial Trough (NET)

TCG from Jun to Sep (2% : 0.30 TC/year, 41kt avg LMI) [1992-2023]



TCG locations in **Red** (first point with intensity of a Moderate TS), in **Cyan** if Subtropical System.

Src : MF BestTrack Data

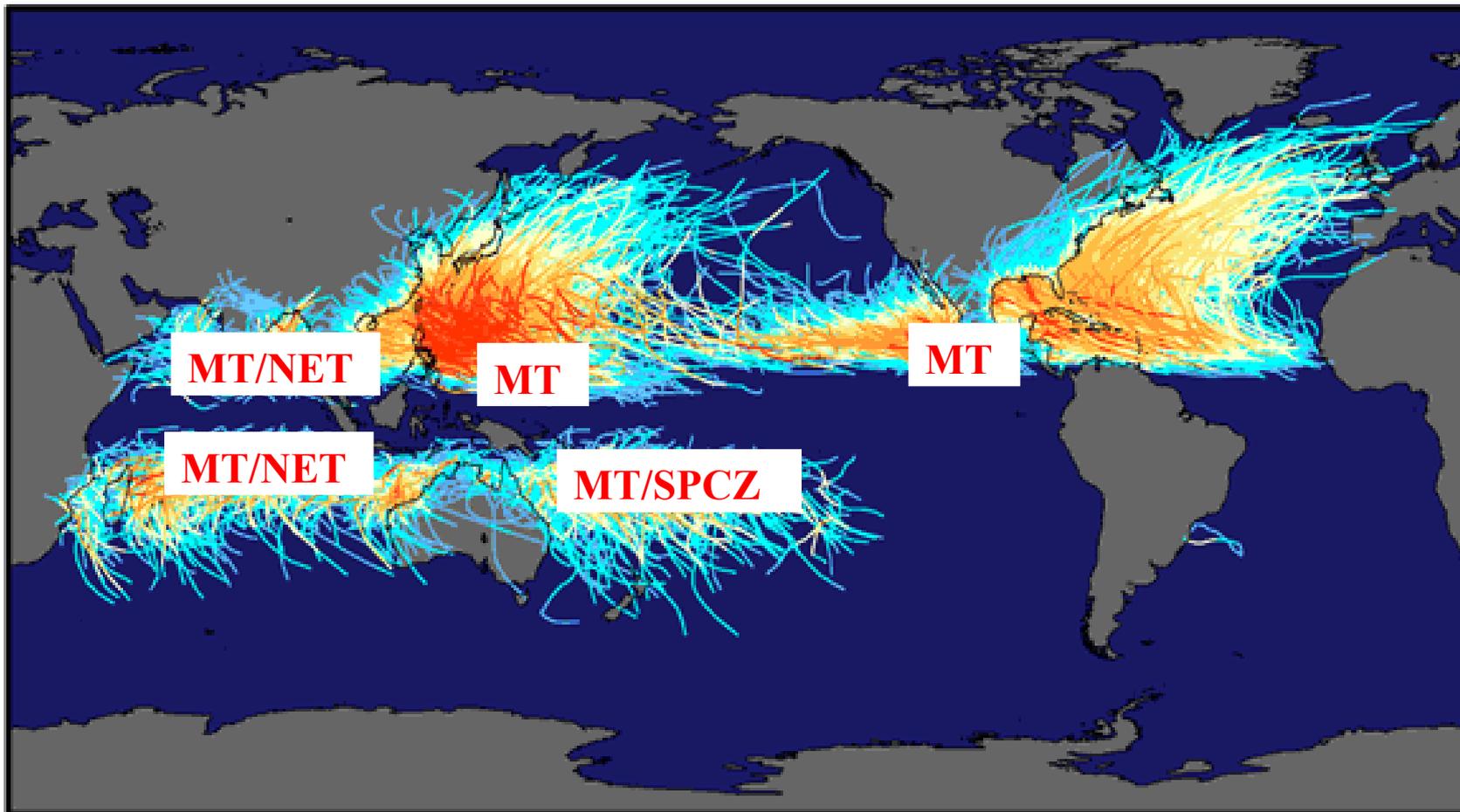
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- **Cyclogenesis : Bassin configurations worldwide**

Tracks and Intensity of Tropical Cyclones, 1851-2006



Saffir-Simpson Hurricane Intensity Scale

Robert A. Rohde, UC Berkeley / NASA's Earth Observatory



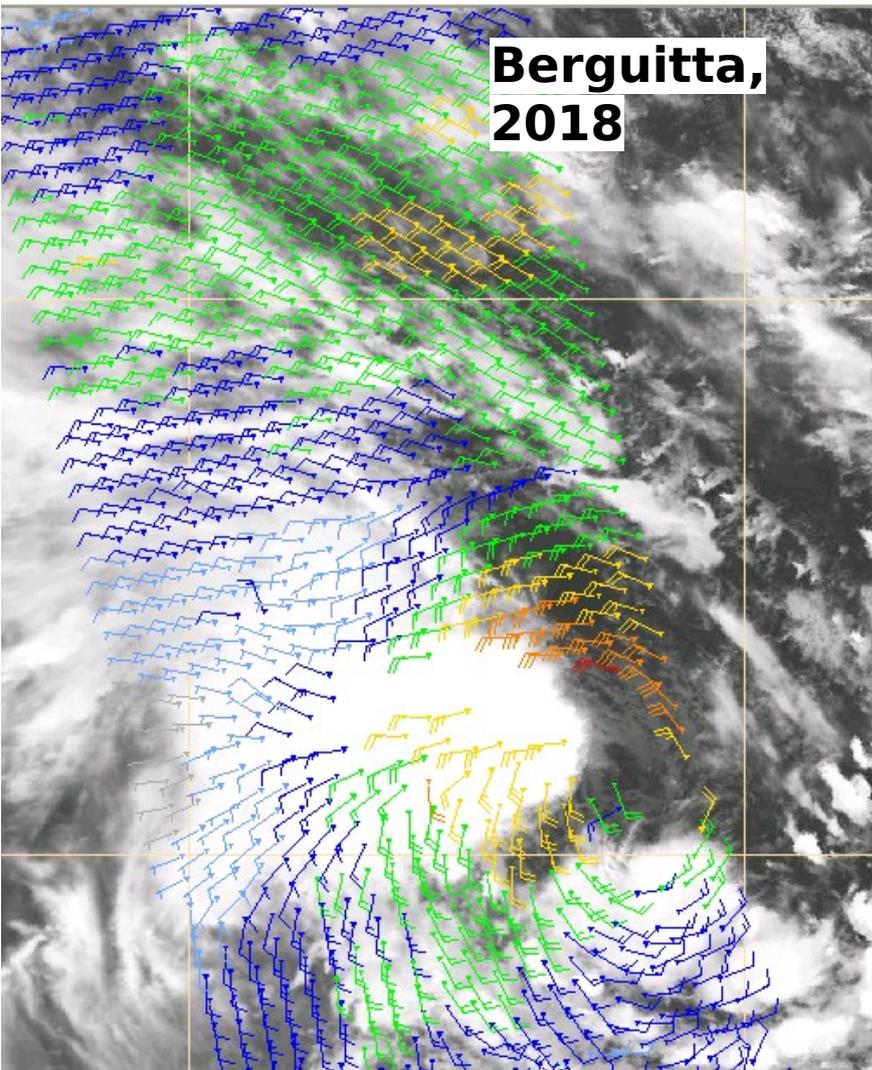
- **Low levels convergence enhancement: the spark that starts the cyclogenesis**

These wind surges can occur on both sides of the traffic:

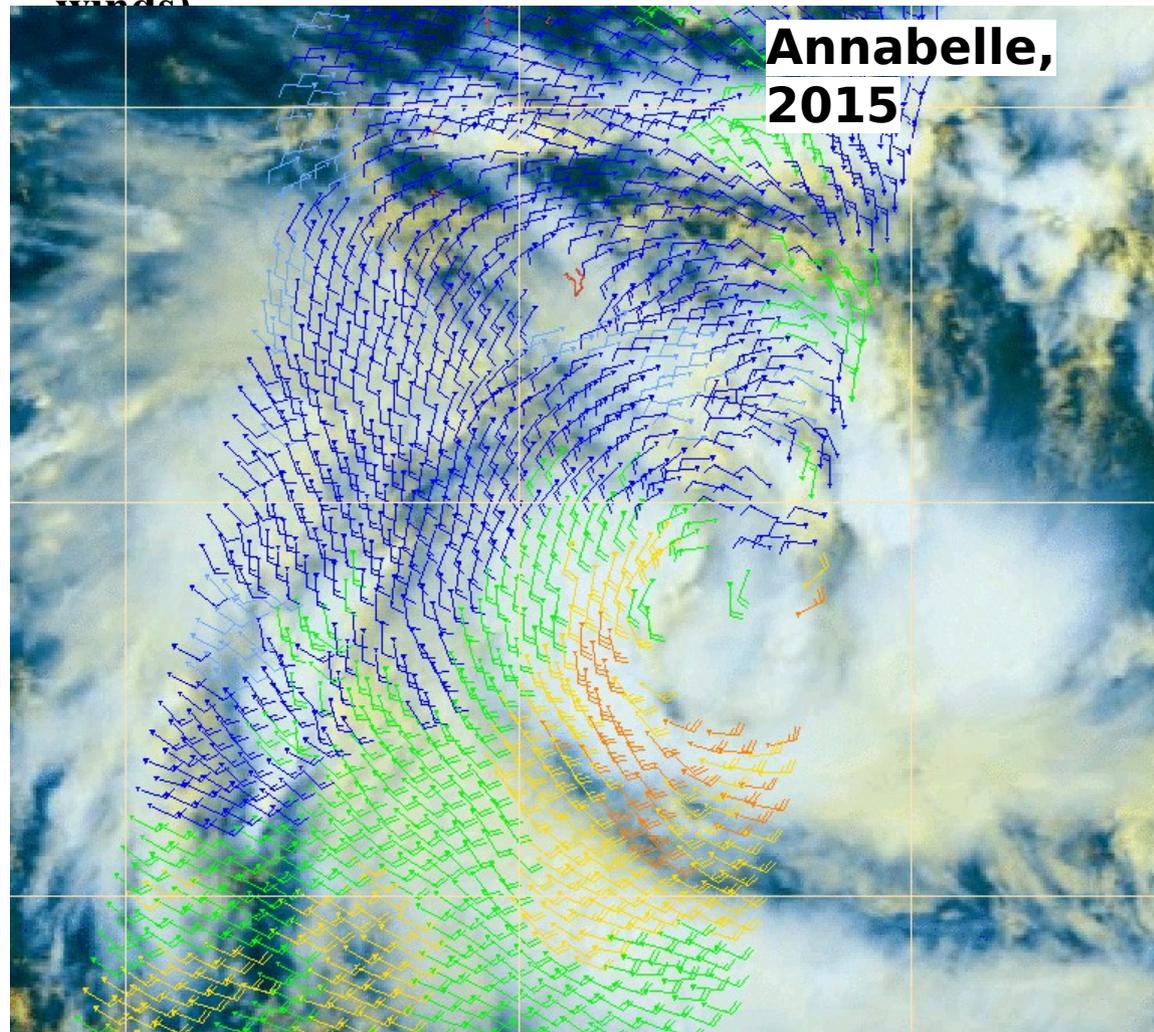
Equatorial (« westerly surge » = monsoon) and/or **tradewinds side** (strengthening trade

winds)

**Berguitta,
2018**



**Annabelle,
2015**

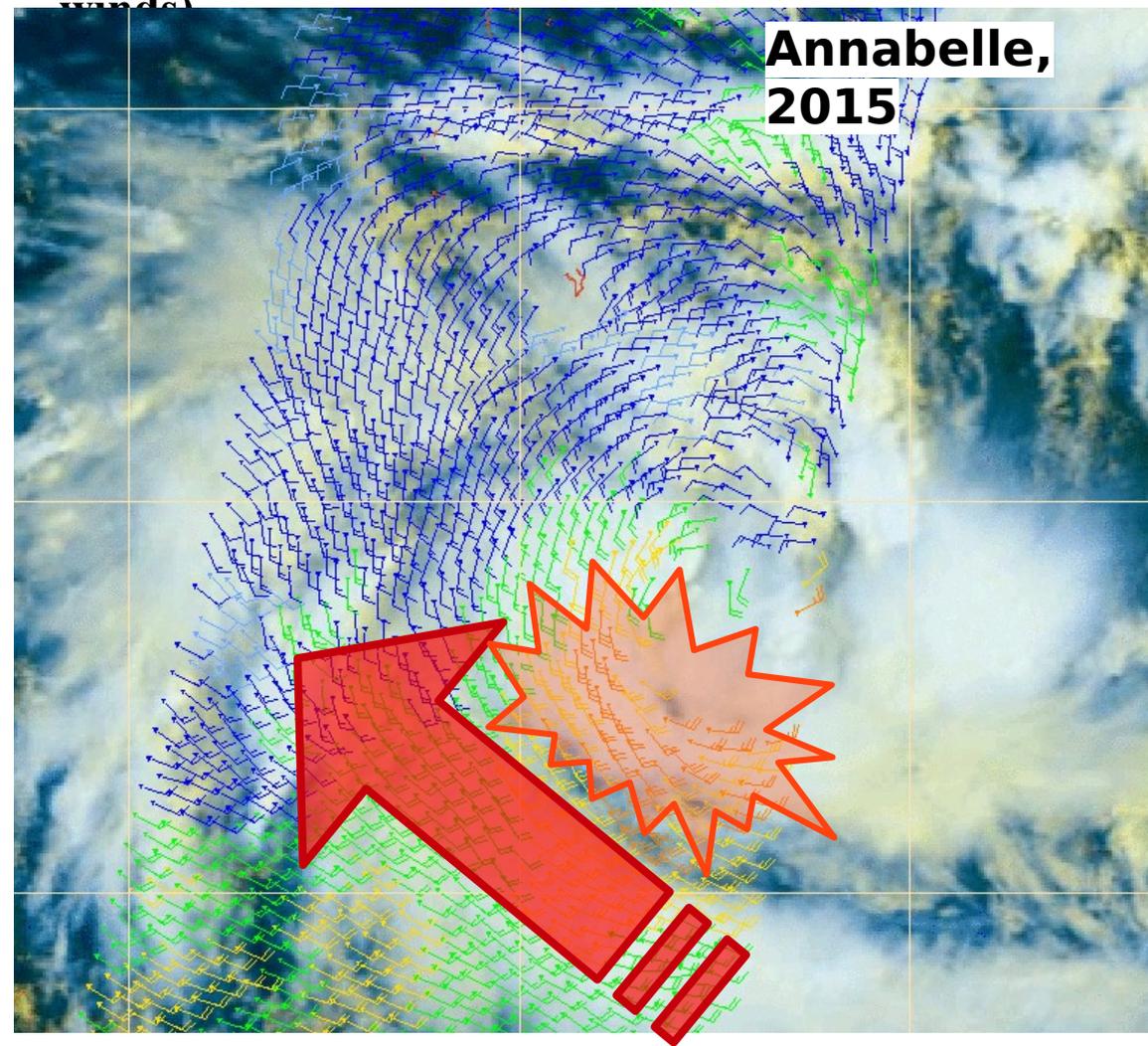
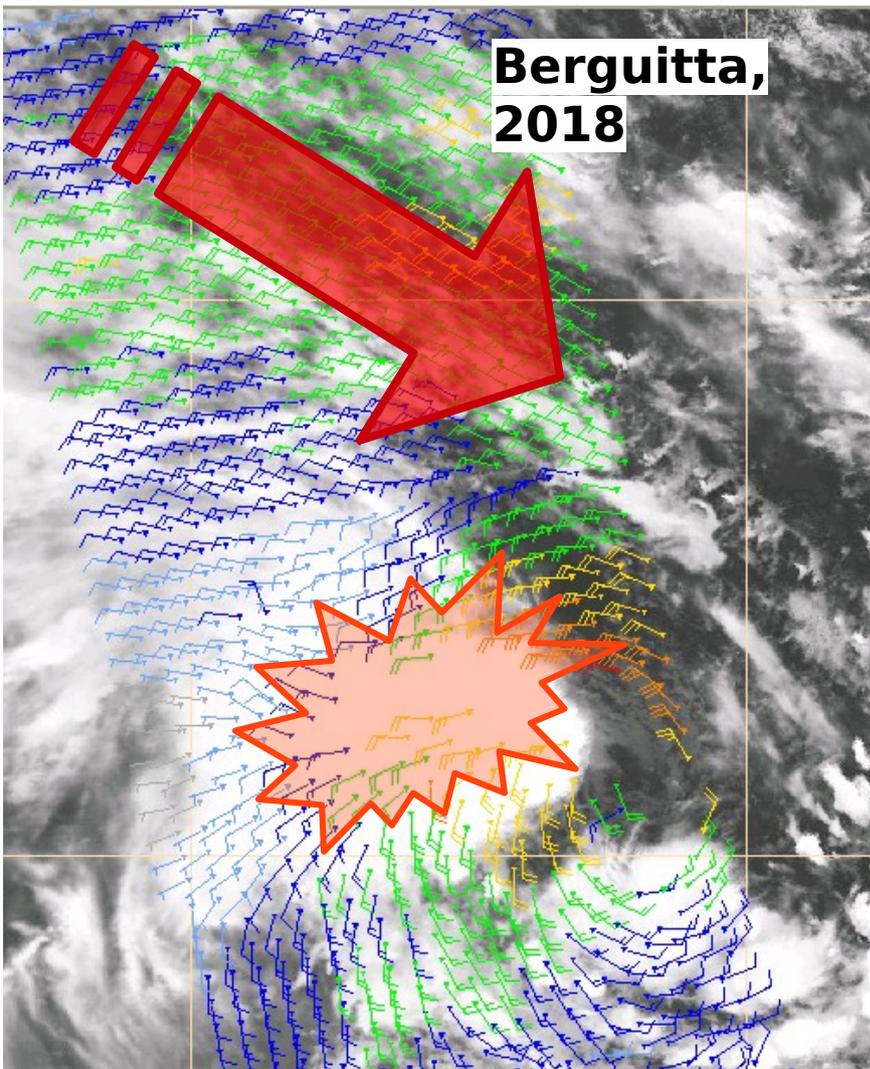


- **Low levels convergence enhancement: the spark that starts the cyclogenesis**

These wind surges can occur on both sides of the traffic:

Equatorial (« westerly surge » = monsoon) and/or **tradewinds side** (strengthening trade

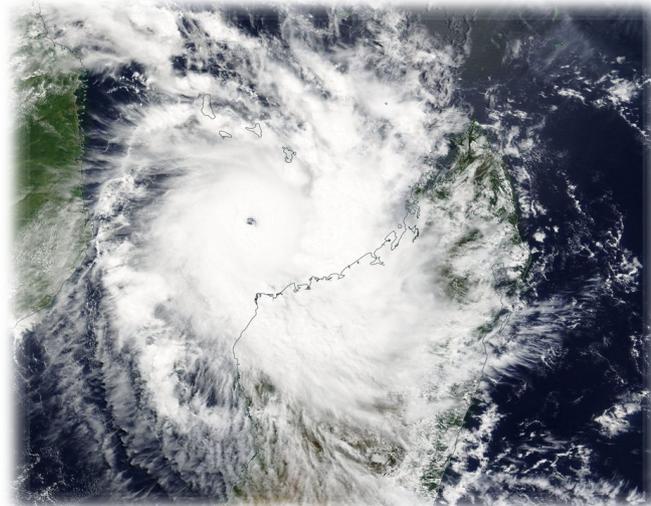
winds)



- **Basin Configuration**
Conclusive remarks

→ The preferred basin configurations in the SWIO basin during the warm season (November to April) are the Monsoon Trough (mid-season) and the Near Equatorial Trough (early/late season)

→ In the SWIO basin, a MT or NET configuration is, in the vast majority of cases, a **prerequisite for cyclogenesis** due to the strengthening of the low-level vorticity and the convergence of moisture



Tropical cyclogenesis

The role of equatorial waves



**METEO
FRANCE**



AI generated (Dall-E via Bing)

Equatorial waves

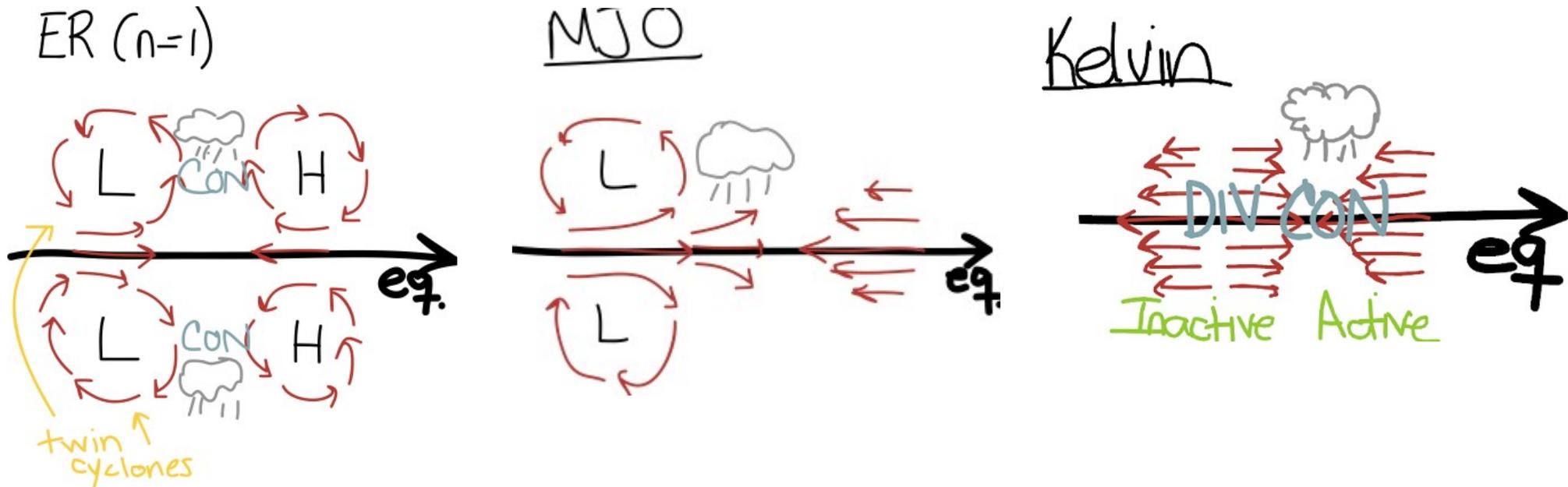
Definition : An equatorial wave materializes the **propagation of an atmospheric disturbance** on a planetary scale. It is **coupled to convection**: strong convective burst give rise to it and the propagation of the wave favors in turn convection. It remains channeled in the near equatorial zone ($\pm 15^\circ$) by the equatorial waveguide but also by the seasonal shift of the ITCZ.



Src : BOM YouTube

Equatorial waves

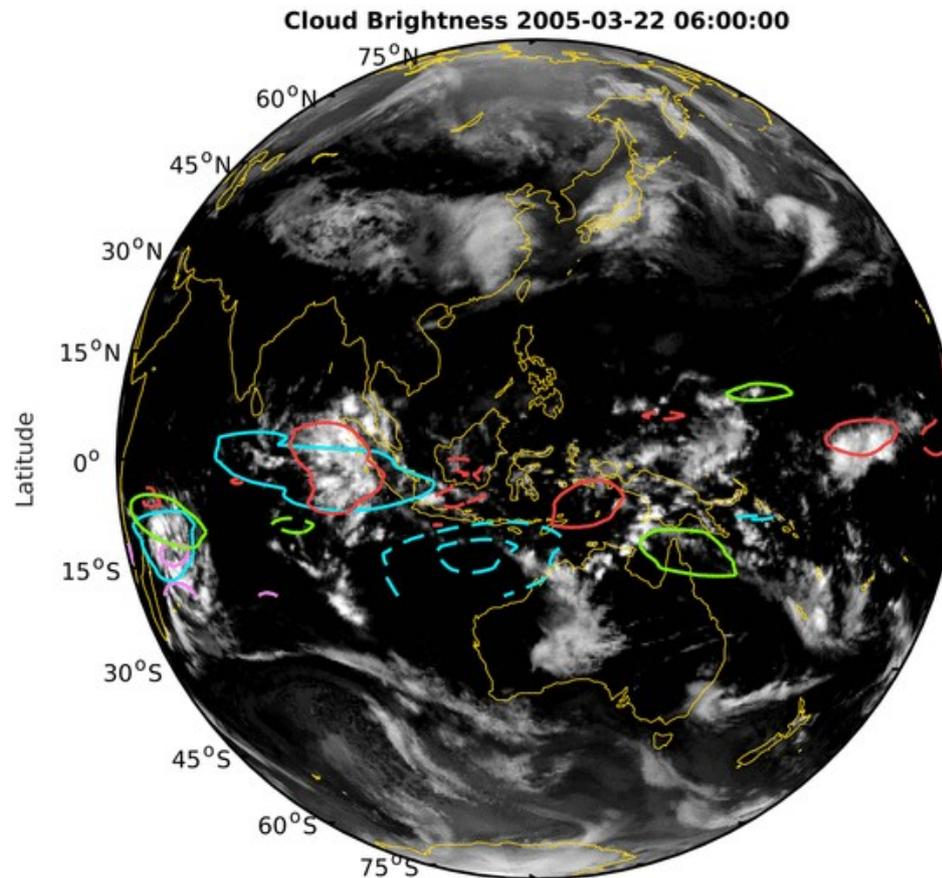
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Equatorial waves bring predictability on an intra-seasonal (monthly) scale.

Equatorial waves

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MJO

Kelvin waves
Kelvin waves in the
MJO envelope.

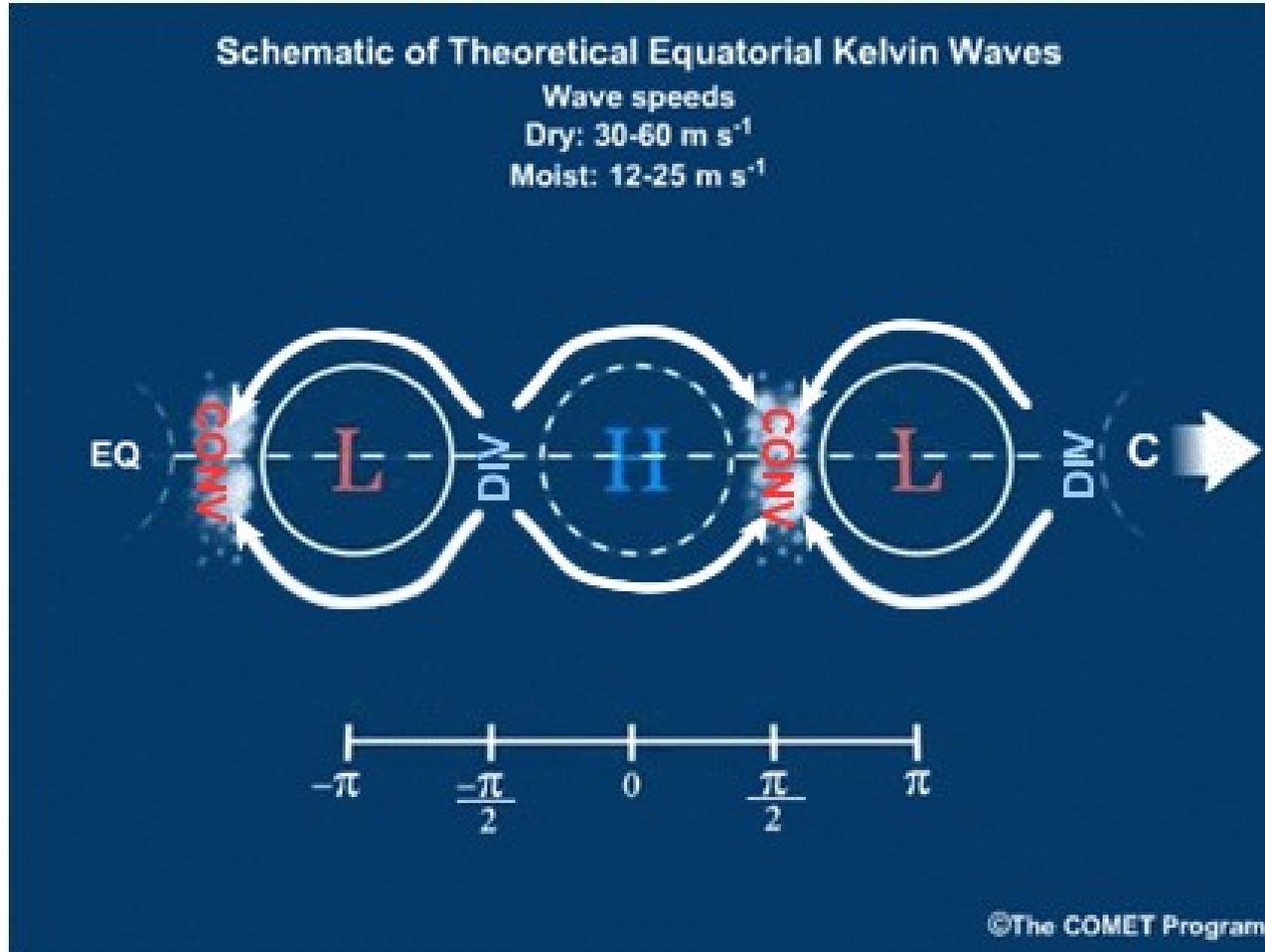
Equatorial Rossby waves

Kelvin waves

Propagation : Eastwards

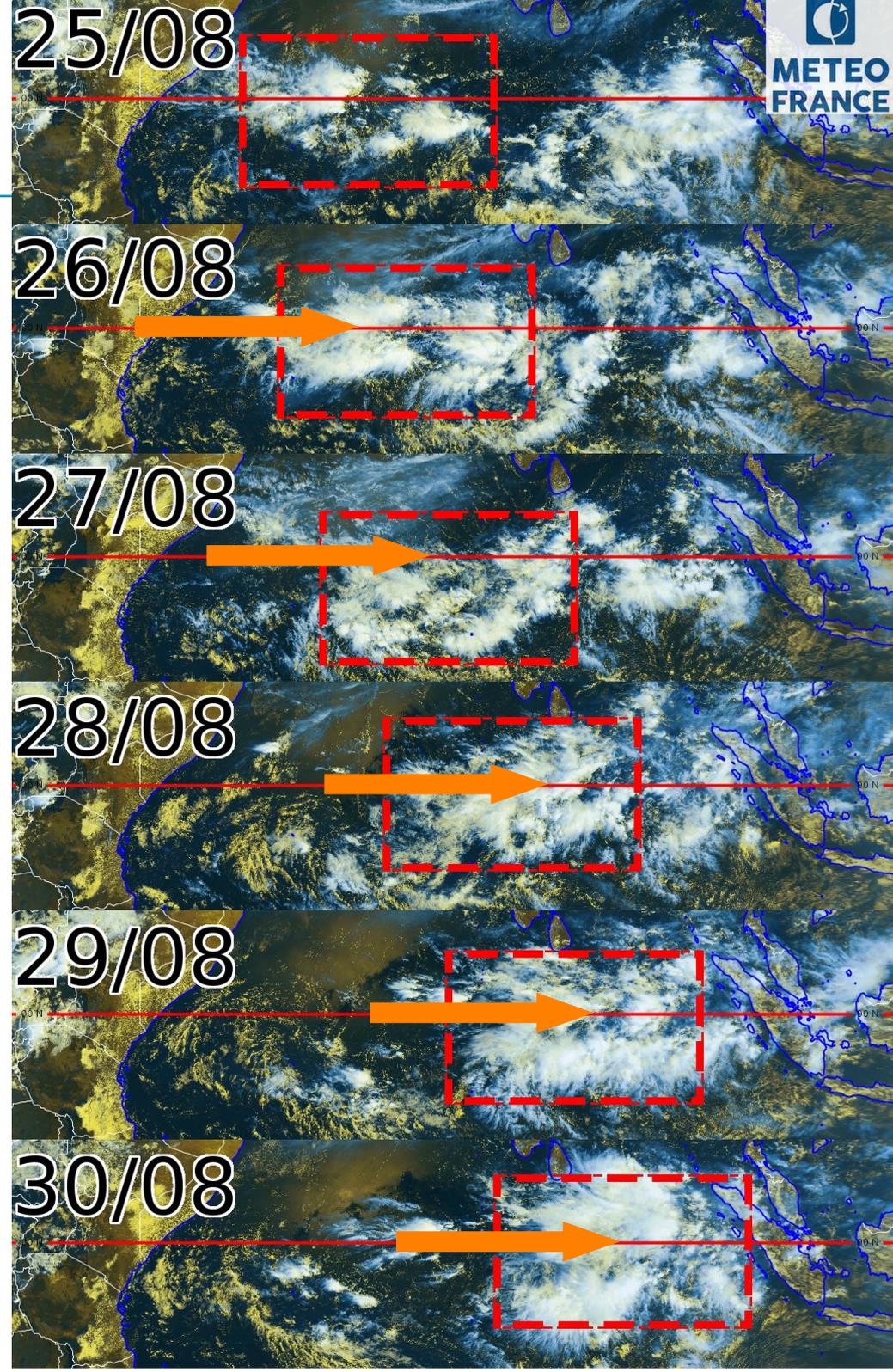
Period $\approx 3/7$ days

Consequences : enhanced convection ahead of a westerly wind surge



Kelvin wave

Propagation of a Kelvin wave over the Indian ocean

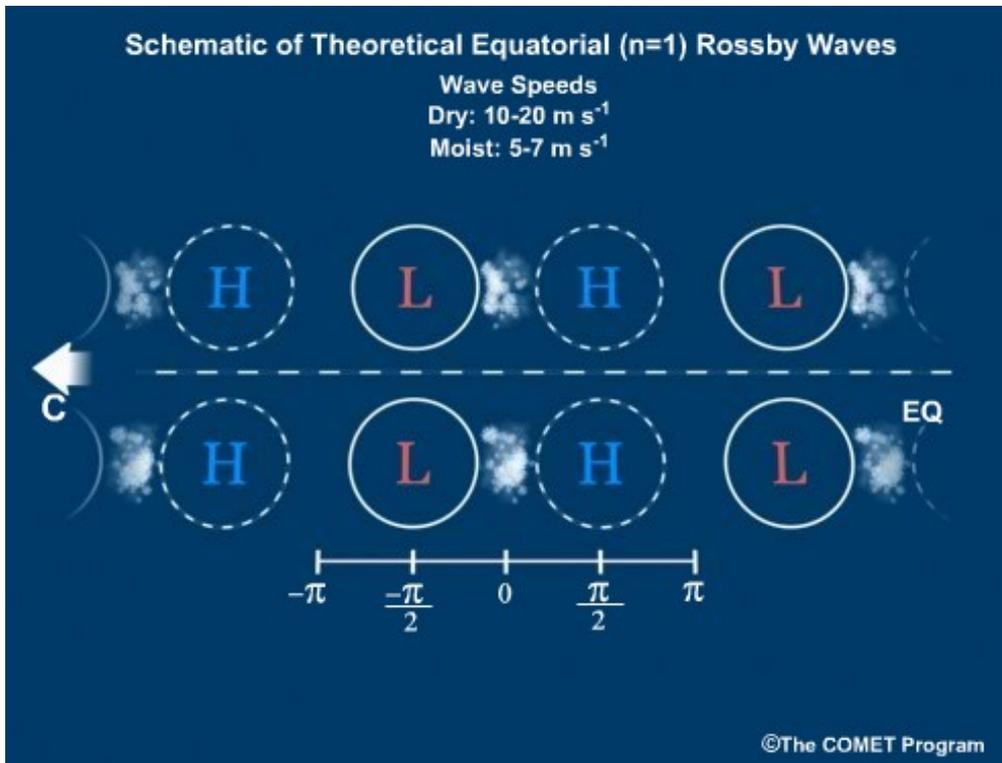


Equatorial Rossby (ER)

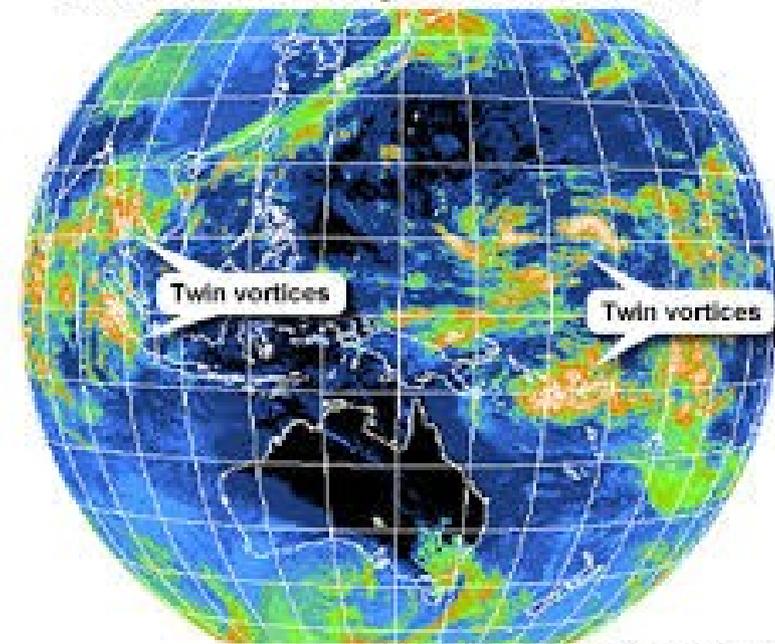
Propagation : **Westwards**

Period \approx **10/20 days**

Consequences : **Symmetrical cyclonic vortexes on both sides of the meteorological equator (varies with the season)**



Enhanced IR Satellite Image at 0000 UTC 7 Oct 2002



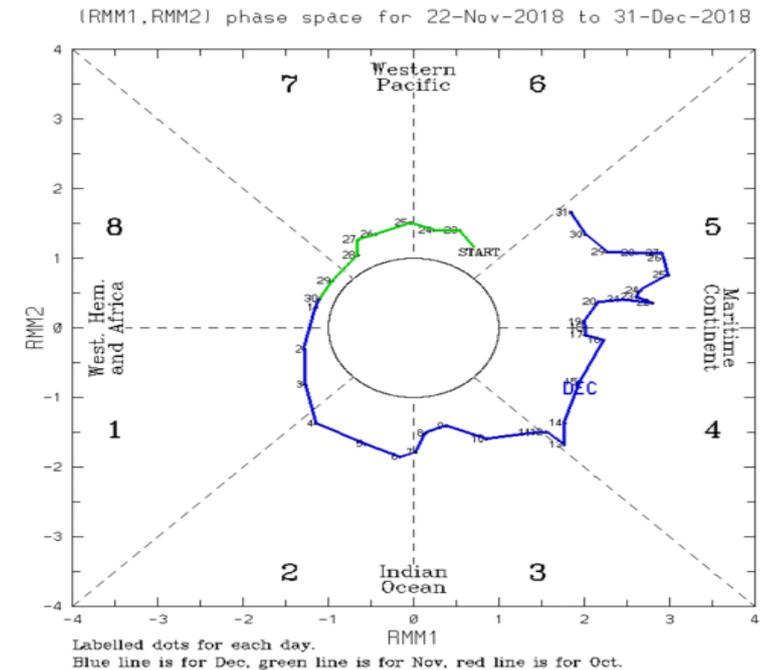
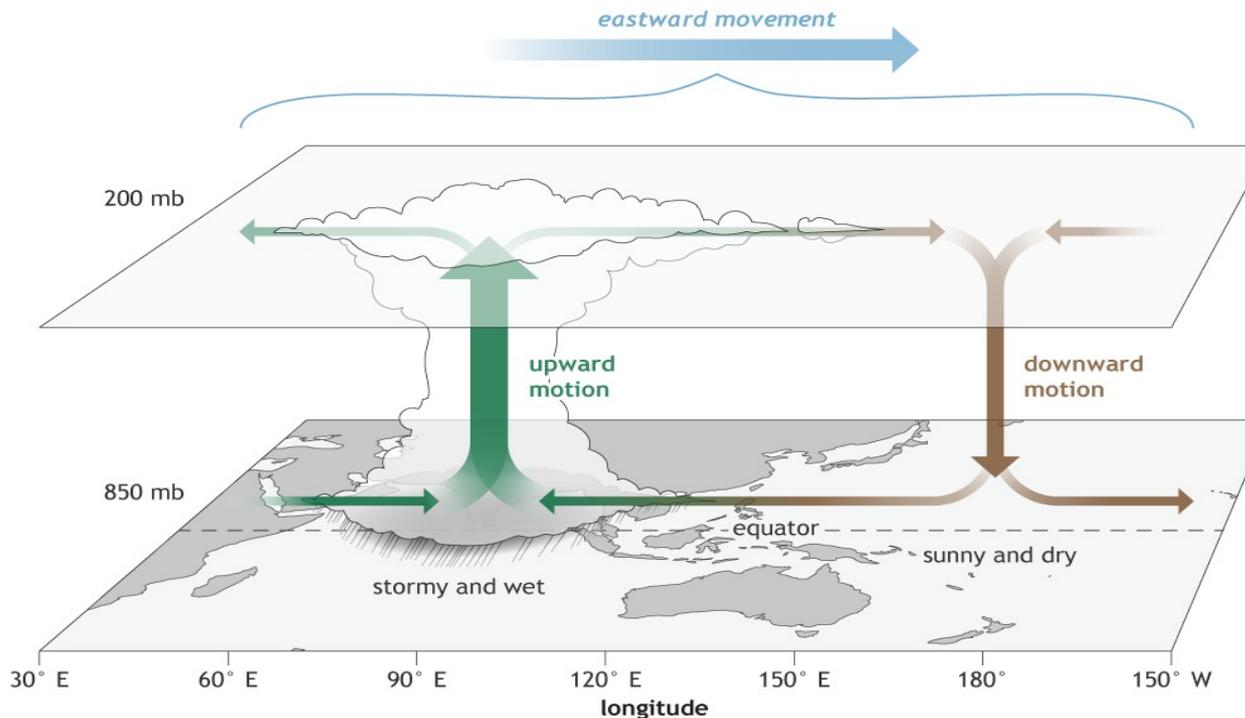
Australian Bureau of Meteorology / JMA

Madden-Julian Oscillation (MJO)

Propagation : Eastwards

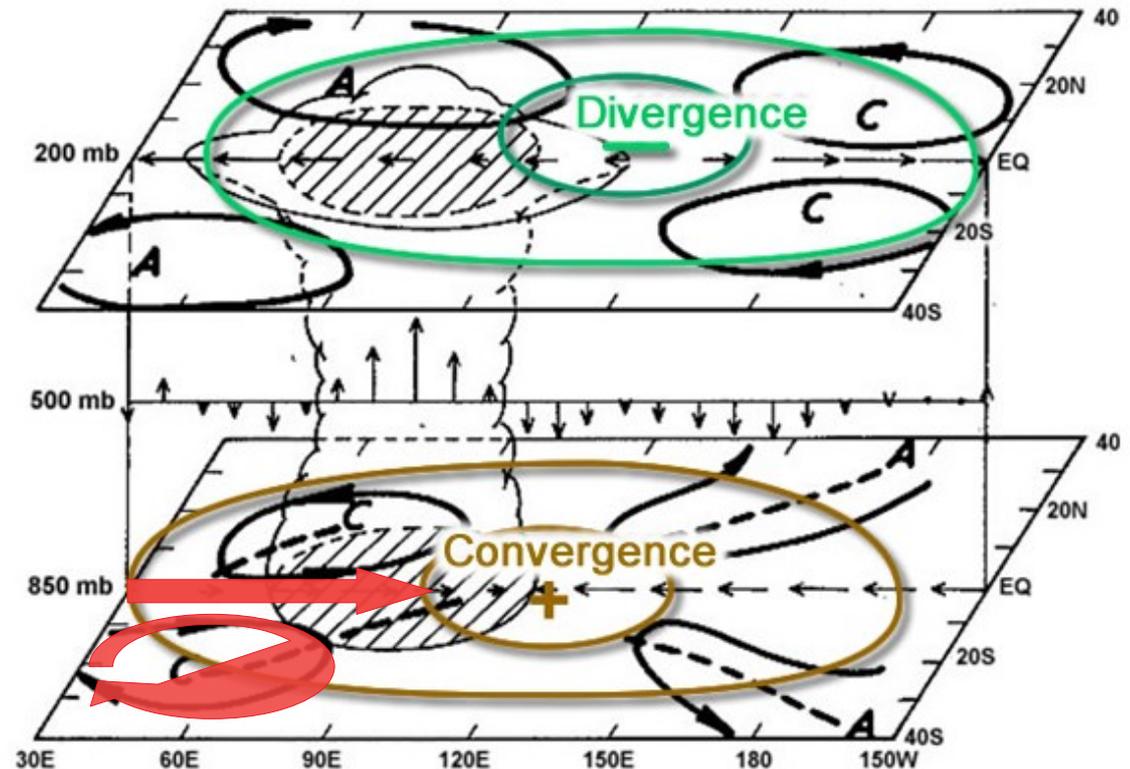
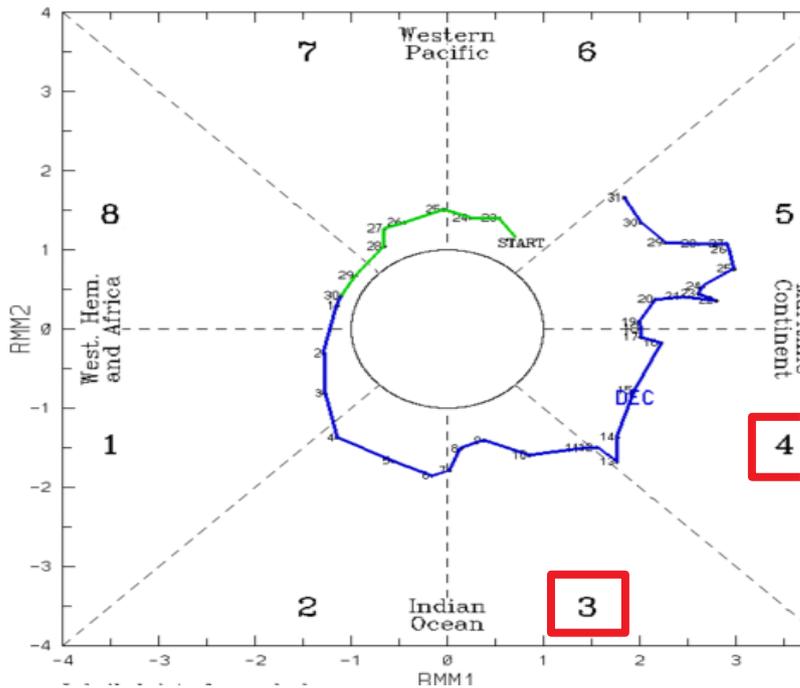
Period > 30 jours

Consequences : Succession of enhanced / suppressed large scale deep convection over the Indo-Pacific area within slow moving Walker cells. Modulating influence on tropical cyclogenesis at intra-seasonal time-scale for this region.



Madden-Julian Oscillation (MJO)

Schematic Depiction of the Large-scale Wind Structure of the MJO



Rui and Wang 1990

Over the SWIO, cyclogenesis often occurs at the back of the active MJO phase, when the RMM index is in position 3 and 4.

Tropical Waves influence on TCG

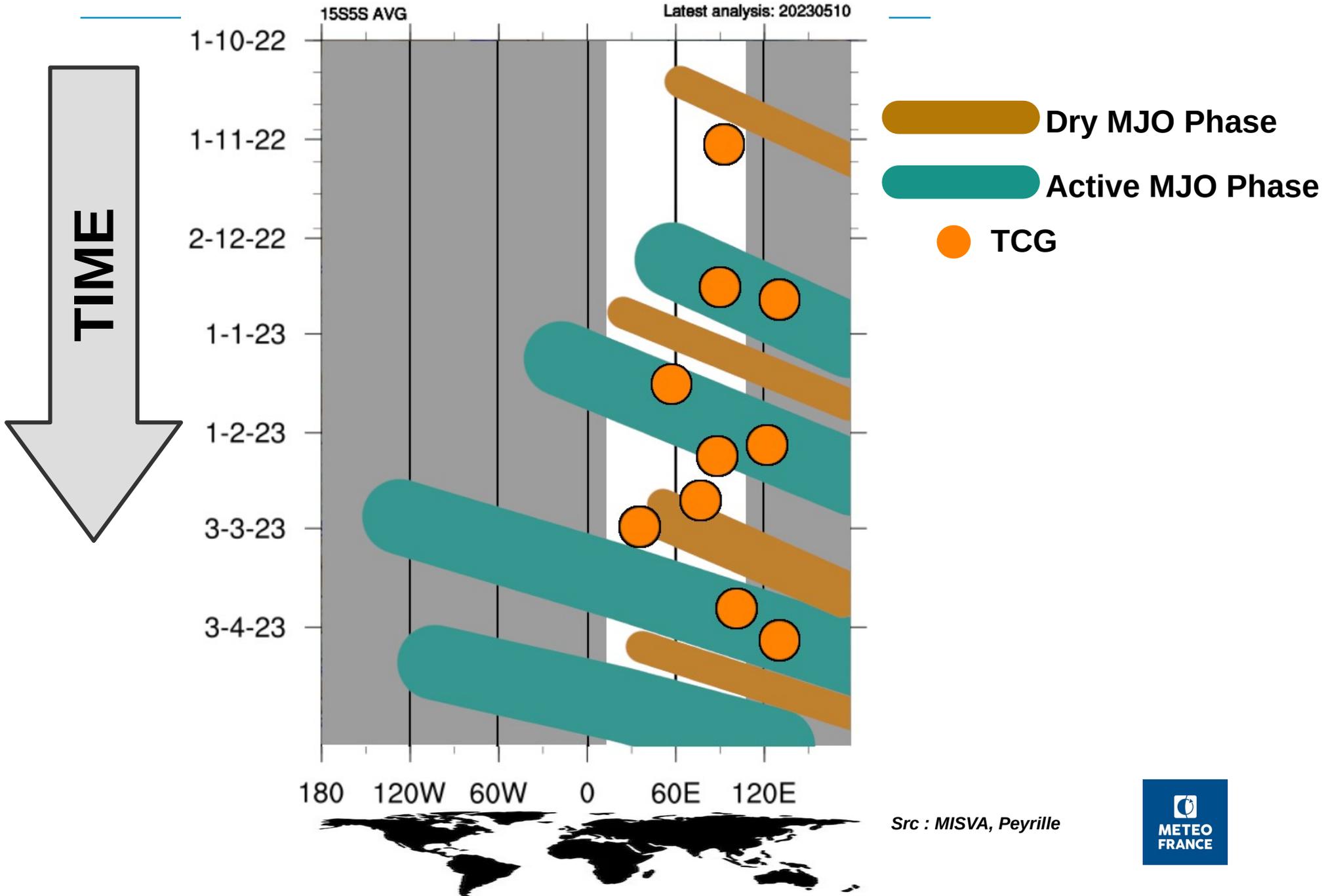
Tropical Waves favor TCG by :

→ Increasing deep convection and creating conducive conditions

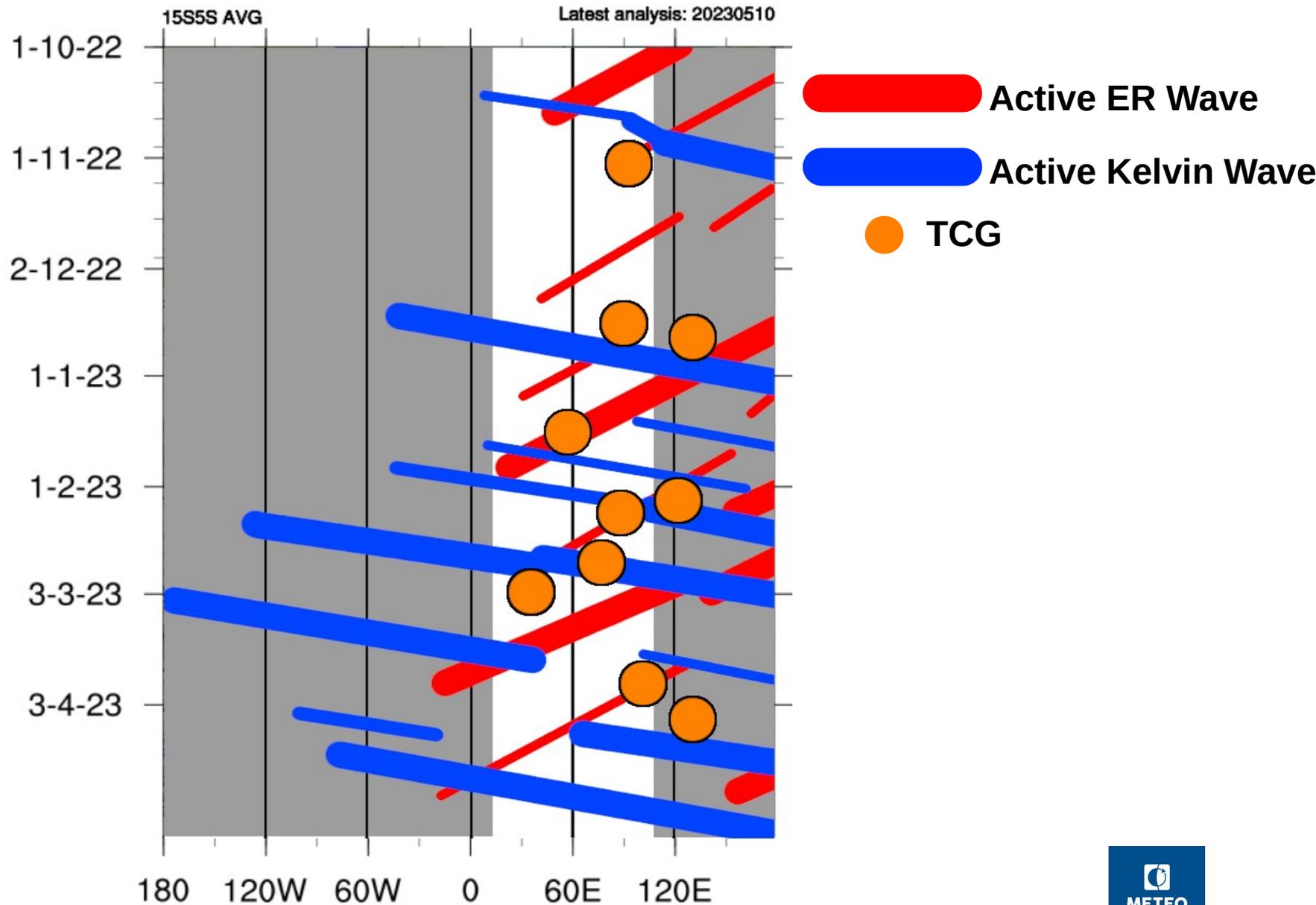
→ Creating background vorticity that decreases Rossby's radius of deformation

It is especially true when two waves cross path, typically Eastward moving Kelvin or Mjo with a westward moving Equatorial Rossby

Remember last summer...

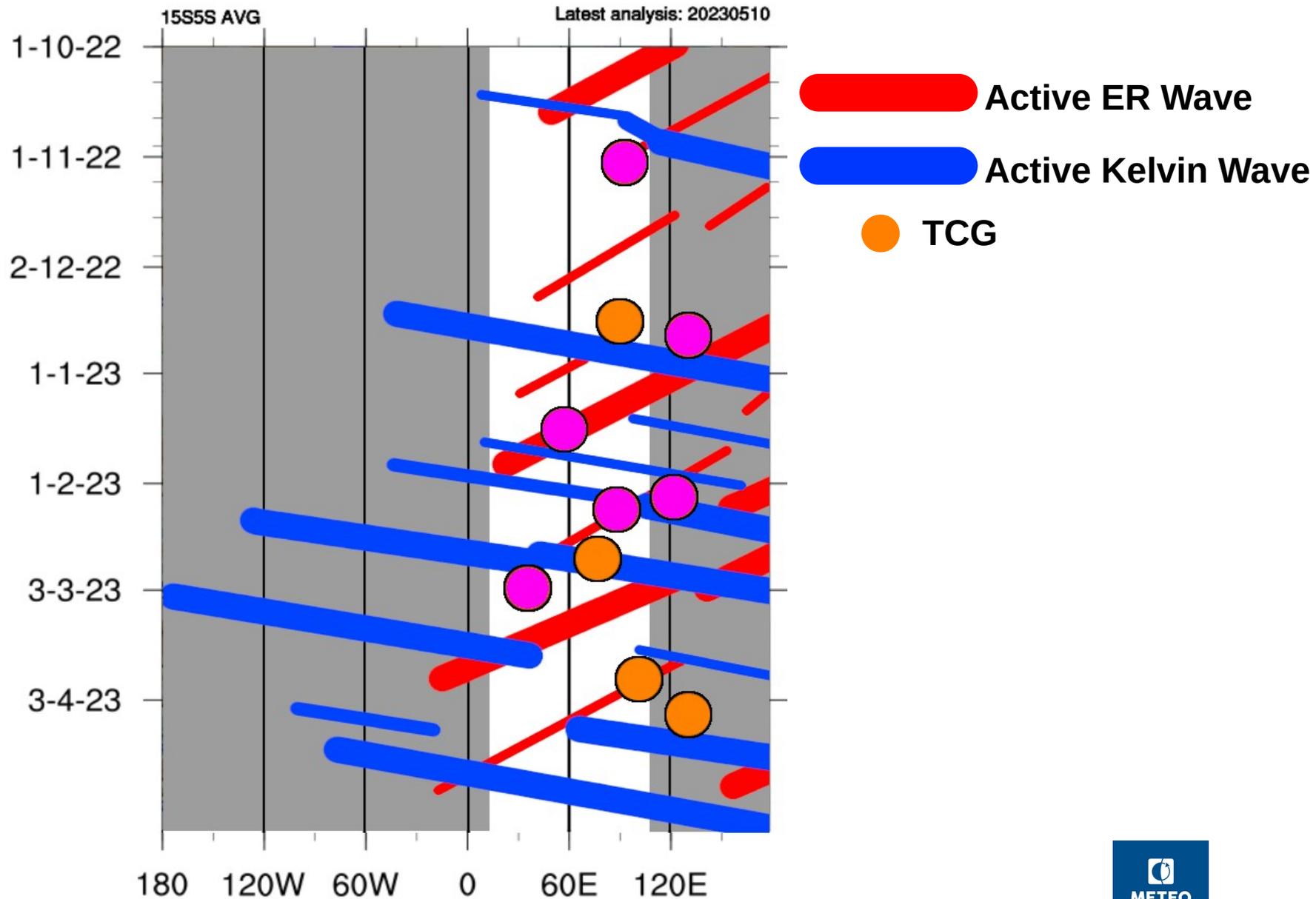


Remember last summer...



Src : MISVA, Peyrille

Remember last summer...



Src : MISVA, Peyrille

Tropical cyclogenesis

A few advanced notions



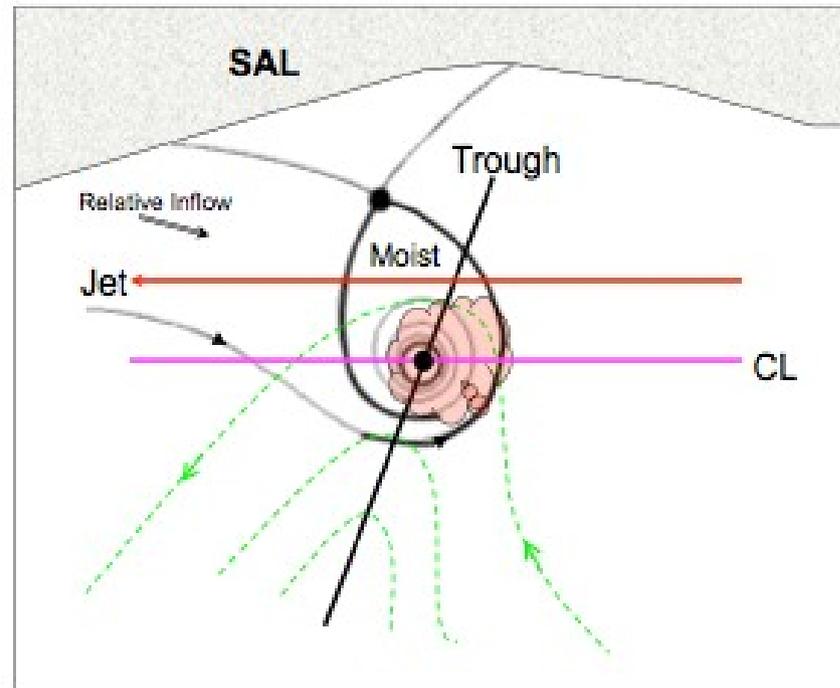
**METEO
FRANCE**



AI generated (Dall-E via Bing)

The Marsupial Paradigm

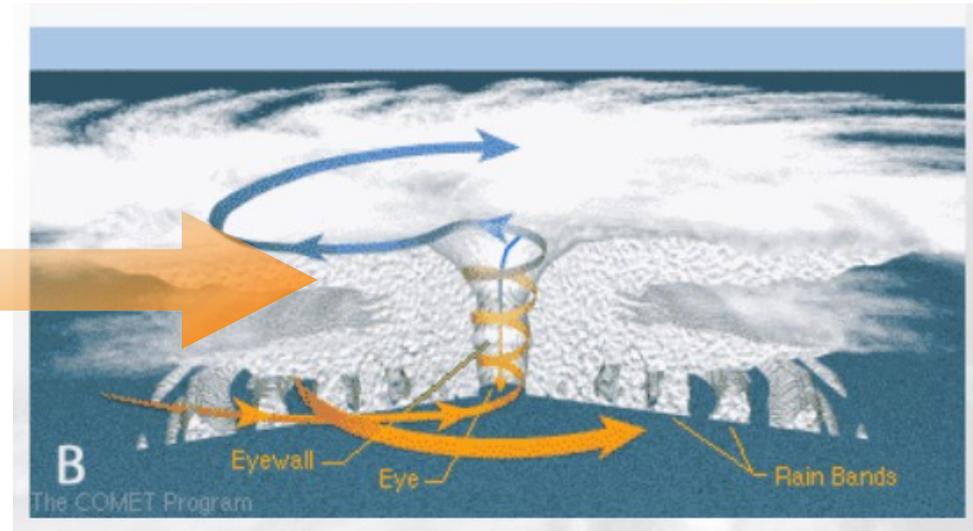
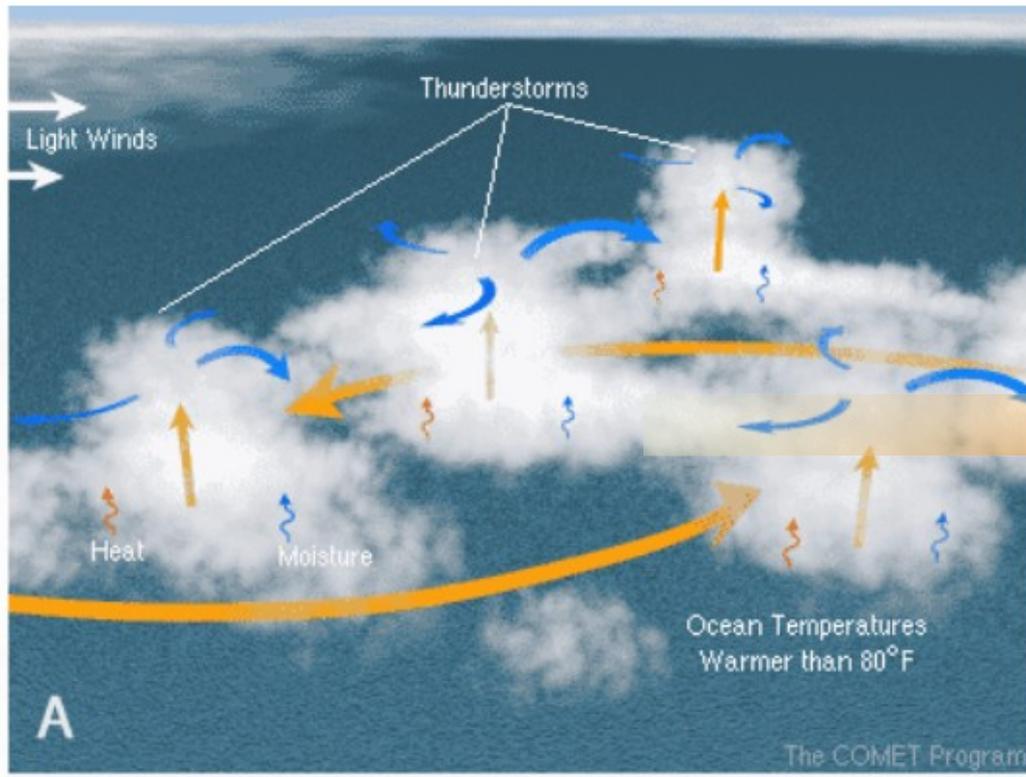
The areas of vorticities associated with Tropical Waves are thought to form a protective « pouch » of conducive environment, with little to no exchange with the outside.



From Wang et al. 2010

- Air inside of the pouch is continually moistened by convective bursts.
- Pouch offers protection from dry air intrusions.
- Originally created from Easterly waves over North Atlantic
- Currently being generalized to other basins and types of waves (**ER**)

In reality, pre-TC systems are not really symmetric...



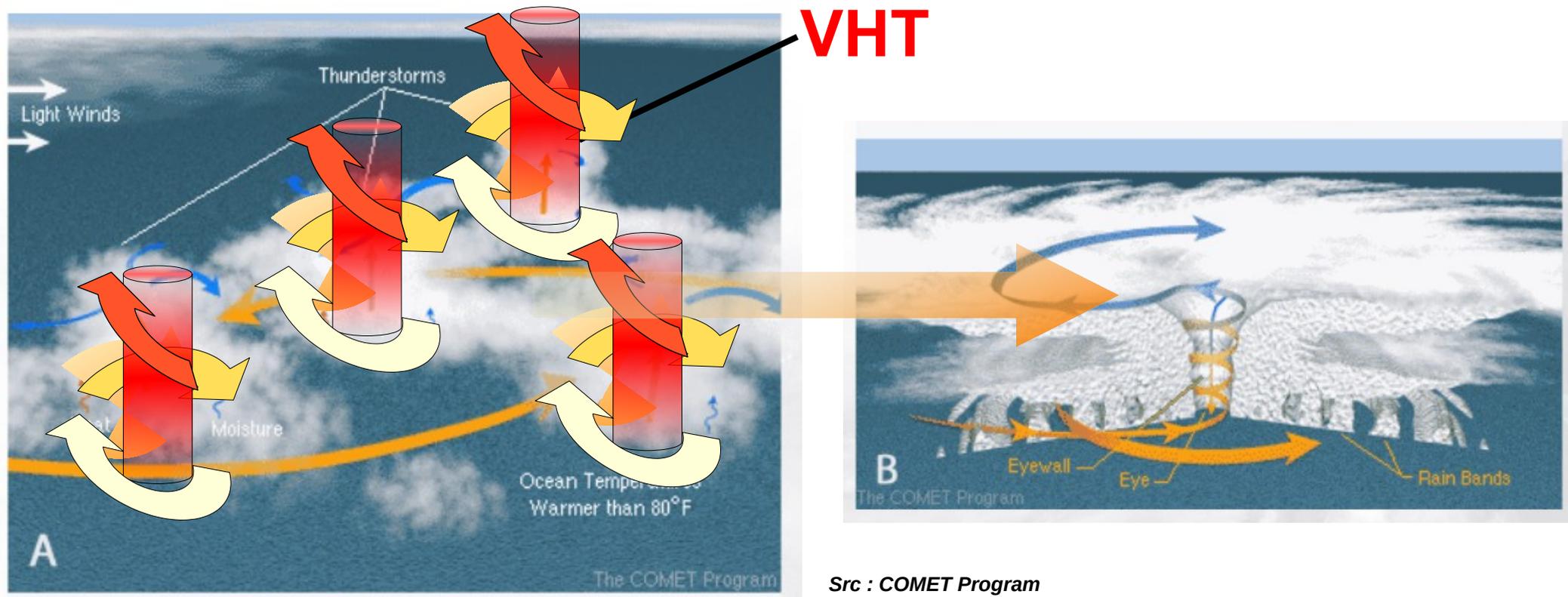
Src : COMET Program

Vortical Hot Towers (VHTs) develop inside the pouch circulation. They are the ‘building blocks’ of the future TC vortex, as they gradually merge and migrate towards the center.

This VHT pathway first formalized by Montgomery et al, 2006 appears to be :

- **Closer to real-world TCG with a 3D framework**
- **Taking into account asymmetries observed in real-life TCG scenarios**

How do we go from a few MCS to a self-sustained TS ?



Src : COMET Program

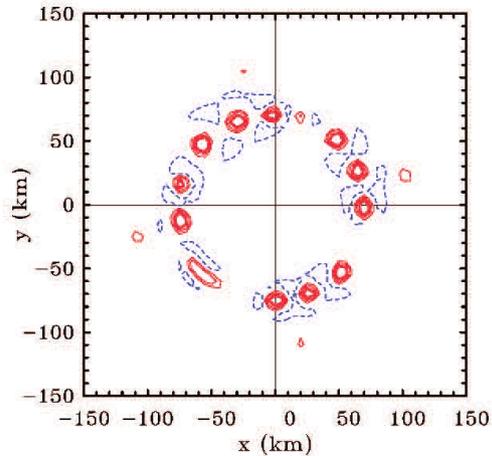
Vortical Hot Towers (VHTs) develop inside the precursor circulation. They are the 'building blocks' of the future TC vortex, as they gradually merge and migrate towards the center.

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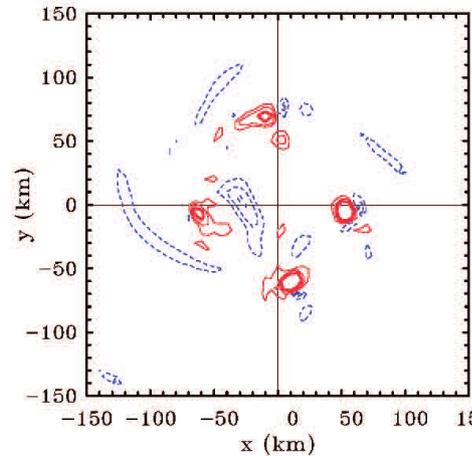
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How do we go from a few MCS to a self-sustained TS ?

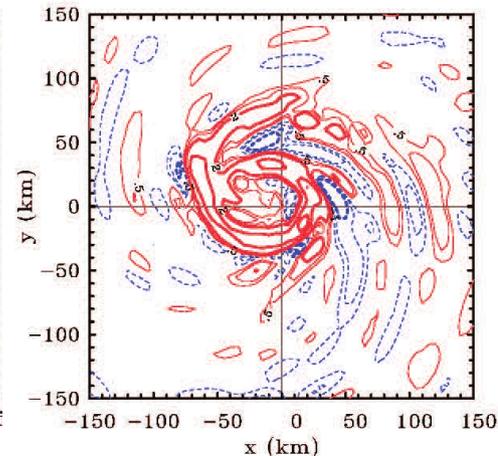
9,75 hrs



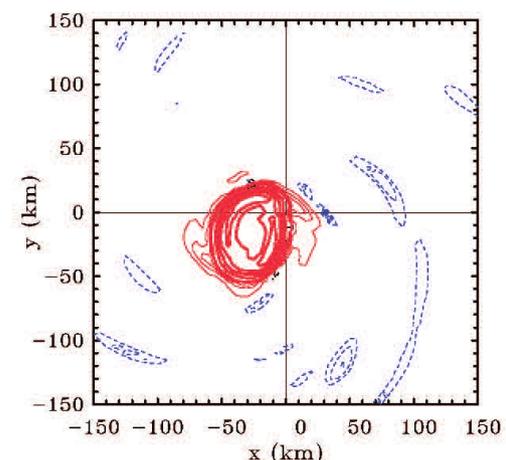
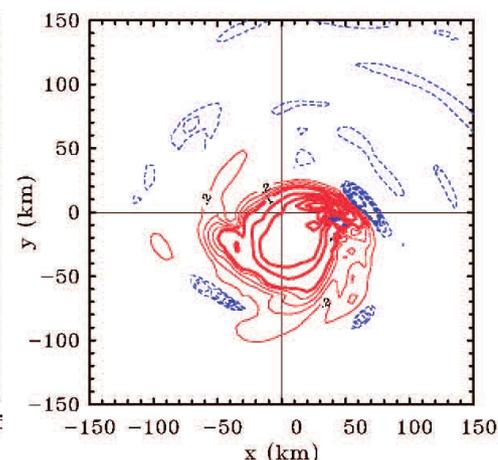
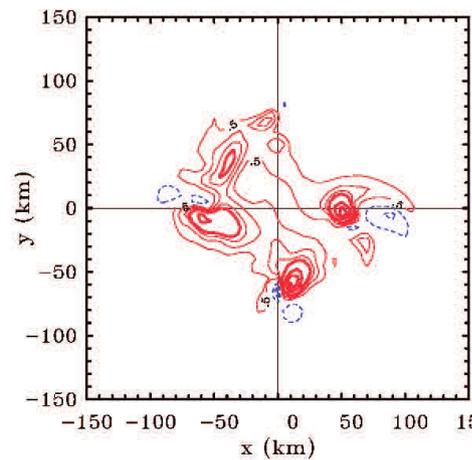
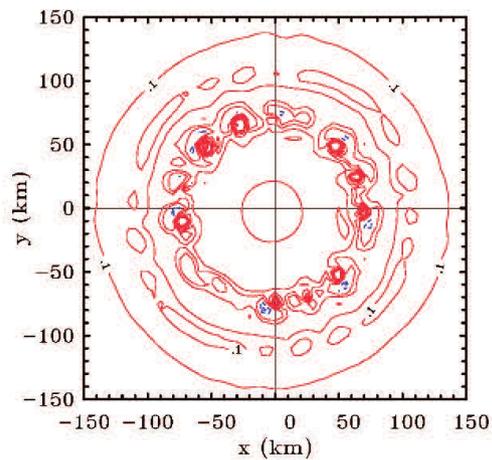
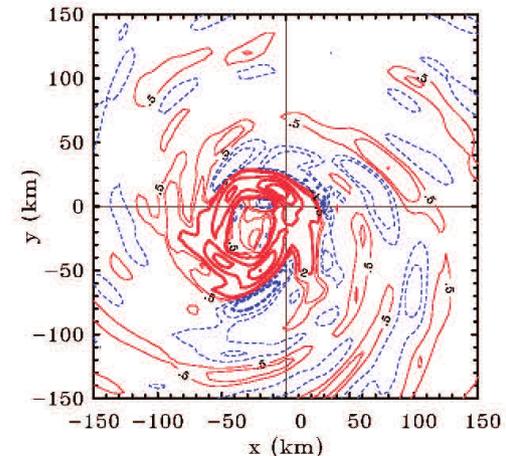
24 hrs



48 hrs



96 hrs

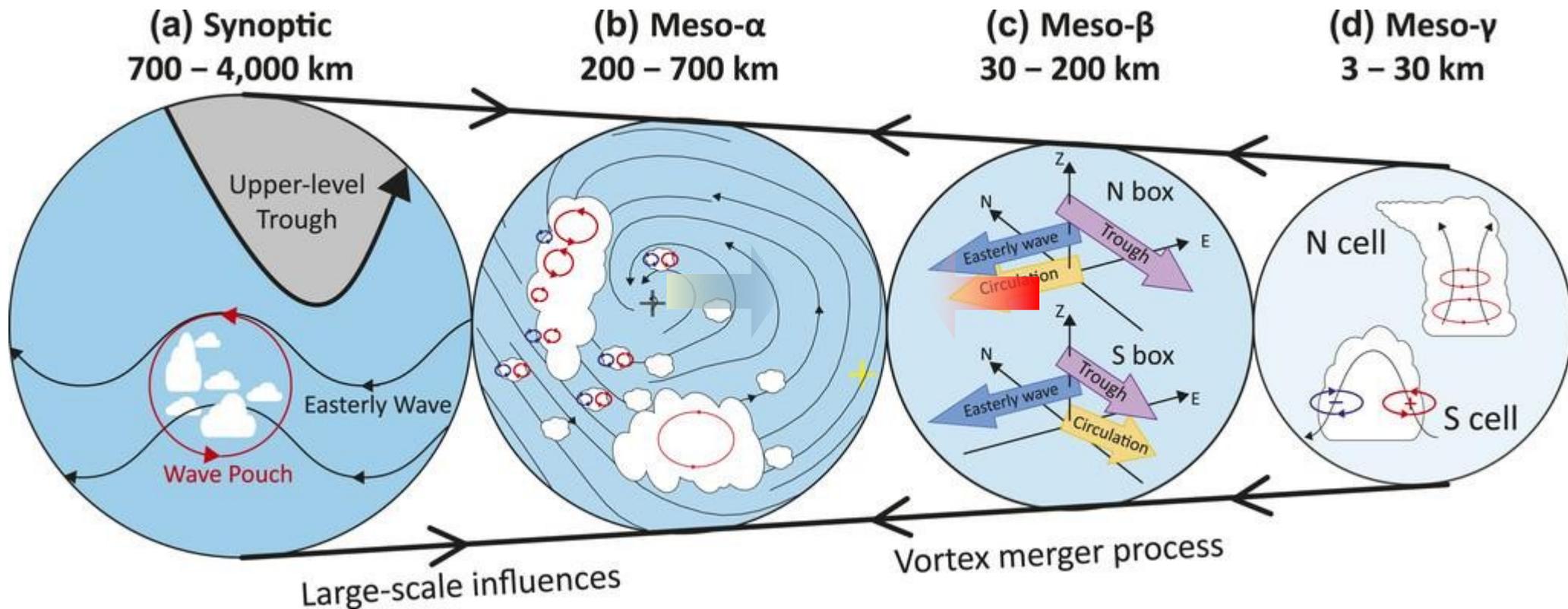


Figures from Montgomery,

2014

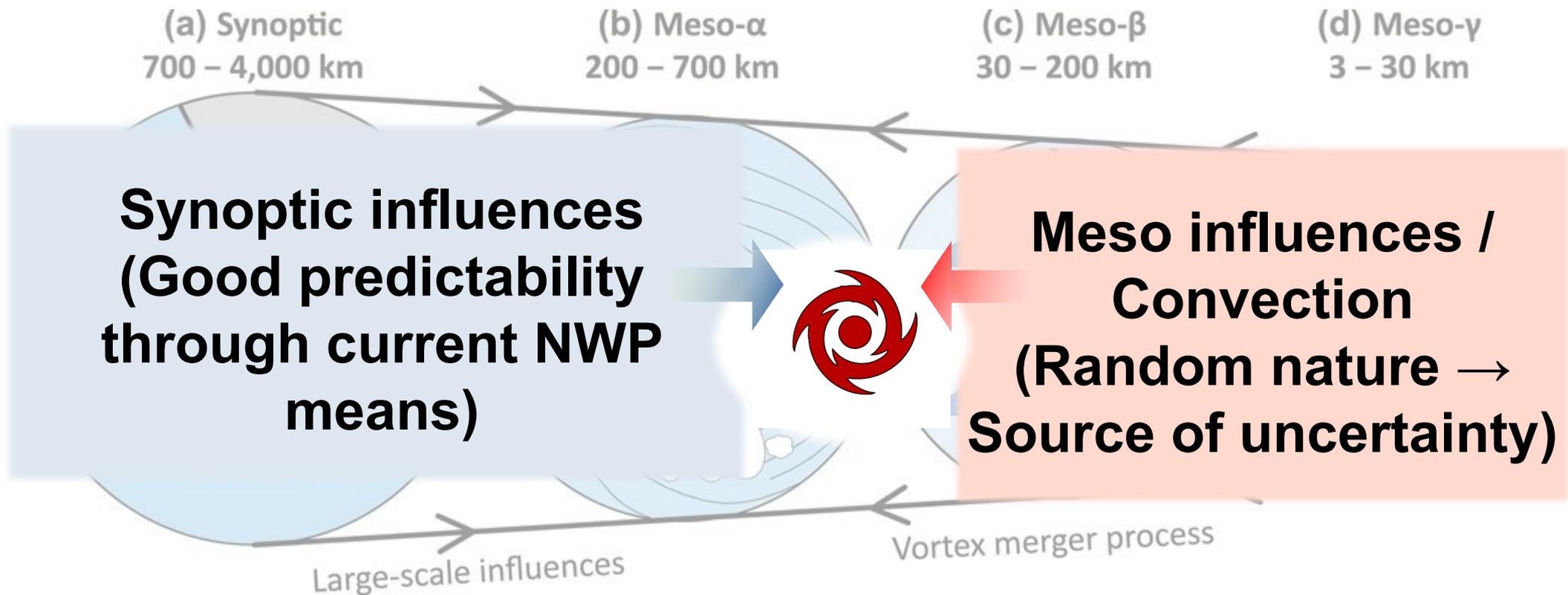
Vertical velocities (850 hPa): top line
Vorticity (850 hPa) : bottom line

How do we go from a few MCS to a self-sustained TS ?



Adapted from Nam and Bell (2021).

How do we go from a few MCS to a self-sustained TS ?



Adapted from Nam and Bell (2021).

Take away messages

- Cyclogenesis corresponds to the **initial intensification phase** of a tropical low pressure system.
- Once initiated, the **positive feedback loop** between the release of latent heat by deep convection, lowering of central pressure and strengthening of surface winds, does not need input from the environment to continue. **TC becomes autonomous.**
- The presence of a **well defined vorticity precursor** in the low-levels is key in the default pathway to cyclogenesis over the SWIO.
- **Specific basin configurations and tropical waves** influence TCG by creating precursors and promoting deep convection.
- Research is still under way on the subject of TCG on various aspects of the problem : interactions between waves, cloud microphysics.



AI generated (Dall-E via Bing)

